Chapter 2: Terrestrial and Freshwater Ecosystems and their Services

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Executive Summary

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Chapter 2, building upon prior assessments¹ provides a global assessment of observed impacts and projected
 risks of climate change to terrestrial and freshwater ecosystems, including their component species and the
 services they provide to people. Where possible, differences among regions, taxonomic groups and
 ecosystems types are presented. Adaptation options to reduce risks to ecosystems and people are assessed.

Observed Impacts

Multiple lines of evidence, combined with strong and consistent trends observed on every continent, 10 make it very likely² that many observed changes in ranges, phenology, physiology and morphology of 11 terrestrial and freshwater species can be attributed to regional and global climate changes, 12 particularly increases in frequency and severity of extreme events (very high confidence³) {2.3.1; 13 2.3.3.5; 2.4.2; 2.4.5; Table 2.2; Table 2.3; Table 2.S.1; Cross-Chapter Box EXTREMES this Chapter }. The 14 most severe impacts are occurring in the most vulnerable species and ecosystems, characterized by inherent 15 physiological, ecological or behavioural traits that limit their abilities to adapt and those most exposed to 16 climatic hazards (high confidence) {2.4.2.2; 2.4.2.6; 2.4.2.8; 2.4.5; 2.6.1; Cross-Chapter Box EXTREMES 17 this Chapter}. 18 19

New studies since AR5 and SR1.5 (now >12,000 species globally) show changes consistent with climate-20 change. Where attribution was assessed (>4,000 species globally) approximately half of species had 21 shifted their ranges to higher latitudes or elevations and two-thirds of spring phenology had advanced, 22 driven by regional climate changes (very high confidence). Shifts in species ranges are altering 23 community make-up, with exotic species exhibiting greater ability to adapt to climate change than natives, 24 especially in more northern latitudes, potentially leading to newly invasive species {2.4.2.3.3}. New 25 analyses demonstrate that prior reports underestimated impacts due to complex biological responses to 26 climate change (high confidence). {2.4.2.1; 2.4.2.3; 2.4.2.4; 2.4.2.5; 2.4.5; Table 2.2; Table SM2.1; Table 27 2.3} 28

29 Responses in freshwater species are strongly related to changes in the physical environment (high 30 confidence) {2.3.3; 2.4.2.3.2}. Global coverage of quantitative observations in freshwater ecosystems has 31 increased since AR5. Water temperature has increased in rivers (up to 1°C decade⁻¹) and lakes (up to 0.45°C 32 decade⁻¹) $\{2.3.3.1;$ Figure 2.2 $\}$. Extent of ice cover has declined by 25% and duration by >2 weeks $\{2.3.3.4;$ 33 Figure 2.4}. Changes in flow have led to reduced connectivity in rivers (high confidence) {2.3.3.2; Figure 34 2.3. Indirect changes include alterations in river morphology, substrate composition, oxygen concentrations 35 and thermal regime in lakes (very high confidence) {2.3.3.2; 2.3.3.3}. Dissolved oxygen concentrations have 36 typically declined and primary productivity increased with warming. Warming and browning (increase in 37 organic matter) have occurred in boreal freshwaters with both positive and negative repercussions on water 38 temperature profiles (lower vs. upper water)(high confidence) and primary productivity (medium confidence) 39 and reduced water quality (high confidence) {2.4.4.1; Figure 2.5}. 40

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¹ Previous IPCC assessments include the IPCC Fifth Assessment Report (AR5) (IPCC, 2013; IPCC, 2014c; IPCC, 2014d; IPCC, 2014a), the Special Report on Global Warming of 1.5°C (SR1.5) (IPCC, 2014b), the Special Report on Ocean and Cryosphere in a Changing Climate (SROCC) (IPCC, 2019b) and the IPCC Sixth Assessment Report Working Group I (AR6 WGI).

² In this Report, the following terms have been used to indicate the assessed likelihood of an outcome or a result: Virtually certain 99–100% probability, Very likely 90–100%, Likely 66–100%, About as likely as not 33–66%, Unlikely 0–33%, Very unlikely 0–10%, and Exceptionally unlikely 0–1%. Additional terms (Extremely likely: 95– 100%, More likely than not >50–100%, and Extremely unlikely 0–5%) may also be used when appropriate. Assessed likelihood is typeset in italics, e.g., *very likely*). This Report also uses the term '*likely* range' to indicate that the assessed likelihood of an outcome lies within the 17-83% probability range.

³ In this Report, the following summary terms are used to describe the available evidence: limited, medium, or robust; and for the degree of agreement: low, medium, or high. A level of confidence is expressed using five qualifiers: very low, low, medium, high, and very high, and typeset in italics, e.g., *medium confidence*. For a given evidence and agreement statement, different confidence levels can be assigned, but increasing levels of evidence and degrees of agreement are correlated with increasing confidence.

Chapter 2 FINAL DRAFT IPCC WGII Sixth Assessment Report Climate change has increased wildlife diseases (high confidence). Experimental studies provide high 1 confidence in attribution of observed increased disease severity, outbreak frequency and emergence of novel 2 vectors and their diseases into new areas to recent trends in climate and extreme events. Many vector-borne 3 diseases, and those caused by ticks, helminth worms and Bd (chytrid) fungus, have shifted poleward, 4 upward, and are emerging in new regions (high confidence). In the high Arctic and high elevations of Nepal, 5 there is high confidence that climate change has driven expansion of vector-borne diseases that infect 6 humans. {2.4.2.7, 7.2.2.1, 9.8.2.4, 10.4.7.1, 12.3.1.4, 13.7.1.2, 14.4.6.4, Cross-Chapter Box ILLNESS this 7 Chapter} 8 9 Forest insect pests have expanded northward and severity and outbreak extent has increased in 10 northern North America, northern Eurasia, due to warmer winters reducing mortality and longer 11 growing seasons favouring more generations per year (high confidence) {2.4.2.1} 12 13 Climate-caused local population extinctions have been widespread among plants and animals, detected 14 in 47% of 976 species examined and associated with increases in hottest yearly temperatures (very high 15 confidence) {2.4.2.2}. Climate-driven population extinctions have been higher in tropical (55%), than 16 temperate habitats (39%), higher in freshwater (74%), than in marine (51%) or terrestrial (46%) habitats and 17 higher in animals (50%) than in plants (39%). Extreme heat waves has led to local fish kills in lakes 18{2.3.3.5}. Intensification of droughts contributes to disappearance of small or ephemeral ponds, which often 19 hold rare and endemic species. {2.4.2.2; Cross-Chapter Box EXTREMES this Chapter} 20 21 Global extinctions or near-extinctions have been linked to regional climate change in three 22 documented cases {2.4.2.2}. The cloud-forest-restricted Golden toad (Incilius periglenes) was extinct by 23 1990 in a nature preserve in Costa Rica following successive extreme droughts (medium confidence). The 24 white sub-species of the lemuroid ringtail possum (Hemibelideus lemuroides) in Queensland, Australia, 25 disappeared after heatwaves in 2005 (high confidence): intensive censuses found only 2 individuals in 2009. 26 The Bramble Cays Melomys (Melomys rubicola), was not seen after 2009 and declared extinct in 2016, with 27 SLR and increased storm surge, associated with climate change, the most probable drivers (high confidence). 28 The interaction of climate change and chytrid fungus (Bd) has driven many of the observed global 29 amphibian declines and species' extinctions (robust evidence, high agreement) {2.4.2.7.1}. 30 31 A growing number of studies document genetic evolution within populations in response to recent climate 32 change (very high confidence). To date, genetic changes remain within the limits of known variation for 33 species (high confidence). Controlled selection experiments and field observations indicate that 34 evolution would not prevent a species becoming extinct if its climate space disappears globally (high 35 confidence). Climate hazards outside those to which species are adapted are occurring on all continents 36 (high confidence). More frequent and intense extreme events, superimposed on longer-term climate trends, 37 have pushed sensitive species and ecosystems towards tipping points, beyond ecological and evolutionary 38 capacity to adapt, causing abrupt and possibly irreversible changes (medium confidence). {2.3.1; 2.3.3; 39 2.4.2.6; 2.4.2.8; 2.6.1; Cross-Chapter Box ILNESS this Chapter, Cross-Chapter Box EXTREMES this 40 Chapter} 41 42 Since AR5, biome shifts and structural changes within ecosystems have been detected at an increasing 43 number of locations, consistent with climate change and increasing atmospheric CO₂ (high 44 confidence). New studies document changes that were projected in prior reports, including upward shifts in 45 the forest/alpine tundra ecotone, northward shifts in the deciduous/boreal forest ecotones, increased woody 46 vegetation in sub-Arctic tundra, and shifts in thermal habitat in lakes. A combination of changes in grazing, 47 browsing, fire, climate, and atmospheric CO₂ are leading to observed woody encroachment into grasslands 48 49 and savannas, consistent with projections from process-based models driven by precipitation, atmospheric CO₂ and wildfire (high confidence). {2.4.3; Table 2.3; Table 2.S.1; Box 2.1; Figure Box 2.1.1; Table Box 50 2.1.1} There is high agreement between projected changes in earlier reports and recent observed trends for 51 areas of increased tree death in temperate and boreal forests and of woody encroachment in savannas, 52 grasslands and tundra {2.5.4; Box 2.1; Figure Box 2.1.1; Table Box 2.1.1}. Observed changes impact 53 structure, functioning and resilience of ecosystems, and ecosystem services such as climate regulation (high 54 confidence). {2.3; 2.4.2; 2.4.3; 2.4.4, 2.5.4, Figure 2.11, Table 2.5, Box 2.1; Figure Box 2.1.1; Table Box 55 2.1.156

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Regional increases in area burned by wildfires (up to double natural levels), tree mortality up to 20%, and biome shifts up to 20 km latitudinally and 300 m upslope, have been attributed to anthropogenic climate change in tropical, temperate and boreal ecosystems around the world (*high confidence*), damaging key aspects of ecological integrity. This degrades vegetation survival, habitat for biodiversity, water supplies, carbon sequestration, and other key aspects of the integrity of ecosystems and their ability to provide services for people (*high confidence*). {2.4.3.1, 2.4.4.2, 2.4.4.3, 2.4.4.4; Table 2.3; Table 2.S.1}

- B.11 Fire seasons have lengthened on one-quarter of vegetated area since 1979 as a result of increasing 8 temperature, aridity, and drought (medium confidence). Field evidence shows that anthropogenic climate 9 change has increased the area burned by wildfire above natural levels in western North America from 10 1984–2017 by double for the Western USA and 11 time higher than natural in one extreme year in 11 British Columbia (high confidence). Burned area has increased in the Amazon, the Arctic, Australia, and 12 parts of Africa and Asia, consistent with, but not formally attributed to anthropogenic climate change. 13 Wildfires generate up to one-third of global ecosystem carbon emissions, a feedback that exacerbates climate 14 change (*high confidence*). Deforestation, peat draining, agricultural expansion or abandonment, fire 15 suppression, and inter-decadal cycles such as the El Niño-Southern Oscillation, can exert a stronger 16 influence than climate change on increasing or decreasing wildfire. {2.4.4.2; Table 2.3; Table 2.S.1; 17 FAQ2.1}. Increases in wildfire from levels to which ecosystems are adapted degrades vegetation, habitat for 18 biodiversity, water supplies, and other key aspects of the integrity of ecosystems and their ability to provide 19
- services for people (*high confidence*). {2.4.3.1, 2.4.4.2, 2.4.4.3, 2.4.4.4; Table 2.3; Table 2.S.1}

Anthropogenic climate change has caused drought-induced tree mortality of up to 20% in the period 22 1945–2007 in three regions in Africa and North America. It has also potentially contributed to over 100 23 other cases of drought-induced tree mortality across Africa, Asia, Australia, Europe, and North and South 24 America (high confidence). Field observations document post-mortality vegetation shifts (high confidence). 25 Timber cutting, agricultural expansion, air pollution, and other non-climate factors also contribute to tree 26 death. Climate-change driven increases in forest insect pests have contributed to mortality and changes in 27 carbon dynamics in many temperate and boreal forest areas (very high confidence). The direction of changes 28 in carbon balance and wildfires following insect outbreaks depends on the local forest-insect communities 29 (medium confidence). {2.4.4.3; Table 2.3; Table 2.S.1}. 30

31 32 **Terrest**

Terrestrial ecosystems currently remove more carbon from the atmosphere, 2.5–4.3 Gt y⁻¹, than they emit. Intact tropical rainforests, Arctic permafrost, and other healthy high carbon ecosystems provide 33 a vital global ecosystem service of preventing release of stored carbon (high confidence). Terrestrial 34 ecosystems contain stocks of ~3500 GtC in vegetation, permafrost, and soils, three to five times the amount 35 of carbon in unextracted fossil fuels (high confidence), and >4 times the carbon currently in the atmosphere 36 (high confidence). Tropical forests and Arctic permafrost contain the highest ecosystem carbon, with 37 peatlands following (high confidence). Deforestation, draining and burning of peatlands, and thawing of 38 Arctic permafrost due to climate change shifts these ecosystems from carbon-sinks to carbon-sources (high 39 confidence). {2.4.3.6; 2.4.3.8; 2.4.3.9. 2.4.4.4} 40 41

Evidence indicates that climate change is affecting many species, ecosystems, and ecological processes that provide ecosystem services connected to human health, livelihoods, and well-being *(medium*)

confidence). These services include climate regulation, water and food provisioning, pollination of crops,
 tourism and recreation. It is difficult establish end-to-end attribution from climatic changes to changes in a
 given ecosystem service and to identify the location and timing of impacts. This limits specific adaptation
 planning, but protection and restoration of ecosystems could build resilience of service provision. {2.2; 2.3;
 2.4.2.7; 2.4.5; 2.6.3; 2.6.4; 2.6.5; 2.6.6; 2.6.7; Cross-Chapter Box NATURAL this Chapter; Cross-Chapter
 Box ILLNESS this Chapter; Cross-Chapter Box EXTREMES this Chapter; Cross-Chapter Box COVID in
 Chapter 7; Cross-Chapter Box MOVING PLATE in Chapter 5}

52 Projected Risks

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 54 Climate change increases risks to fundamental aspects of terrestrial and freshwater ecosystems, with
- ⁵⁵ the potential for species' extinctions to reach 60% at 5°C GSAT warming (*high confidence*), biome
- shifts (changes in the major vegetation form of an ecosystem) on 15% (at 2°C warming) to 35% (at
- 57 4°C warming) of global land (medium confidence), and increases in the area burned by wildfire of 35%

(at 2°C warming) to 40% (at 4°C warming) of global land (medium confidence). {2.5.1; 2.5.2; 2.5.3; 1 2.5.4; Figure 2.6; Figure 2.7; Figure 2.8; Figure 2.9; Figure 2.11; Table 2.5; Table 2.S.2; Table 2.S.4; Cross-2

Chapter Box DEEP in Chapter 1; Cross-Chapter Paper 1} 3

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Extinction of species is an irreversible impact of climate change, the risk of which increases steeply

with rises in global temperature. It is *likely* that the percentage of species at high risk of extinction (median 6

and maximum estimates) will be 9% (max 14%) at 1.5°C, 10% (max 18%) at 2°C, 12% (max 29%) at 3.0°C, 7 13% (max 39%) at 4°C and 15% (max 48%) at 5°C (Figure 2.7). Among groups containing largest numbers 8

of species at high risk of extinctions for mid-levels of warming (3.2°C) are: invertebrates (15%), specifically 9

pollinators (12%), amphibians (11%, but salamanders are at 24%) and flowering plants (10%). All groups 10

fare substantially better at 2°C, with extinction projections reducing to <3% for all groups, except 11

- salamanders at 7% (medium confidence) (Figure 2.8a). Even the lowest estimates of species' extinctions 12 (9%) are 1000x natural background rates. Projected species' extinctions at future global warming levels are
- 13 consistent with projections from AR4, but assessed on many more species with much greater geographic 14
- coverage and a broader range of climate models. {2.5.1.3; Figure 2.6; Figure 2.7; Figure 2.8; Cross-Chapter 15
- Box DEEP in Chapter1; Cross-Chapter Paper 1} 16

17 Species are the fundamental unit of ecosystems, and increasing risk to species increases risk to 18

ecosystem integrity, functioning and resilience with increasing warming (high confidence). As species 19 become rare, their roles in the functioning of the ecosystem diminishes (high confidence). Loss of species 20 reduces the ability of an ecosystem to provide services and lowers its resilience to climate change (high 21 confidence). At 1.58°C (median estimate), >10% of species are projected to become endangered (sensu 22 IUCN); at 2.07°C (median) >20% of species are projected to become endangered, representing high and very 23 high biodiversity risk, respectively (medium confidence) {2.5.4; Figure 2.8b, Figure 2.11; Table 2.5, Table 24 2.S.4}. Biodiversity loss is projected for more regions with increasing warming, and to be worst in northern 25 South America, southern Africa, most of Australia, and northern high latitudes (medium confidence) {2.5.1.3; 26 Figure 2.6}. 27

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Climate change increases risks of biome shifts on up to 35% of global land at \geq 4°C warming, that 29 emissions reductions could limit to <15% for <2°C warming (medium confidence). Under high 30 warming scenarios, models indicate shifts of extensive parts of the Amazon rainforest to drier and 31 lower-biomass vegetation (medium confidence), poleward shifts of boreal forest into treeless tundra 32 across the Arctic, and upslope shifts of montane forests into alpine grassland (high confidence). Area at 33 high risk of biome shifts from climate change and land use change combined can double or triple compared 34 to climate change alone (medium confidence). Novel ecosystems, with no historical analogue, are expected 35 to become increasingly common in future (medium confidence). {2.3, 2.4.2.3.3, 2.5.2; 2.5.4, Figure 2.11; 36 Table 2.5; Table 2.8.4} 37 38

Risk of wildfires increases with global temperature (high confidence). With 4°C warming by 2100 39 wildfire frequency is projected to have a net increase of ~30% (medium confidence). Increased wildfire, 40 combined with soil erosion due to deforestation, could degrade water supplies (medium confidence). For 41 ecosystems with historically low fire frequencies, a projected 4°C global temperature rise increases risks of 42 fire, with potential increases in tree mortality and conversion of extensive parts of Amazon rainforest to drier 43 and lower-biomass vegetation (medium confidence). {2.5.3.2; 2.5.3.3} 44

45 Continued climate change substantially increases risk of carbon stored in the biosphere being released 46 into the atmosphere due to increases in processes such as wildfires, tree mortality, insect pest 47 outbreaks, peatland drying and permafrost thaw (high confidence). These phenomena exacerbate self-48 reinforcing feedbacks between emissions from high-carbon ecosystems (that currently store ~3030-4090 49 GtC) and increasing global temperatures. Complex interactions of climate change, land use change, carbon 50 dioxide fluxes, and vegetation changes, combined with insect outbreaks and other disturbances, will regulate 51 the future carbon balance of the biosphere, processes incompletely represented in current earth system 52 models. The exact timing and magnitude of climate-biosphere feedbacks and potential tipping points of 53 carbon loss are characterized by large uncertainty, but studies of feedbacks indicate increased ecosystem 54 carbon losses can cause large future temperature increases (medium confidence). {AR6 WGI 5.4, Table 5.4, 55 Figure 5.29; 2.5.2.7; 2.5.2.8; 2.5.2.9; 2.5.3.2; 2.5.3.3; 2.5.3.4; 2.5.3.5; Figure 2.10; Figure 2.11; Table 2.4; 56 57

Table 2.5; Table 2.S.2; Table 2.S.4}

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Contributions of Adaptation Measures to Solutions

3 The resilience of biodiversity and ecosystem services to climate change can be increased by human 4 adaptation actions including ecosystem protection and restoration (high confidence). Ecological theory 5 and observations show that a wide range of actions can reduce risks to species and ecosystem integrity. This 6 includes minimising additional stresses or disturbances, reducing fragmentation, increasing natural habitat 7 extent, connectivity and heterogeneity, maintaining taxonomic, phylogenetic and functional diversity and 8 redundancy; and protecting small-scale refugia where microclimate conditions can allow species to persist 9 (high confidence). Adaptation also includes actions to aid the recovery of ecosystems following extreme 10 events. Understanding the characteristics of vulnerable species can assist in early warning systems to 11 minimise negative impacts and inform management intervention. {2.3; Figure 2.1; 2.5.3.1, 2.6.2, Table 2.6, 12 2.6.5, 2.6.7, 2.6.8}. 13

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There is new evidence that species can persist in refugia where conditions are locally cooler, when they are declining elsewhere (*high confidence*) {2.6.2}. Protecting refugia, for example where soils remain wet during drought or fire risk is reduced, and in some cases creating cooler microclimates, are promising adaptation measures {2.6.3; 2.6.5; CCP1; CCP5.2.1}. There is also new evidence that species can persist

locally because of plasticity, including changes in phenology or behavioural changes that move an individual
 into cooler micro-climates, and genetic adaptation may allow species to persist for longer than might be
 expected from local climatic changes (*high confidence*) {2.4.2.6; 2.4.2.8, 2.6.1}. There is no evidence to
 indicate that these mechanisms will prevent global extinctions of rare, very localised species at their climatic
 limits or species inhabiting climate/habitat zones that are disappearing (*high confidence*). {2.4.2.8, 2.5.1,

24 2.5.3.1, 2.5.4, 2.6.1, 2.6.2, 2.6.5}

Since AR5, many adaptation plans and strategies have been developed to protect ecosystems and biodiversity but there is limited evidence of the extent to which adaptation is taking place and virtually no evaluation of the effectiveness of adaptation measures in the scientific literature (medium confidence). This is an important evidence gap that needs to be addressed to ensure a baseline is available against which to judge effectiveness and develop and refine adaptation in future. Many proposed adaptation measures have not been implemented (low confidence) {2.6.2; 2.6.3; 2.6.4; 2.6.5; 2.6.6; 2.6.8; 2.7}

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Ecosystem restoration and resilience building cannot prevent all impacts of climate change, and 33 adaptation planning needs to manage inevitable changes to species distributions, ecosystem structure 34 and processes (very high confidence). Actions to manage inevitable change include local modification of 35 microclimate or hydrology, adjustment of site management plans and facilitating the dispersal of vulnerable 36 species to new locations, both by increasing habitat connectivity or by active translocations of species. 37 Adaptation can reduce risks but cannot prevent all damaging impacts so is not a substitute for reductions in 38 greenhouse gas emissions (high confidence). {2.2; 2.3; 2.3.1; 2.3.2; 2.4.5; 2.5.1.3; 2.5.1.4; 2.5.2; 2.5.3.1; 39 2.5.3.5; 2.5.4; 2.6.1; 2.6.2; 2.6.3; 2.6.4; 2.6.5; 2.6.6; 2.6.8; Cross-Chapter Box NATURAL this Chapter }. 40 41

Ecosystem-base Adaptation (EbA)can deliver climate change adaptation for people with multiple 42 additional benefits, including for biodiversity (high confidence). An increasing body of evidence 43 demonstrates that climatic risks to people, including from flood, drought, fire and over-heating, can be 44 lowered by a range of Ecosystem-based Adaptation techniques in urban and rural areas (medium confidence). 45 EbA forms part of a wider range Nature-based Solutions (NbS) actions and some have mitigation co-46 benefits, including the protection and restoration of forests and other high-carbon ecosystems, as well as 47 agroecological farming practices {2.6.3; 2.6.5; Cross-Chapter Box NATURAL this Chapter}. However, EbA 48 and other NbS are still not widely implemented. {2.2; 2.5.3.1; 2.6.2; 2.6.3; 2.6.4; 2.6.5; 2.6.6, 2.6.7; Table 49 2.7; Cross-Chapter Box NATURAL this Chapter; Cross-Chapter Paper 1}. 50

To realise potential benefits and avoid harm, it is essential that EbA is deployed in the right places and with the right approaches for that area, with inclusive governance (*high confidence*). Interdisciplinary scientific information and practical expertise, including local and Indigenous knowledge, are essential to effectiveness (*high confidence*). There is a large risk of maladaptation where this does not happen (*high confidence*). {1.4.2; 2.2; 2.6; Table 2.7; Box 2.2; Figure Box 2.2.1; Cross-Chapter Box NATURAL this Chapter; Cross-Chapter Paper1; 5.14.2}.

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Ecosystem-based Adaptation and other Nature-based Solutions are themselves vulnerable to climate change impacts. They need to take account of climate change adaptation if they are to remain effective and will increasingly be under threat at higher warming levels. Nature-based Solutions cannot be regarded as an 4 alternative to, or a reason to delay, deep cuts in greenhouse gas emissions (high confidence). {2.6.3, 2.6.5; 2.6.7; Cross-Chapter Box NATURAL this Chapter} 6

Climate Resilient Development

Protection and restoration of natural and semi-natural ecosystems are key adaptation measures in 10 light of clear evidence that damage and degradation of ecosystems exacerbates the impacts of climate 11 change on biodiversity and people (high confidence). Ecosystem services that are at threat from a 12 combination of climate change and other anthropogenic pressures include climate change mitigation, flood 13 risk management, food provisioning and water supply (high confidence). Adaptation strategies that treat 14 climate, biodiversity and human society as coupled systems will be most effective. {2.3; Figure 2.1; 2.5.4; 15 2.6.2; 2.6.3; 2.6.7; Cross-Chapter Box NATURAL and Cross-Chapter Box ILLNESS this Chapter 16 17

A range of analyses have concluded that ~30% of Earth's surface needs to be effectively conserved to 18maintain biodiversity and ecosystem services (medium confidence). Climate change places additional 19 stress on ecosystem integrity and functioning, adding urgency for taking action. Low intensity, 20 sustainable management, including by Indigenous peoples, is an integral part of some protected areas and 21 can support effective adaptation and maintain ecosystem health. Food and fibre production in other areas 22 will need to be efficient, sustainable and adapted to climate change to meet the needs of the human 23 population (*high confidence*). {Figure 2.1; 2.5.4; 2.6.2; 2.6.3; 2.6.7} 24

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Natural ecosystems can provide carbon storage and sequestration at the same time as providing 26 multiple other ecosystem services, including EbA (high confidence) but there are risks of 27 maladaptation and environmental damage from some approaches to land-based mitigation (high 28 confidence). Plantation forests in areas which would not naturally support forest, including savannas, natural 29 grasslands and temperate peatlands, or replacing native tropical forests on peat soils, have destroyed local 30 biodiversity and created a range of problems, including for water supply, food supply, fire risk and 31 greenhouse gas emissions. Large scale deployment of bioenergy, including Bioenergy with Carbon Capture 32 and Storage (BECCS) through dedicated herbaceous or woody bioenergy crops and non-native production 33 forests can damage ecosystems directly or through increasing competition for land. {2.6.3, 2.6.5, 2.6.6, 34 2.6.7; Box 2.2; Cross-Chapter Box NATURAL this Chapter; CCP7.3.2; Cross-Working Group Box 35 BIOECONOMY in Chapter 5}. 36 37

Terrestrial and aquatic ecosystems and species are often less degraded in lands managed by 38

Indigenous Peoples and local communities than in other lands (medium confidence). Including 39 indigenous and local institutions is a key element in developing successful adaptation strategies. Indigenous 40 and local knowledge contain a wide variety of resource-use practices and ecosystem stewardship strategies 41 that conserve and enhance both wild and domestic biodiversity. {2.6.5; 2.6.7; Cross-Chapter Box 42 NATURAL this Chapter} 43

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Extreme events are compressing the timeline available for natural systems to adapt and impeding our 45

ability to identify, develop and implement solutions (medium confidence). There is now an urgent need to 46

build resilience and assist recovery following extreme events. This, combined with long-term changes in 47 baseline conditions means that implementation of adaptation and mitigation measures cannot be delayed if

- 48 49 they are to be fully effective. {2.3; Cross-Chapter Box EXTREMES this Chapter}
- 50

2.1 Introduction

1 2

2.1.1 **Overview** 3

4 Chapter 2: We provide assessments of observed and projected impacts of climate change across species, 5 biomes (vegetation types), ecosystems and ecosystem services, highlighting processes emerging on a global 6 scale. Where sufficient evidence exists, differences in biological responses among regions, taxonomic groups 7 or types of ecosystems are presented, particularly when such differences provide meaningful insights into 8 current or potential future autonomous or human-mediated adaptations. Human interventions that might 9 build resilience of ecosystems and minimise negative impacts of climate change on biodiversity and 10 ecosystem functioning are assessed. Such interventions include adaptation strategies and programmes to 11 support biodiversity conservation and Ecosystem-based Adaptation (EbA). The assessments were done in the 12 context of the Convention on Biological Diversity (CBD) and Sustainable Development Goals (SDGs), 13 whose contributions to climate-resilient development pathways are assessed. This chapter highlights both 14 successes and failures of adaptation attempts and considers potential synergies and conflicts with land-based 15 climate change mitigation. Knowledge gaps and sources of uncertainty are included to encourage additional 16 research. 17

18The Working Group II Summary for Policy Makers of the 5th Assessment Report (WGII AR5 SPM) stated 19 that "many terrestrial and freshwater species have shifted their geographic ranges, seasonal activities, 20 migration patterns, abundances, and species interactions in response to ongoing climate change" (IPCC, 21 2014e). Based on long-term observed changes across the regions, it was estimated that approximately 20-22 30% of plant and animal species are at risk of extinction when global mean temperatures rise 2–3°C above 23 preindustrial levels (Fischlin et al., 2007). In addition, WGII AR5 (IPCC, 2014f) broadly suggested that 24 autonomous adaptation by ecosystems and wild species might occur, and proposed human-assisted 25 adaptation to minimise negative climate change impacts. 26

27

Risk assessments for species, communities, key ecosystems and their services were based on the Risk 28 Assessment Framework introduced in the IPCC AR5 (IPCC, 2014). Assessments of observed changes in 29 biological systems emphasise detection and attribution of climate change on ecological and evolutionary 30 processes with an emphasis on freshwater ecosystems, and assess ecosystem processes that were lightly 31 assessed in previous reports, such as wildfire. Where appropriate, assessment of interactions between climate 32 change and other human activities is provided. 33

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Land-use and land cover change (LULCC), and unsustainable exploitation of resources from terrestrial and 35 freshwater systems continue to be a major factor of natural ecosystem and biodiversity loss (high 36 confidence). Fertiliser input, pollution of waterways, dam construction and extraction of freshwater for 37 irrigation put additional pressure on biodiversity and alter ecosystem function (Shin et al., 2019). Likewise, 38 for biodiversity, invasive alien species have been identified as a major threat, especially in freshwater 39 systems, islands and coastal regions (high confidence) (IPBES, 2018b; IPBES, 2018e; IPBES, 2018c; 40 IPBES, 2018d; IPBES, 2019). Climate change and CO_2 are expected to become increasingly important as 41 drivers of change over the coming decades (Ciais et al., 2013; Settele et al., 2014; IPBES, 2019; IPCC, 42 2019c). 43

2.1.2 **Points of Departure** 45

46 Species diversity and ecosystem function influence each other reciprocally, while ecosystem function forms 47 the necessary basis for ecosystem services (Hooper et al., 2012; Mokany et al., 2016). Drivers of impacts on 48 49 biodiversity, ecosystem function and ecosystem services have been assessed in reports from IPCC, Food and Agriculture Organization (FAO), IPBES and the Global Environmental Outlook (Settele et al., 2014; FAO, 50 2018; IPBES, 2018b; IPBES, 2018c; IPBES, 2018c; IPBES, 2018d; IPBES, 2019; UNEP, 2019; Diversity, 51 2020). Most recently, the IPCC SRCCL provided an assessment on land degradation and desertification, 52 greenhouse gas emissions and food security in the context of global warming (IPCC, 2019c), and the IPBES-53 IPCC joint report on Biodiversity and Climate Change provided a synthesis of current understanding of the 54 interactions, synergies and feedbacks between biodiversity and climate change (Pörtner et al., 2021). This 55 chapter builds on and expands the results from these assessments. 56 57

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1	Assessment of impacts of climate cha	nge on freshwater systems l	has been limited in previous assessments,
1	1	6	1
2	and interlinkages between terrestrial a	and freshwater processes hav	ve not been fully explored (Settele et al.,

2014; IPBES, 2019). Improved treatment of impacts on terrestrial and freshwater systems is critical
 considering the revisions of international sustainability Goals and Targets, especially the conclusion that
 many of the proposed post-2020 Biodiversity-targets of the Convention on Biological Diversity (CBD)

- 6 cannot be met due to climate change impacts (Arneth et al., 2020).
- 7

Previous reports highlighted the possibility of new ecosystem states stemming from shifts in thermal 8 regimes, species composition, and energy and matter flows (Settele et al., 2014; Shin et al., 2019). Projecting 9 such "tipping points" (see glossary) has been identified in previous reports as a challenge since neither 10 monitoring programmes nor field studies, nor ecosystem and biodiversity modelling tools capture the 11 underlying species-species and species-climate interactions sufficiently well to identify how biological 12 interactions within and across trophic levels may amplify or dampen shifts in ecosystem states (Settele et al., 13 2014; Shin et al., 2019). Building on these previous analyses and recent literature, Chapter 2 in this AR6 14 provides new insights compared to previous assessments by (i) emphasising freshwater aspects, and the 15 interlinkages between freshwater and terrestrial systems, (ii) assessing more clearly the link between 16 biodiversity and ecosystem functioning, (iii) assessing impacts associated with climate change mitigation 17 scenarios versus impacts of climate change, including interactions with adaptation, and (iv) where possible, 18 places findings in context of the United Nations Sustainable Development Goals (SDGs) 2030, and services 19 for human societies. 20

21
 22 2.1.3 Guide to Attribution and Traceability of Uncertainty Assessments

For biological systems we use the framework for detection and attribution outlined in AR5 in which attribution of observed biological changes is made not to global, but to local or regional climate changes, (Parmesan et al., 2013; Cramer et al., 2014). However, global distribution of regional responses is desirable to achieve generality, and data in prior reports were concentrated from the northern hemisphere. The critique of "global" studies by (Feeley et al., 2017) argues that their naming is misleading, that most of them are far from global and that considerable geographic and taxonomic bias remains. This bias is diminishing, as data from southern hemisphere regions are added and there is now representation from every continent.

31

Overall confidence in climate change attribution of a biological change can be increased in multiple ways 32 (Parmesan et al., 2013), of which we list four here. First, confidence rises when the time span of biological 33 records is long, such that decadal trends in climate can be compared with decadal trends in biological 34 response and long-term trends can be statistically distinguished from natural variability. Secondly, 35 confidence can be increased by examining a large geographic area, which tends to diminish the effects of 36 local confounding factors (Parmesan et al., 2013; Daskalova et al., 2020). Third, confidence is increased 37 when there is experimental or empirical evidence of a mechanistic link between particular climate metrics 38 and biological response. Fourth, confidence is increased when particular fingerprints of climate change are 39 documented that uniquely implicate climate change as the causal driver of the biological change (Parmesan 40 and Yohe, 2003). These conditions constitute multiple lines of evidence which, when they converge, can 41 provide very high confidence that climate change is the causal driver of an observed change in a particular 42 biological species or system (Parmesan et al., 2013). 43

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Important factors that may confound or obscure effects of climate change are presence of invasive species, 45 changes in land use (LULCC), and, in freshwater systems, eutrophication (IPCC, 2019a). Temporal and 46 spatial scale of the study also affect estimates of impacts. The most extreme published estimates of 47 biological change tend to be derived from smaller areas and/or shorter timeframes (Daskalova et al., 2020), 48 49 and a recent large global analysis of data for 12,415 species found that differences in study methodology accounted for most of the explained variance in reported range shifts (Lenoir et al., 2020). The importance of 50 land-use change is frequently stressed, but there is a paucity of studies that actually quantify the relative 51 effects of climate change and land-use change on species and communities. Sirami et al. (2017) found only 52 13 such studies, among which four concluded that effects of land-use change over-rode those of climate 53 change, four found that the two drivers independently affected different species and five found that they 54 acted in synergy. 55

1 2

2.2 Connections of Ecosystem Services to Climate Change

Ecosystems provide services essential for human survival and well-being. The Millennium Ecosystem Assessment defined ecosystem services as "the benefits people obtain from ecosystems" including "provisioning services such as food and water; regulating services such as regulation of floods, drought, land degradation, and disease; supporting services such as soil formation and nutrient cycling; and cultural services such as recreational, spiritual, religious, and other nonmaterial benefits" (Assessment, 2005).

- 8 The Intergovernmental Science-Policy Platform on Biodiversity and Ecosystem Services (IPBES) re-named 9 the concept "nature's contributions to people" and broadened the definition to "the contributions, both 10 positive and negative, of living nature (i.e. diversity of organisms, ecosystems, and their associated 11 ecological and evolutionary processes) to the quality of life for people. Beneficial contributions from nature 12 include such things as food provision, water purification, flood control, and artistic inspiration, whereas 13 detrimental contributions include disease transmission and predation that damages people or their assets" 14 (IPBES, 2019). IPBES modified the concept to include more social viewpoints and broaden analyses beyond 15 narrow economic stock-and-flow valuation approaches (Díaz et al., 2018). IPBES developed a classification 16 of 18 categories of ecosystem services (see Table 2.1). 17
- 18

19 When anthropogenic climate change affects ecosystems, it can also affect ecosystem services for people.

20 Climate change connects to ecosystem services through three links: climate change—species—ecosystems—

ecosystem services. This IPCC chapter assesses these connections through all three links when end-to-end published scientific analyses are available for terrestrial and freshwater ecosystems. This type of robust

published scientific analyses are available for terrestrial and freshwater ecosystems. This type of robust
 evidence exists for some key ecosystem services (Section 2.5.4), and is assessed in specific report sections:

evidence exists for some key ecosystem services (Section 2.5.4), and is assessed in specific report sections
 biodiversity habitat creation and maintenance (Sections 2.4, 2.5), regulation of detrimental organisms and

biological processes (Sections 2.4.2.3, 2.4.2.7, 2.4.4, 2.5.3, 2.6.4, Cross-Chapter Box ILLNESS this

26 Chapter), regulation of climate through ecosystem feedbacks in terms of carbon storage (Sections 2.4.4.4,

- 27 2.5.2.10, 2.5.3.4, 2.5.3.5) and albedo (Section 2.5.3.5), and provision of freshwater from ecosystems to 28 people (Section 2.5.3.6).
- 29

For ecosystem services that do not have published scientific information to establish unambiguous links to climate change, the climate—species—ecosystem links are assessed. Global ecological assessments, including the Global Biodiversity Assessment (Programme, 1995), the Millennium Ecosystem Assessment (Assessment, 2005), and the IPBES Global Assessment Report (IPBES, 2019) have synthesised scientific information on the ecosystem—ecosystem services link, but full assessment from climate change to ecosystem services is often impeded by limited quantitative studies that span this entire spectrum, see (Mengist et al., 2020) for a review of this gap in montane regions.

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IPCC and IPBES are collaborating to address gaps in knowledge on the effects of climate change on
 ecosystem services (Services and Ecosystem, 2021). Table 2.1 provides a guide for finding information on
 climate change and individual ecosystem services in the IPCC Sixth Assessment Report.

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Table 2.1: Connections of ecosystem services to climate change, indicating the 18 categories of nature's contributions to people of the IPBES (IPBES, 2019), the most relevant sections in the IPCC Sixth Assessment Report, and the level of evidence in this report for attribution to anthropogenic climate change of observed impacts on ecosystem services. The order of services in the table follows the order presented by IPBES and does not denote importance or priority.

The order of services in the table follows the order presented by IPBES and does not denote importance or priority.
 Connections denote observed impacts, future risks, and adaptation. The order of connections follows the relevance or

48 <u>the order of sections</u>.

Ecosystem service	Connections to climate change
Habitat creation and maintenance	Species extinctions (2.4.2.2, 2.5.1.3), Species range shifts (2.4.2.1, 2.4.2.5), Ecological changes in freshwater ecosystems (2.3.3, 2.4.2.3.2, 2.4.4.1, 2.4.4.5.2, 2.5.1.3.2, 2.5.3.5, 2.5.4, 2.5.3.6, 2.5.5.8), Vegetation changes (2.4.3, 2.4.4.2.5, 2.4.4.3, 2.4.4.4, 2.4.4.5.1, 2.5.2, 2.5.3.3), Biome shifts (2.4.3.2, 2.5.4), Wildfire (2.4.4.2, 2.5.3.2), Tree mortality (2.4.4.3, 2.5.3.3) (<i>robust evidence</i>)

FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
Pollination and dispersal of seeds and other propagules	Species extinctions (2.4.2.2, 2.5.1. Phenology changes (2.4.2.4, 2.4.2. (<i>medium evidence</i>)	3), Species range shifts (2.4.2.1, 2.4.2.5), 5)
Regulation of air quality	Wildfire (2.4.4.2, 2.5.3.2, Chapter (<i>medium evidence</i>)	7), Tree mortality (2.4.4.3, 2.5.3.3)
Regulation of climate	Working Group I, Chapter 5), Am 2.4.4.4.2, 2.5.2.6, 2.5.2.10, 2.5.3.3 2.5.3.5, 2.5.4), biome shifts (2.4.3,	ns, and removals (2.4.4.4, 2.5.3.4, IPCC AR6 azon rainforest dieback (2.4.3.6, 2.4.4.3.2,), Tundra permafrost thaw (2.4.4.4, 2.5.2.8, , 2.5.2, 2.5.3.2.2), Wildfire (2.4.4.2, 2.5.3.2), rimary productivity changes (2.4.4.5, 2.5.3.5)
Regulation of ocean acidification	Ocean acidification (IPCC AR6 W marine species distribution and ab (<i>robust evidence</i>)	Vorking Group I, Chapter 5), Changes in undance (Chapter 3)

Regulation of freshwater quantity, location, and timing	Physical changes in freshwater systems (2.3.3), Ecological changes in freshwater ecosystems (2.4.2.3.2, 2.4.4.1, 2.4.4.5.2, 2.5.1.3.2, 2.5.3.7), Tree mortality (2.4.4.3, 2.5.3.3), Freshwater supply from ecosystems (2.5.3.6) (<i>medium evidence</i>)
Regulation of freshwater and coastal water quality	Coastal ecosystem changes (Chapter 3), Physical changes in freshwater systems (2.3.3), Ecological changes in freshwater ecosystems (2.4.2.3.2; 2.4.4.1, 2.4.4.5.2, 2.5.1.3.2, 2.5.3.7) (<i>robust evidence</i>)
Formation, protection, and decontamination of soils and sadiments	Agricultural ecosystem changes (Chapter 5), Physical changes in freshwater systems (2.3.1), Vegetation changes (2.4.3, 2.5.4), Wildfire (2.4.4.2, 2.5.3.2) (madium avidance)

- sediments (medium evidence) Coastal ecosystem changes (Chapter 3), Vegetation changes (2.4.3, 2.5.2), Regulation of hazards and extreme events Wildfire (2.4.4.2, 2.5.5.2), Summary of hazards (2.3), Cross-Chapter Box EXTREMES this Chapter (medium evidence)
 - Inter-species interactions (2.4.2), Control of disease vectors (2.4.2.7, 2.5.1, 2.6.4), Insect pest infestations (2.4.4.3), Cross-Chapter Box ILLNESS this Chapter (medium evidence)

Forestry plantation changes (Chapter 5), Biomass changes in natural ecosystems (2.4.4.4), Bioeconomy (Cross-Working Group Box BIOECONOMY in Chapter 5), Tree mortality (2.4.4.3, 2.5.3.3) *(limited evidence)*

Agricultural ecosystem changes (Chapter 5), Species extinctions (2.4.2.2, 2.5.1.3), Species range shifts (2.4.2.1), Nature-based services from natural ecosystems (Cross-Chapter Box NATURAL this chapter), shifts in commercial food species Cross-Chapter Box Moving Plate in Chapter 5) (medium evidence)

Materials, companionship, and labour	Forestry plantation changes (Chapter 5), Species extinctions (2.4.2.2, 2.5.1.3), Species range shifts (2.4.2.1), Tree mortality (2.4.4.3, 2.5.3.3) (<i>limited evidence</i>)
Medicinal, biochemical, and genetic resources	Species extinctions (2.4.2.2, 2.5.1.3), Species range shifts (2.4.2.1) (<i>limited evidence</i>)
Learning and inspiration	All observed impacts (2.4) and future risks (2.5) in terrestrial and freshwater ecosystems

Regulation of detrimental

organisms and biological

processes

Energy

Food and feed

FINAL DRAFT

(limited evidence)

FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
Physical and psychological experiences	All observed impacts (2.4) and future risks (2.5) in terrestrial and freshwater ecosystems (<i>limited evidence</i>)	
Supporting identities	All observed impacts (2.4) and freecosystems (<i>limited evidence</i>)	uture risks (2.5) in terrestrial and freshwater
Maintenance of options	ecosystems, Nature-based servic	uture risks (2.5) in terrestrial and freshwater es from natural ecosystems (Cross-Chapter ross-Chapter Box DEEP in Chapter 17

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2.3 Hazards and Exposure

In AR6, WGI describes changes in physical climate systems using the term 'climatic impact-drivers' (CIDs), which can have detrimental, beneficial or neutral effects on a system. In contrast, the literature on natural systems tends to focus on hazards, which include natural or human-induced physical events, impacts or trends with the potential to cause negative effects on ecosystems and environmental resources. Hazards are affected by current and future changes in climate, including altered climate variability and extreme events (WGI Chapter 12). Hazards can occur suddenly (e.g., a heat wave or heavy rain event), or more slowly (e.g., land loss, degradation and erosion linked to multiple climate hazards compounding). Observed exposure and risks to protected areas is assessed in Section 2.5.3.1.1. See also Cross-Chapter Box EXTREMES this Chapter.

Non-climatic hazards such as land use change, habitat fragmentation, pollution and invasive species have
been the primary drivers of change in terrestrial and freshwater ecosystems in the past (*high confidence*)
(Figure 2.1). These impacts have been extensively documented in reports by the Intergovernmental SciencePolicy Platform on Biodiversity and Ecosystem Services (IPBES, 2021). However, whilst climate change has
not been the predominant influence to date, its relative impact is increasing (IPCC Special Report on Climate
Change and Land(SRCCL)), with greater interactive effects of non-climate and climate hazards now
occurring (Birk et al., 2020).

23 2.3.1 Observed Changes to Hazards and Extreme Events

The major climate hazards at the global level are generally well understood (WGI AR6 Chapter 12; WGI AR6 Interactive Atlas). Increased temperatures and changes to rainfall and runoff patterns; greater variability in temperature, rainfall, river flow and water levels; rising sea-levels and increased frequency of extreme events means that greater areas of the world are being exposed to climate hazards outside those to which they are adapted (*high confidence*) (Lange et al., 2020).

Extreme events are a natural and important part of many ecosystems and many organisms have adapted to 31 cope with long-term and short-term climate variability, within the disturbance regime experienced during 32 their evolutionary history (high confidence). However, climate changes, disturbance regimes change and the 33 magnitude and frequency of extreme events such as floods, droughts, cyclones, heatwaves and fire have 34 increased in many regions (high confidence). These disturbances affect ecosystem functioning, biodiversity 35 and ecosystem services (high confidence) but are, in general, poorly captured in impact models (Albrich et 36 al., 2020b), although this should improve as higher-resolution climate models that better capture smaller-37 scale processes and extreme events become available (WGI AR6, Chapter 11). Extreme events pose large 38 challenges for Ecosystem-based Adaptation (IPCC Special Reports on Extremes, Section 2.6.3). Ecosystem 39 functionality on which such adaptation measures rely may be altered or destroyed by extreme episodic 40 41 events (Handmer et al., 2012; Lal et al., 2012; Pol et al., 2017).

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There is *high confidence* that the combination of internal variability, superimposed on longer-term climate trends, is pushing ecosystems to tipping points, beyond which abrupt and possibly irreversible changes are occurring (Harris et al., 2018a; Jones et al., 2018; Hoffmann et al., 2019b; Prober et al., 2019; Berdugo et al., 1 2020; Bergstrom et al., 2021). Increases in the frequency and severity of heatwaves, droughts and aridity, 2 floods, fires and extreme storms have been observed in many regions (Seneviratne et al., 2012; Ummenhofer

floods, fires and extreme storms have been observed in many regions (Seneviratne et al., 2012; Ummer and Meehl, 2017) and these trends are projected to continue (*high confidence*) (Section 3.2.2.1, Cross-

Chapter Box EXTREMES this Chapter, AR6 WGI Chapter 11;, SR1.5, Hoegh-Guldberg et al., 2018b).

4 5

While the major climate hazards at the global level are generally well described with high confidence, there 6 is less understanding about the importance of hazards on ecosystems when they are superimposed (Allen et 7 al., 2010; Anderegg et al., 2015; Seidl et al., 2017; Dean et al., 2018), and the outcomes are difficult to 8 quantify in future projections (Handmer et al., 2012). Simultaneous or sequential events (coincident or 9 compounding events) can lead to an extreme event or impact, even if each event is not in themselves extreme 10 (Denny et al., 2009; Hinojosa et al., 2019). For example, the compounding effects of sea-level rise, extreme 11 coastal high tide, storm surge, and river flow can substantially increase flooding hazard and impacts on 12 freshwater systems (Moftakhari et al., 2017). On land, changing rainfall patterns and repeated heat waves 13 may interact with biological factors such as altered plant growth and nutrient allocation under elevated CO_2 , 14 affecting herbivory rates and insect outbreaks leading to widespread dieback of some forests (e.g. in 15 Australian Eucalypt forests) (Gherlenda et al., 2016; Hoffmann et al., 2019a). Risk assessments typically 16 only consider a single climate hazard without changing variability, potentially underestimating actual risk (17 Milly et al., 2008; Sadegh et al., 2018; Zscheischler et al., 2018; Terzi et al., 2019; Stockwell et al., 2020;). 18 19

Understanding impacts associated with the rapid rate of climate change is less developed and more uncertain than changes in mean climate. High climate velocity (Loarie et al., 2009) is expected to be associated with distribution shifts, incomplete range filling and species extinctions (*high confidence*) (Sandel et al., 2011; Burrows et al., 2014), although not all species are equally at risk from high velocity (see Sections 2.4.2.2, 2.5.1.3). It is generally assumed that the more rapid the rate of change, the greater the impact on species and ecosystems, but responses are taxonomically and geographically variable (*high confidence*) (Kling et al.,

26 27 2020).

For example, strong dispersers are less at risk, while species with low dispersal ability, small ranges and long 28 lifespans (e.g. many plants, especially trees, many amphibians and some small mammals) are more at risk 29 (Hamann et al., 2015) (IPCC, 2014). This is likely to favour generalist and invasive species, altering species 30 composition, ecosystem structure and function (Clavel et al., 2011; Büchi and Vuilleumier, 2014). The 31 ability to track suitable climates is substantially reduced by habitat fragmentation and human modifications 32 of the landscape such as dams on rivers and urbanisation (high confidence). Freshwater systems are 33 particularly at risk of rapid warming given their naturally fragmented distribution. Velocity of change in 34 surface temperature of inland standing waters globally has been estimated as 3.5 ± 2.3 km per decade from 35 1861 to 2005. This is projected to increase from 2006 to 2099 from between 8.7 ± 5.5 km per decade (RCP 36 2.6) to 57.0 ± 17.0 km per decade (RCP 8.5) (Woolway and Maberly, 2020). Although the dispersal of aerial 37 adult stage of some aquatic insects can surpass these climate velocities, rates of change under mid- and high 38 emissions scenarios (RCP4.5, RCP6.0, RCP8.5) are substantially higher than known rates of active dispersal 39 of many species (Woolway and Maberly, 2020). Many species, both terrestrial and freshwater, are not 40 expected to be able to disperse fast enough to track suitable climates under mid- and high emissions 41 scenarios (medium confidence) (RCP4.5, RCP6.0, RCP8.5; Brito-Morales et al., 2018). 42

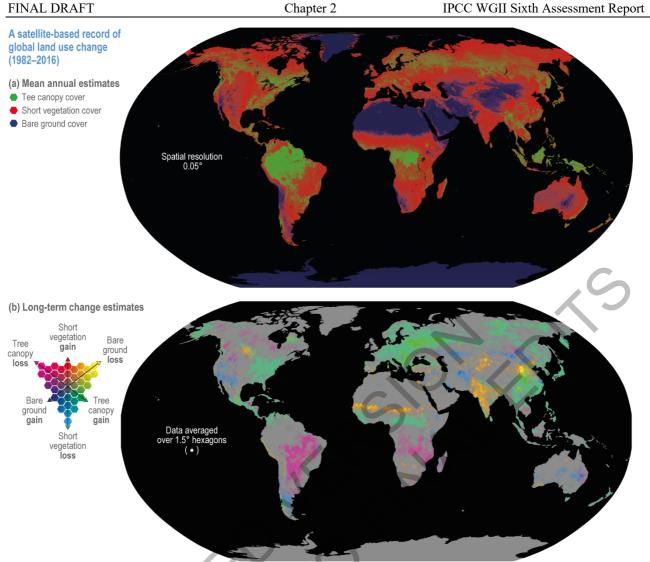


Figure 2.1: Map of global land use change from 1982–2016. Based on satellite records of global tree canopy (TC), short vegetation (SV) and bare ground (BG) cover (from Song et al., 2018). a) Mean annual estimates of cover (% of pixel area at 0.05° resolution). b) Long-term change estimates (% of pixel area at 1.5° resolution), with pixels showing a statistically significant trend (n = 35 years, two-sided Mann–Kendall test, P < 0.05) in TC, SV or BG. The dominant changes are Tree canopy gain with Short vegetation loss; Bare Ground gain with Short vegetation loss; Tree canopy gain with Bare Ground loss; Bare Ground gain with Tree canopy loss; 5, Short vegetation gain with Bare Ground loss; and Short vegetation gain with TC loss. Grey indicates areas with no significant change between 1982–2016.

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2.3.2 Projected Impacts of Extreme Events

Understanding of the large-scale drivers and the local to regional feedback processes that lead to extreme 13 events is still limited and projections of extremes and coincident or compounding events remain uncertain 14 (Prudhomme et al., 2014; Sillmann et al., 2017; Hao et al., 2018; Miralles et al., 2019). Extreme events are 15 challenging to model because they are by definition rare and often occur at spatial and temporal scales much 16 finer than the resolution of climate models (Sillmann et al., 2017; Zscheischler et al., 2018). Additionally, 17 the processes that cause extreme events often interact, as is the case for drought and heat events, and are 18 19 spatially and temporally dependent, for example, as is the case in soil moisture and temperature (Vogel et al., 2017). Understanding feedbacks between land and atmosphere also remains limited. For example, positive 20 feedbacks between soil and vegetation, or between evaporation, radiation and precipitation are important in 21 the preconditioning of extreme events such as heatwaves and droughts, increasing the severity and impact of 22 extreme events (Miralles et al., 2019). 23

Despite recent improvements in observational studies and climate modelling (Santanello et al., 2015;
Stegehuis et al., 2015; PaiMazumder and Done, 2016; Basara and Christian, 2018; Knelman et al., 2019), the
potential to quantify or infer formal causal relationships between multiple drivers and/or hazards remains

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report			
1 2			Kleinman et al., 2019; Miralles et al., 2019; rlying the response are difficult to identify			
3			g species and at different life stages, and an			
4	•		nd Kneeshaw, 2018). Additionally, hazards			
5			agricultural water demands, requiring more			
6	sophisticated modelling of the physic					
7		2 1	guides to how future events may evolve, teract to influence both system functioning			
8 9	and the climate state that have not be	e				
10	and the enhance state that have not be	en experienced in the past (s	see Chapter 11, wor Alto).			
11	2.3.3 Biologically Important Phys	ical Changes in Freshwate	r Systems			
12		-				
13			of biological organisation, from individual			
14			specific to freshwater systems which are			
15	not documented elsewhere in AR6 ar	e summarised here.	Co			
16	2221 Observed Changes in Theme	al Habitat and Omeron Anai	lability			
17 18	2.3.3.1 Observed Change in Thermo	u naduai ana Oxygen Avai	lability			
18	Since AR5, evidence for changes in t	emperatures of lakes and riv	vers has continued to increase. Global			
20			o 0.45°C per decade between 1970-2010,			
21						
22	exceeding SST trends of 0.09°C per decade between 1980–2017 (<i>robust evidence, high agreement</i>) (Figure 2.2; Schneider and Hook, 2010; Kraemer et al., 2015; O'Reilly et al., 2015; Woolway et al., 2020b).					
23			regions (O'Reilly et al., 2015) but more			
24	homogeneous than changes of deep v	vater temperature (Pilla et al	., 2020). Because temperature trends in			
25			hanges have occurred in the amounts of			
26	habitat available to aquatic organisms	s at particular depths and ter	nperatures (Kraemer et al., 2021).			
27						
28	Changes in river water temperatures					
29			006; Kaushal et al., 2010; Jurgelėnaitė et al.,			
30 31			and Pius, 2016). The more rapid increase in l winters (O'Reilly et al., 2015) can in part			
32			regions (<i>robust evidence, high agreement</i>)			
33	(Figure 2.2b; Screen and Simmonds,					
34	(g	,, _,				
35	Shifts in thermal regime: Since AR5	the trend that lake waters m	ix less frequently continues (Butcher et al.,			
36	2015; Adrian et al., 2016; Richardsor	et al., 2017; Woolway et al	l., 2017). This results from greater warming			
37	-		he loss of ice during winter which prevents			
38			dence, high agreement) (Adrian et al., 2009;			
39	Winslow et al., 2015; Adrian et al., 2	016; Schwefel et al., 2016; I	Richardson et al., 2017).			
40		· · · · · · · · · · · · · · · · · · ·	······································			
41			ixing cause a decrease in dissolved oxygen.			
42 43	1980 and 2017 (Jane et al., 2021). Th		ed by 4.1 and 16.8%, respectively between			
43 44			ing and fewer complete mixing events, with			

- 44 hypoxic conditions by more than 25% due to reduced winter mixing and fewer complete mixing events, with 45 strong repercussions on nutrient dynamics and loss of thermal habitat (*robust evidence, high agreement*)
- 46 (Straile et al., 2010; Zhang et al., 2015; Schwefel et al., 2016; Adrian et al., 2016; Kraemer et al., 2021).
- 47 48

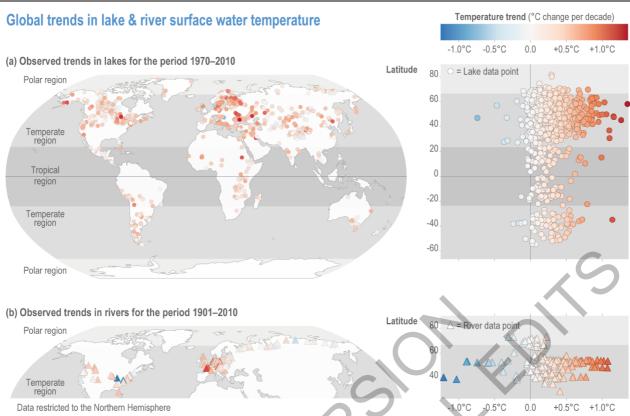


Figure 2.2: Observed global trends in lake and river surface water temperature. a) Left panel: Map of temperatures of lakes (1970-2010). b) Left panel: Map of temperatures of rivers (1901–2010). Note that the trends of river water temperatures are not directly comparable within rivers or directly comparable to lakes since time periods are not consistent across river studies. Right panels in a) and b) depict water temperature trends along a latitudinal gradient highlighting the above average warming rates in northern Polar Regions (polar amplification). Data sources for lakes: (O'Reilly et al., 2015; Carrea and Merchant, 2019; Woolway et al., 2020a; Woolway et al., 2020b). Data sources for rivers: (Webb and Walling, 1992; Langan et al., 2001; Daufresne et al., 2004; Moatar and Gailhard, 2006; Lammers et al., 2007; Patterson et al., 2010; Webb and Nobilis, 2007; Durance and Ormerod, 2009; Kaushal et al., 2010; Pekárová et al., 2011; Jurgelėnaitė et al., 2012; Markovic et al., 2013; Arora et al., 2016; Latkovska and Apsīte, 2016; Marszelewski and Pius, 2016; Jurgelėnaitė et al., 2017).

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2.3.3.2 Observed Changes in Water Level

Depending on how the intensification of the global water cycle affects individual lake water budgets, the 16 amount of water stored in specific lakes may increase, decrease, or have no substantial cumulative effect 17 (Notaro et al., 2015; Pekel et al., 2016; Rodell et al., 2018; Busker et al., 2019; Woolway et al., 2020b). The 18 magnitude of hydrological changes that can be assuredly attributed to climate change remains uncertain 19 (Hegerl et al., 2015; Gronewold and Rood, 2019; Kraemer et al., 2020). Attribution of water storage 20 variation in lakes due to climate change is facilitated when such variations occur coherently across broad 21 geographic regions and long timescales, preferably absent other anthropogenic hydrological influences 22 (Watras et al., 2014; Kraemer et al., 2020). There is increasing awareness that climate change contributes to 23 the loss of small temporary ponds, which cover a greater global area than lakes (Bagella et al., 2016). 24

Lakes fed by glacial meltwater are growing in response to climate change and glacier retreat (*robust evidence*, *high agreement*) (Shugar et al., 2020). Water storage increases on the Tibetan Plateau (Figure 2.3a) have been attributed to changes in glacier melt, permafrost thaw, precipitation and runoff, in part as a result of climate change (Huang et al., 2011; Meng et al., 2019; Wang et al., 2020a). *High confidence* in attribution of these trends to climate change is supported by long-term ground survey data and observations from the GRACE satellite mission (Ma et al., 2010; Rodell et al., 2018; Kraemer et al., 2020).

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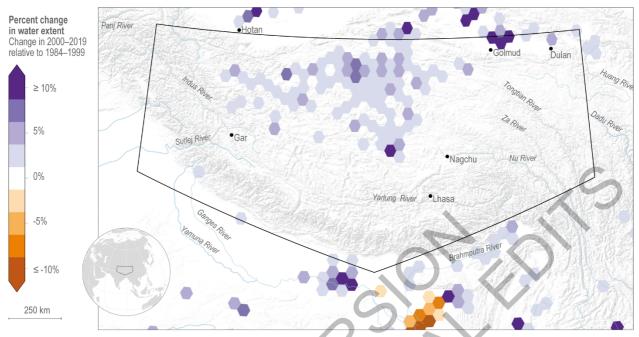
25

In the Arctic, lake area has increased in regions with continuous permafrost and decreased in regions where permafrost is thinner and discontinuous *(robust evidence, high agreement)* (see Chapter 4; Smith et al., 2005; Andresen and Lougheed, 2015; Nitze et al., 2018; Mekonnen et al., 2021).

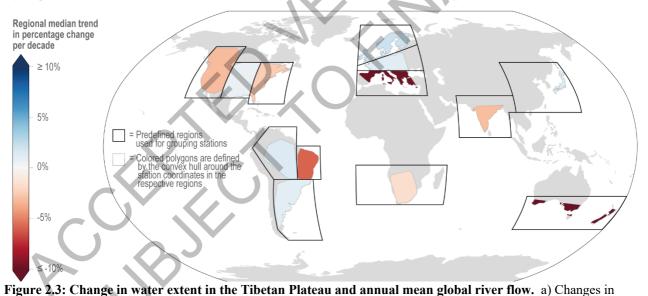
1 2

Change in water extent in the Tibetan Plateau & annual mean global river flow

(a) Change in water extent: Qinghai-Tibetan Plateau, Asia



(b) Regional median trend in annual mean river flow derived from 7,250 observatories around the world (period 1971-2010)



water storage on the Tibetan Plateau. Map of the Qinghai-Tibetan Plateau, Asia showing the percent change in surface

water extent from 1984-2019 based on LANDSAT imagery. Increases in surface water extent in this region are mainly

caused by climate-change mediated increases in precipitation and glacial melt (Source: EC JRC/Google; (Pekel et al.,

2016). b) Global map of the median trend in annual mean river flow derived from 7250 observatories around the world (period 1971–2010). Some regions are drying (northeast Brazil, southern Australia, and the Mediterranean) and others

evapotranspiration and alterations of the timing of snow accumulation and melt driven by rising temperatures (Source:

are wetting (northern Europe) mainly caused by large-scale shifts in precipitation, changes in factors that influence

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2.3.3.3 Observed Changes in Discharge

Analysis of river flows from 7,250 observatories around the world covering the years 1971 to 2010 and identified spatially complex patterns, with reductions in northeastern Brazil, southern Australia and the

Gudmundsson et al., 2021).

Mediterranean and increases in northern Europe (medium evidence, medium agreement) (Figure 2.3b; 1 Gudmundsson et al., 2021). More than half of global rivers undergo periodic drying that reduces river 2 connectivity (medium evidence, medium agreement). Increased frequency and intensity of droughts may 3 cause perennial rivers to become intermittent and intermittent rivers to disappear (medium evidence, medium 4 agreement), threatening freshwater fish in habitats already characterised by heat and droughts (Datry et al., 5 2016; Schneider et al., 2017; Jaric et al., 2019). In high altitude/latitude streams, reduced glacier and 6 snowpack extent, earlier snowmelt and altered precipitation patterns, attributed to climate change, have 7 increased flow intermittency (Vorosmarty et al., 2010; Siebers et al., 2019; Gudmundsson et al., 2021) 8 Patterns in flow regimes can be directly linked to a variety of processes shaping freshwater biodiversity, 9 hence any climate-change induced changes on flow regimes and river connectivity are expected to alter 10 species composition, as well as having societal impacts (See Chapter 3 of IPCC SR1.5; Bunn and 11 Arthington, 2002; Thomson et al., 2012; Chessman, 2015; Kakouei et al., 2018) 12 13

14 2.3.3.4 Observed Loss of Ice

Studies since AR5 have confirmed ongoing and accelerating loss of lake and river ice in the northern 16 hemisphere (robust evidence, high agreement) (Figure 2.4). In recent decades, systems have been freezing 17 later in winter and thawing earlier in spring, reducing ice duration by >2 weeks per year and leading to 18increasing numbers of years with loss of perennial ice cover, leading to intermittent ice-cover or even 19 absence of ice (Adrian et al., 2009; Kirillin et al., 2012; Paquette et al., 2015; Adrian et al., 2016; Park et al., 20 2016; Roberts et al., 2017; Sharma et al., 2019). The global extent of river ice declined by 25% between 21 1984 and 2018 (Yang et al., 2020). This trend has been more pronounced at higher latitudes, consistent with 22 enhanced polar warming (large geographic coverage) (Du et al., 2017). Empirical long-term and remote 23 sensing data gathered in an increasingly large number of freshwater systems supports very high confidence in 24 25 attribution of these trends to climate change. For declines of glaciers, snow and permafrost see AR6 WGII Chapter 4 and SROCC report. 26

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Global ice cover trends of lakes & rivers

(a) Future changes in lakes that experience intermittent winter ice cover in the Northern Hemisphere.

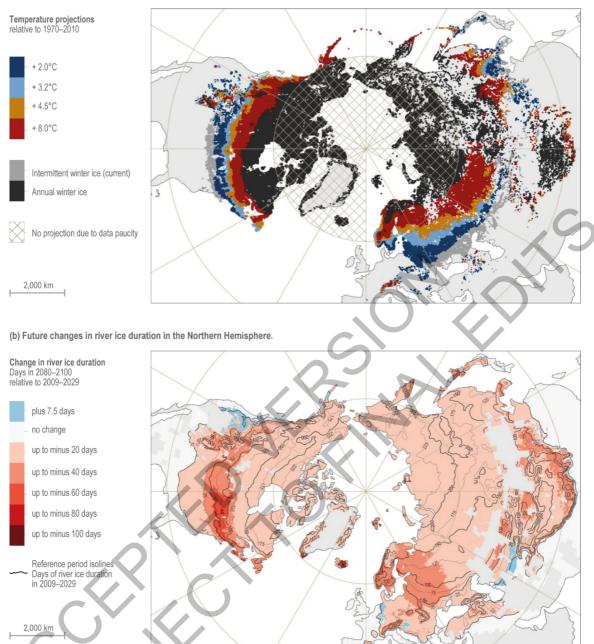


Figure 2.4: Global ice cover trends of lakes and rivers. a) Spatial distribution of current (light grey areas) and future 2 (coloured areas) Northern Hemisphere lakes that may experience intermittent winter ice cover with climate warming. 3 Projections were based on current conditions (1970-2010) and four established air temperature projections (Data 4 source: (Sharma et al., 2019). b) Spatial distribution of projected change in Northern Hemisphere river ice duration 5 under the RCP 4.5 emission scenario by 2080-2100 relative to the period 2009-2029. White areas refer to rivers without 6 ice cover in the period 2009-2029 (zero days). Reference period isolines indicate river ice duration in the period 2009-7 2029. Coloured areas depict loss of ice duration in days. Blue areas depict a projected increase in river ice duration. 8 Grey land areas indicate a lack of Landsat-observable rivers (Data source: Yang et al., 2020). 9

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2.3.3.5 Extreme Weather Events and Freshwater Systems

Since AR5, numerous drastic short-term responses have been observed in lakes and rivers, both to expected
seasonal extreme events and to unexpected supra-seasonal extremes extending over multiple seasons.
Consequences for ecosystem functioning are not well understood (Bogan et al., 2015; Death et al., 2015;
Stockwell et al., 2020). Increasing frequencies of severe floods and droughts attributed to climate change are
major threats for river ecosystems (Peters et al., 2016; Alfieri et al., 2017). While extreme floods cause

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massive physical disturbance, moderate floods can have positive effects, providing woody debris that 1 contributes to habitat complexity and diversity, flush fine sediments, dissolving organic carbon and 2 providing important food sources from terrestrial origins (Peters et al., 2016; Talbot et al., 2018). Droughts 3 reduce river habitat diversity and connectivity, threatening aquatic species, especially in deserts and arid 4 regions (Bogan et al., 2015; Death et al., 2015; Ledger and Milner, 2015; Jaric et al., 2019). 5 6 Rivers already stressed by human activities such as urban development and farming on floodplains are prone 7 to reduced resilience to future extreme events (medium confidence) (Woodward et al., 2016; Talbot et al., 8 2018). Thus, potential for floods to become catastrophic for ecosystem services are exacerbated by land-use 9 changes (Peters et al., 2016; Talbot et al., 2018). However, biota can recover rapidly from extreme flood 10 events if river geomorphology is not reformed. If instream habitat is strongly affected, recovery, if it occurs, 11 takes much longer, resulting in decline in biodiversity (medium confidence) (Thorp et al., 2010; Death et al., 12 2015; Poff et al., 2018). 13 14 However, not all extreme events will have a biological impact, depending in particular on the timing, 15 magnitude, frequency of events and the antecedent conditions (Bailey and van de Pol, 2016; Stockwell et al., 16 2020; Jennings et al., 2021; Thayne et al., 2021). For instance, an extreme wind event may have little impact 17 on phytoplankton in a lake, which was fully mixed prior to the event. Conversely, storm effects on 18phytoplankton communities may compound when lakes are not yet recovered from a previous storm or if 19 periods of drought alternate with periods of intense precipitation (limited evidence) (Leonard, 2014; 20 Stockwell et al., 2020). 21 22 In summary, extreme events (heat waves, storms, loss of ice) affect lakes in terms of water temperature, 23 water level, light, oxygen concentrations and nutrient dynamics, that in turn affect primary production, fish 24 communities and greenhouse gas emissions (high confidence). These impacts are modified by levels of solar 25 radiation, wind speed and precipitation (Woolway et al., 2020a). Droughts have negative impact on water 26 quality in streams and lakes by increasing water temperature, salinity, the frequency of algal blooms and 27 contaminant concentrations, and reducing concentrations of nutrients and dissolved oxygen (medium 28 confidence) (Peters et al., 2016; Alfieri et al., 2017) (Woolway et al., 2020a). Understanding how these 29 pressures subsequently cascade through freshwater ecosystems will be essential for future projections of 30 their resistance and resilience towards extreme events (Leonard, 2014; Stockwell et al., 2020). See Table 31 SM2.1 for specific examples of observed changes. 32

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34 2.3.3.6 Projected Changes in Physical Characteristics of Lakes and Rivers

Given the strength of relationship between past GAST and warming trends at lake surfaces (Figure 2.2; 36 section 2.3.3.1), and projected increases in heatwaves, surface water temperatures are projected to continue 37 to increase (Woolway et al., 2021). Mean May to October lake surface temperatures in 46,557 European 38 lakes were projected to be 2.9°, 4.5°, and 6.5°C warmer by 2081-2099 compared to historic (1981-1999) 39 under RCPs 2.0, 6.0, and 8.5, respectively (Woolway et al., 2020a). Under RCP 2.6, average lake heatwave 40 intensity increases from 3.7° to 4.0°C and average duration from 7.7 to 27.0 days, relative to the historical 41 period (1970-1999). For RCP 8.5, warming increases to 5.4°C and duration increases dramatically to 95.5 42 days (medium confidence) (Woolway et al., 2021). 43

- Worldwide alteration of lake mixing regimes in response to climate change are projected (Kirillin, 2010).
 Most prominently, monomictic lakes–undergoing one mixing event in most years–will become permanently
 stratified, while lakes that are currently dimictic–mixing twice per year–will become monomictic by 20802100 (*medium confidence*) (Woolway and Merchant, 2019). Nevertheless, predicting mixing behavior
 remains an important challenge and attribution to climate change remains difficult (Schwefel et al., 2016;
 Bruce et al., 2018).
- 51

Under climate projections of 3.2°C warming, 4.6% of the ice covered lakes in the northern hemisphere could switch to intermittent winter ice cover (Figure 2.4a; Sharma et al., 2019). Unfrozen and warmer lakes lose more water to evaporation (Wang et al., 2018b). By 2100, global annual lake evaporation will increase by 16%, relative to 2006-2015, under RCP 8.5 (Woolway et al., 2020b). Moreover, melting of ice decreases the ratio of sensible to latent heat flux, thus channelling more energy into evaporation (*medium confidence*)

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1 2	(Wang et al., 2018b). Between 2009-202 by 7.3 and 16.7 days under the RCP 4.5		ge river ice duration is projected to decline 4b; Yang et al., 2020).
3			
4	Projections of lake water are limited by	-	
5			ertainty is reflected in the widely divergent
6	projections for lake water storage in resp		
7			2; Notaro et al., 2015). Selecting models that
8 9	perform well when comparing hindcaste water storage projection uncertainty (Ar	1	r storage variation often does little to reduce This wide range of potential changes
10		e i i	projected changes in the global water cycle
11			ter and their implications on water quality
12			the role of weather and climate extremes on
13	the global water cycle, see Chapter 11 o		
14			
15	In summary, with ongoing climate warn	ning and an increase in the	he frequency and intensity of extreme
16			d shifts in thermal regime, are projected to
17	continue (high confidence).		
18			
19			
20	[START CROSS-CHAPTER BOX EX]	FREMES HERE]	
21			
22	Cross-Chapter Box EXTREMES: Rai		Extremes for Marine, Terrestrial,
23	Freshwater and Polar Natural S	Systems	
24		0.	
25	Authors: Rebecca Harris (Australia, Cha		
26	(Germany, Chapter 2), Jörn Birkmann (
27			r 3), Guy Midgley (South Africa, Chapter
28			2), Dieter Piepenburg (Germany, Chapter
29			apter 3), Björn Rost (Germany, Chapter 3,
30	CCP6), David Schoeman (Australia, Ch	apter 3), Maarten van Aa	aist (The Netherlands, Chapter 16).
31 32	Introduction	\vee ()	
32 33	Innounction		
33 34	Extreme events are now causing profou	nd negative effects acros	s all realms of the world (marine, terrestrial,
35			; WGII AR6 Section 2.3.1, 2.3.2, 2.3.3.5,
36			becies distributions, local extirpations and
37			ts in the composition, structure and function
38		-	er, 2018; Harris et al., 2018a; Maxwell et al.,
39			cations for ecosystems and the services they
40			ent cycling and water quality, with knock-on
41		d other natural resources	(Kaushal et al., 2018; Heinze et al., 2021;
42	Pörtner et al., 2021).		
43			
44			es projected for mean conditions for 2100,
45	regardless of emissions scenario (Figure		TREMES.1). This has collapsed the timeline
11	organisms and natural communities have	a to goolimate or adopt to	a climate change (moduum contidonce)

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- 46 organisms and natural communities have to acclimate or adapt to climate change (*medium confidence*).
 - Consequently, rather than having decades to identify, develop and adopt solutions, there is now an urgent
 need to build resilience and assist recovery following extreme events.
 - 49
 - Recent extremes highlight characteristics that enable natural systems to resist or recover from events, helping
 natural resource managers to develop solutions to improve resilience of natural communities and identify
 limits to adaptation (Bergstrom et al., 2021).
- 53

54 *Marine Heatwaves*

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Consensus is emerging that anthropogenic climate change has significantly increased the likelihood of recent
 marine heat waves (MHWs) (*medium confidence*) (WGI AR6, Chapter 9; Oliver et al., 2018). A widespread

MHW occurred in the NE Pacific between 2013–2015, with upper ocean temperature anomalies of up to 1 6.2°C relative to 2002-2012 (Gentemann et al., 2017). This event, termed the "Blob", enhanced surface 2 water stratification, decreasing nutrient supply, primary and community production, leading to widespread 3 changes to open ocean and coastal ecosystems, with geographical shifts of key species across trophic levels, 4 mass strandings of marine mammals, seabird mortalities and closures of commercially important fisheries 5 (Cavole et al., 2016; Piatt et al., 2020). The heatwave reappeared in 2019 ("Blob 2.0") (Amaya et al., 2020), 6 with similarly high temperature anomalies extending from Alaska to California, but the ecological effects of 7 this event are expected to differ because the Blob originated in winter, and Blob2.0 intensified in summer 8 (Amaya et al., 2020). Modelling suggests rapid shifts in the geographic distributions of important fish 9 species in response to MHWs (Cheung & Frölicher, 2020), with projected decreased biomass and 10distributional shifts of fish at least four times faster and larger than the effects of decadal-scale mean changes 11 throughout the 21st century under RCP8.5 (high confidence) (Cheung & Frölicher, 2020). Marine heatwaves 12 can also dramatically increase CH4 emissions from oceans, a significant positive feedback on global 13 warming (See also Chapter 3; Borges et al., 2019). 14 15 The Arctic region is warming more than twice as fast as the global mean, and polar organisms and 16

ecosystems are likely to be particularly vulnerable to heatwaves due to their specific thermal niches and 17 physiological thresholds and the lack of poleward 'refugia' (high confidence). The consequences of MHWs 18are exacerbated by concomitant sea-ice melting and freshening of surface waters, leading to secondary 19 effects due to osmotic stress and failing pH homeostasis. Since sea-ice associated organisms are often critical 20 components of polar food chains, cascading effects up to top predators are expected. In 2015–2016 a MHW 21 occurred in the Gulf of Alaska/Bering Sea (Walsh et al., 2018) which was unprecedented in terms of surface 22 temperatures and ocean heat content, geographical extent, depth range and persistence, impacting the entire 23 marine food web. Persistent warming favoured some phytoplankton species and triggered one of the largest 24 algal blooms recorded in this region, with concomitant oyster farm closures due to uncommon paralytic 25 shellfish poisoning events (Walsh et al., 2018). There were also massive die-offs of common murres (Uria 26 aalge) and puffins (Fratercula cirrhata), attributed to starvation resulting from warming-induced effects on 27 food supply (Jones et al., 2019). A 2017 survey found a 71% decline in abundance of Pacific cod (Gadus 28 *macrocephalus*) since 2015, likely due to an increase in metabolic demand and reduced prey supply during 29 the MHWs (Barbeaux et al., 2020). 30

32 Terrestrial Heatwaves

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33 Heatwaves are now regularly occurring that exceed the physiological thresholds of some species, including 34 birds and other small endotherms such as flying-foxes (high confidence) (Sections 2.4.2.2, 2.4.2.6). 35 Heatwaves in Australia, North America and southern Africa have caused mass mortality events due to lethal 36 hyperthermia and dehydration (Saunders et al., 2011; Conradie et al., 2020; McKechnie et al., 2021), 37 reducing fitness (du Plessis et al., 2012; Andrew et al., 2017; Sharpe et al., 2019; van de Ven et al., 2019; 38 van de Ven et al., 2020), breeding success and recruitment (Kennedy et al., 2013; Wiley and Ridley, 2016; 39 Ratnayake et al., 2019) and affecting daily activity and geographic distributions (Albright et al., 2017). They 40 also place enormous demands on wildlife management agencies and pose human health risks (Welbergen et 41 al., 2008). 42 43

Recent mortality events affected 14 species of bird and fruit bats (Epomophorus wahlbergi) in South Africa 44 when maximum air temperatures exceeded 43–45°C in 2020 (McKechnie et al., 2021). Passerine birds seem 45 more vulnerable to lethal hyperthermia due to the relative inefficiency of panting to lose heat (McKechnie et 46 al., 2021) and their small size, as heat tolerance generally increases with body mass (McKechnie et al., 47 2017). Several mass mortality events of flying-foxes (Pteropus poliocephalus, P. alecto) have occurred in 48 49 eastern Australia when maximum air temperatures exceeded 42°C (Welbergen et al., 2008). Nineteen such events occurred between 1994 and 2008, compared to three events prior to 1994. In January 2002, maximum 50 temperatures exceeded the 30-year average mean daily maximum by up to 16.5°C and killed more than 3500 51 individuals (Welbergen et al., 2008). In 2014, an estimated 45,500 flying-foxes died in a single day, when 52 average maximum temperatures were 8°C or more above average (Meteorology, 2014). Drought compounds 53 the impacts, as mortality increases when water availability is low (Welbergen et al., 2008; Mo and Roache, 54 2020; McKechnie et al., 2021). 55

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 1
 Antarctica encountered its first recorded heatwave in 2020. Record-high temperatures occurred in East

Antarctica (Robinson et al., 2020), with a maximum (9.2°C) temperature \sim 7°C above the mean maximum, and minimum temperatures > 0°C. Record-high temperatures (18.3°C) were also recorded in West Antarctica (Robinson et al., 2020). It is too soon to know the impacts on polar life, but such abrupt heating is expected to have wide-ranging effects on biota, from flash-flooding and dislodgement of plants, to excess melt waters supplying moisture to arid polar ecosystems (CCP Polar). Heatwaves in Siberia in 2016, 2018 and 2020,

with air temperature anomalies >6°C, were associated with extensive wildfires, pest infestations and melting
 permafrost (Overland and Wang, 2021).

10 Freshwater Extremes

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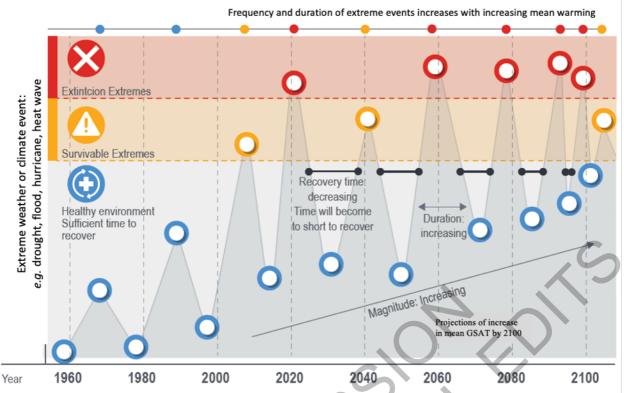
Heatwaves, storms and floods affect the thermal regime and biogeochemical functioning of lakes and rivers 12 (Woolway and Merchant, 2017; Vicente-Serrano et al., 2020). Extreme heatwaves lead to abnormally high 13 water temperatures (Till et al., 2019) and reduce mixing of lakes (Woolway et al., 2021), causing a decrease 14 in oxygen and deep-water oxygen renewal (Zhang et al., 2015). Ectotherms such as fish and invertebrates are 15 particularly susceptible to such temperature and oxygen stress (Stoks et al., 2014). Their metabolic demands 16 increase with rising temperature and suitable habitat is eroded due to both high temperatures and lower 17 oxygen concentrations in lakes and rivers. Till et al. (2019) attributed 502 fish kill events in Wisconsin lakes 18 (USA) to warmer summers in lakes that experienced abnormally high water temperatures. Such events are 19 predicted to double by 2041-2059 and increase fourfold by 2081-2099 compared to historical levels (Till et 20 al., 2019). This anticipated increase in die-offs may facilitate warm-water fish species displacing cool-water 21 species (Hansen et al., 2017; Jennings et al., 2021). Floods mobilise nutrients and sediment, and aid dispersal 22 of invasive species in rivers (Death et al., 2015), while drought extremes reduce river connectivity, 23

threatening biodiversity in rivers (Section 2.3.3.5; Tickner et al., 2020).

26 Learnings from Recent Extremes

27 These examples show that the impact of an extreme event is a function of its characteristics and those of the 28 exposed ecosystem. The timing, frequency, absolute magnitude and geographic extent of the extreme event, 29 relative to antecedent conditions, the life-cycle, resistance and resilience of the natural community, all 30 determine the biological response (Figure Cross-Chapter Box EXTREMES.2; Hillebrand et al., 2018; Gruber 31 et al., 2020). Impacts appear to be greater when extreme events occur more frequently, particularly when the 32 interval between events is insufficient to allow recovery to previous population sizes (e.g. frequent fire, coral 33 bleaching), or coincides with vulnerable life cycle stages, even when populations are adapted to cope with 34 such disturbances. Events occurring over large spatial areas reduce the potential for recolonisation from 35 nearby populations (e.g. regional droughts causing widespread declines). Often the magnitude of extreme 36 events exceeds historical levels, so organisms are less likely to be adapted to them, particularly when several 37 extremes coincide (e.g. high water temperatures, drought) (Duke et al., 2017). When hazards occur 38 simultaneously (compound events), impacts of extremes can be substantially aggravated, triggering 39 cascading effects in ecosystems (Gruber et al., 2020). 40

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1 Figure Cross-Chapter Box EXTRMES.1: A conceptual illustration of how extinction risk is affected by changes in 2 the frequency, duration and magnitude of extreme weather or climate events (e.g., drought, fire, flood, heatwaves). 3 Many organisms have adapted to cope with long-term and short-term climate variability, but as the magnitude and 4 frequency of extreme events increases, superimposed on the long-term climate trend, the threshold between survivable 5 extreme weather events (yellow) and extremes that have high risk of causing population or species extinctions (red) is 6 7 crossed more frequently. This can lead to local extinction events with insufficient time between to enable recovery, resulting in long-term, irreversible changes to the composition, structure and function of natural systems. When the 8 9 extreme event occurs over a large area relative to the distribution of a species (e.g., a hurricane impacting an island which is the only place a given species occurs), a single extreme event can drive the global extinction of a species. 10

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Several characteristics of natural systems are associated with greater vulnerability to extreme events (Figure 13 Cross-Chapter Box EXTREMES.2), knowledge of which can inform solutions to build resilience and aid 14 recovery (Robinson et al., 2020). Resilience can be built prior to an event by minimising additional 15 disturbances, such as water extraction from river systems, pollution of aquatic systems, land-use change and 16 fragmentation. Managing landscapes to reduce fragmentation and increase habitat extent, connectivity and 17 18 heterogeneity, by increasing the number and extent of reserves, may provide local refugia from extreme events and enhance post-event recolonisation, but may be less effective for marine systems (Chapter 3, 19 Section 3.6). Maintaining taxonomic, phylogenetic and functional diversity is important, as more diverse 20 systems may be more stable in the face of disturbance (Pimm, 1984; García-Palacios et al., 2018). 21

22 Several characteristics increase vulnerability: low or narrow thermal tolerances, high habitat specificity, low 23 dispersal ability, long generation times, low competitive ability and lifecycle constraints that limit recovery 24 or recolonisation. Populations living close to one or more limiting factors near range edges are also 25 vulnerable (Arafeh-Dalmau et al., 2019). Understanding these characteristics can inform management 26 intervention to aid recovery following an extreme event. For instance, knowledge of the flying-fox's 27 physiological temperature threshold led to successful interventions, including misting of populations to 28 reduce mortality (Mo and Roache, 2020) and the development of a 'heat stress forecaster', an online tool 29 which uses weather forecasts to identify roosts at risk of extreme heat events (Ratnayake et al., 2019). This 30 early-warning system increases the preparedness of wildlife management and conservation agencies, 31 enabling efficient allocation of management resources towards locations that are likely to be most affected. 32 Monitoring following extreme events can help identify immediate impacts and the potential for cascading 33 interactions, such as changes to competitive interactions following range shifts, impacts on freshwater 34 ecosystems following wildfires and the spread of invasive species. Ongoing monitoring of recovery and 35 effectiveness of management intervention is important, focussing on habitat-forming species (eg. kelp, 36

The acute impacts of extreme events, in addition accelerating and amplifying the biological effect globally and in all realms where life exists. Extr systems to adapt and impeding our ability to ide the urgent need to mitigate global greenhouse g impacts on natural systems (Díaz et al., 2020).	ts of climate change. This amplific reme events are compressing the ti- entify, develop and adopt solutions	cation is being observed meline available for na . Recent events highlig
Characteristics associated with vulnerability to	Solutions to build resilience and	d aid recovery
extreme events	Pre event	Post event
 Small habitat patches High fragmentation Inhospitable matrix 	Reserve/protect to maintain diversity and extent of natural system Increase connectivity to improve recolonisation and recovery	1º
 High disturbance levels Low diversity Habitat 	Minimise disturbance/other stressors (e.g. water extraction for rivers; pollution of waterways and ocean); Maximise heterogeneity to provide refuges (thermal and structural); Manage invasive species	Monitor to assess recovery, identify management requi Assess success of managem
 Population at range limit Low genetic diversity Low competitive ability Poor dispersal ability Narrow thermal tolerance Long generation times 	Minimise population stresses (e.g. hunting, fisheries quotas); Identify vulnerable species and plan (e.g. save seed stock) Forecast events and manage stress where possible (e.g. misting of flying-fox colonies to reduce heat)	 Identify intervention option translocation of warm – ada genotypes; reseeding; remo invasive or range shifting sp Assess cost, likelihood of intervention success Assess feasibility of ongoing management objectives

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Figure Cross-Chapter Box EXTREMES.2: Characteristics of natural systems that affect vulnerability and help identify solutions – both prior to and after extreme events - to build resistance, resilience and recovery.

1617 [END CROSS-CHAPTER BOX EXTREMES HERE]

2.4 Observed Impacts of Climate Change on Species, Communities, Biomes, Key Ecosystems and their Services

2.4.1 Overview

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Global meta-analyses of terrestrial systems in AR3 and AR4 concentrated on long time frames (>20 years) and findings from relatively undisturbed areas, where confidence in attributing observed changes to climate change is high. Recent global and regional meta-analyses (AR5 and later) have been broader, including data from degraded and disturbed areas and studies with shorter time frames (Tables 2.2a,b).

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- By the time of AR5, >4000 species with long-term observational data had been studied in the context of climate change (Parmesan, 2006; Parmesan and Hanley, 2015). Since then, hundreds of new studies have been added, leading to higher confidence in climate change attribution (Table 2.2; Scheffers et al., 2016; Wiens, 2016; Cohen et al., 2018; Feeley et al., 2020). Freshwater habitats have been under-represented in
- prior reports, but new long-term data sets, coupled with laboratory and field experiments, are improving our
- ³⁵ understanding and this assessment stresses observations from lakes and streams. As numbers of studies
- ³⁶ increase and data is increasingly extracted from areas with high LULCC, attribution is more difficult as

habitat loss and fragmentation are known major drivers of changes in terrestrial and freshwater species

(IPBES) (Gardner and Finlayson, 2018; Grill et al., 2019; Zarfl et al., 2019; Tickner et al., 2020).

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2.4.2 Observed Responses to Climate Change by Species and Communities (Freshwater and Terrestrial)

2.4.2.1 Observed Range Shifts Driven by Climate Change

7 Poleward and upward range shifts were already attributable to climate warming with high confidence in 8 AR5. Publication of observed range shifts in accord with climate change have accelerated since AR5 and 9 strengthened attribution. Ongoing latitudinal and elevational range shifts driven by regional climate trends 10 are now well-established globally across many groups of organisms, and attributable to climate change with 11 very high confidence due to very high consistency across a now very large body of species and studies and 12 in-depth understanding of mechanisms underlying physiological and ecological responses to climate drivers 13 (Table 2.2; Table 2.3, Table SM2.1; Pöyry et al., 2009; Chen et al., 2011; Grewe et al., 2013; Gibson-14 Reinemer and Rahel, 2015; MacLean and Beissinger, 2017; Pacifici et al., 2017; Anderegg et al., 2019b). 15 Range shifts stem from local extinctions along warm-range-boundaries (Anderegg et al., 2019b), as well as 16 from colonisation of new regions at cold-range-boundaries (Ralston et al., 2017). 17

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Many studies since AR4 have tended not to be designed as attribution studies, particularly recent large scale, 19 multispecies meta-analyses. That is, all data available were included in such studies (from both undisturbed 20 and from highly degraded lands, and including very short term datasets of <20 years) with little attempt to 21 design the studies to differentiate effects of climate change from effects of other potential confounding 22 variables. These studies tended to find greater lag and lower proportion of species changing in directions 23 expected from climate change, with authors concluding that LULCC, particularly habitat loss and 24 fragmentation, was impeding wild species from effectively tracking climate change (Lenoir and Svenning, 25 2015; Rumpf et al., 2019; Lenoir et al., 2020). 26

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Unprecedented outbreaks of spruce beetles occurring from Alaska to Utah in the 1990s were attributed to warm weather that, in Alaska, facilitated a halving of the insect's life cycle from two years to one (Logan et al., 2003). Milder winters and warmer growing seasons were likewise implicated in poleward range expansions and increasing outbreaks of several forest pests (Weed et al., 2013), leading to the current prediction that 41% of major insect pest species will increase their damage further as climate warms, and only 4% will reduce their impacts, while the rest will show mixed responses (Lehmann et al., 2020).

During their range shifts, forest pests remain climate-sensitive. For example, the distribution of Western 35 Spruce Budworm is limited at its warm range edges by adverse effects of mild winters on overwinter 36 survival, and at its cool range limits by ability to arrive at a cold-resistant stage before winter arrives 37 (Régnière and Nealis, 2019). We might therefore expect tree mortality from insect outbreaks to be most 38 severe in sites climatically less suitable for the plants, where plants would be under more stress. However, 39 (Jaime et al., 2019), using separate SDMs (MaxEnt) for the insects and plants, found that mortality of Scots 40 Pine from bark beetles was highest in sites most climatically suitable for the trees as well as for the insects. 41 In a study of tree mortality in California, bark beetles selectively killed highly-stressed fir trees but killed 42 pines according to their size, irrespective of stress status (Stephenson et al., 2019). 43

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Range shifts in a poleward and upward direction, following expected trajectories given the local and regional 45 climate trends, are strongly occurring in freshwater fish populations in North America (Lynch et al., 2016b), 46 Europe (Comte and Grenouillet, 2013; Gozlan et al., 2019) and Central Asia (Gozlan et al., 2019). Cold 47 water fish, such as coregonids and smelt have been negatively affected at the equatorial borders of their 48 distributions (Jeppesen et al., 2012). Upward elevational range shifts in rivers and streams have been 49 observed. Systematic shifts towards higher elevation and upstream were found for 32 stream fish species in 50 France following regional variation in climate change (Comte and Grenouillet, 2013). Bull trout (Salvelinus 51 confluentus) in Idaho (USA), were estimated to have lost 11–20% (8–16% decade⁻¹) of the headwater stream 52 lengths necessary for cold water spawning and early juvenile rearing, with the largest losses occurring in the 53 coldest habitats (Isaak et al., 2010). Range contractions of the same species have been found in the Rocky 54 Mountain watershed (Eby et al., 2014). Likewise, the distribution of the stonefly Zapada glacier, endemic to 55 alpine streams of Glacier National Park in Montana (USA), has been reduced over several decades by 56

- upstream retreat to higher, cooler sites as water temperatures have increased and glacial masses decreased
 (Giersch et al., 2015).
 The melting of glaciers has led to a change in water discharge associated with community turnover in
 glacier-fed streams (Cauvy-Fraunié and Dangles, 2019). For instance, glacier-obligate macroinvertebrates
 have started disappearing when glacial cover drops below approximately 50% (*robust agreement, high*
- *confidence*), reviewed in (Hotaling et al., 2017). For freshwater invertebrates, no meaningful trends have
 been detected in geographic extent or population size for most species (Gozlan et al., 2019).

An invasive freshwater cyanobacterium in lakes, *Cylindrospermopsis raciborskii*, originating from the tropics, has spread to temperate zones over the last few decades due to climate change-induced earlier increase of water temperature in spring (Wiedner et al., 2007), aided by a competitive advantage in eutrophic systems (Ekvall et al., 2013; Urrutia-Cordero et al., 2016).

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2.4.2.2 Observed Local Population and Global Species' Extinctions Driven by Climate Change

16 Disappearances of local populations within a species range are more frequent and better documented than 17 whole species' extinctions, and attribution to climate change is possible for sites with minimal confounding 18 non-climatic stressors. Changes of temperature extremes are often more important to these local extinction 19 rates than changes of mean annual temperature (see Sections 2.3.1, 2.3.2, 2.3.3.5, 2.4.2.6, Cross-chapter Box 20 EXTREMES this Chapter; Parmesan et al., 2013). In a study of 538 plant and animal species, sites with local 21 extinctions were associated with smaller changes of mean annual temperature but larger and faster changes 22 of hottest yearly temperatures than sites where populations persisted (Román-Palacios and Wiens, 2020). 23 Near warm range limits, 44% of species had suffered local extinctions. In both temperate and tropical 24 regions, sites with local extinction had greater increases in maximum temperatures than those without (T_{max} 25 increased 0.456°C and 0.316°C vs. T_{mean} increase of 0.153 °C and 0.061 °C for temperate (n=505 sites) and 26 tropical (n=76 sites), respectively, P < 0.001) (Román-Palacios and Wiens, 2020). 27

Wiens (2016) assumed that population extinctions were primarily driven by climate change when they 29 occurred at elevational or latitudinal "warm edge" range limits, and were in relatively undisturbed sites that 30 were stated by authors to be under increasing climatic stress. By this criterion, climate-caused local 31 extinctions were widespread among plants and animals, detected in 47% of 976 species examined. The 32 percentage of species suffering these extinctions was higher in the tropics (55%), than in temperate habitats 33 (39%), higher in freshwater (74%), than in marine (51%) or terrestrial (46%) habitats and higher in animals 34 (50%) than in plants (39%). The difference between plants and animals varied with latitude: in the temperate 35 zone a much higher proportion of animals than plants suffered range-limit extinctions (38.6% of 207 animal 36 species versus 8.6% of 105 plants, p < 0.0001) while at tropical sites local extinction rates were 37 (nonsignificantly) higher in plants (59% of 155 species) than in animals (52% of 349 species), the reverse of 38 their temperate zone relationship. Rates varied among animal groups, from 35% in mammals through 43% in 39 birds to 56% in insects and 59% in fish (Wiens, 2016). 40

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Freshwater population extinctions are mainly due to habitat loss, introduction of alien species, pollution, 42 over-harvesting (Gozlan et al., 2019; IPBES, 2019) and climate change induced epidemic diseases (Pounds 43 et al., 2006)(see Section 2.4.2.7.1). Climate warming particularly through intensification and severity of 44 droughts, contributes to the disappearance of small ponds, which hold rare and endemic species (Bagella et 45 al., 2016). Systematic data on the extent and biology of small ponds is, however, lacking at a global scale. 46 Extreme heat waves can lead to large local fish kills in lakes (see Section 2.3.3.5), when water temperature 47 and oxygen concentrations surpass critical thresholds, threatening cold water fish and amphibians 48 49 (Thompson et al., 2012). Evidence for a local extinction of some invertebrate species with a $1.4^{\circ}-1.7^{\circ}$ C rise in mean annual stream winter temperature from 1981–2005 was reported in (Abrahams et al., 2013). 50 Population declines of specialist species in glacier-fed streams, such as the non-biting midge Diamesa davisi 51 (Chironomidae), can be attributed to glacier retreat given climate change (Cauvy-Fraunié and Dangles, 52 2019), and the flatworm Crenobia alpina (Planariidae) has been reported as locally extinct in the Welsh Llyn 53 Brianne river (Durance and Ormerod, 2010; Larsen et al., 2018). 54

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Many high montane possums in Australia have low physiological tolerance to heatwaves, with death occuring due to heat-driven dehydration at temperatures exceeding 29° - 30° C for >4–5 hours over several

days (Meade et al., 2018; Turner, 2020), with major declines recorded for several species, and population extinctions at lower elevations, since the early 2000s (Chandler, 2014; Weber et al., 2021).
Two terrestrial and freshwater species have gone extinct, with climate change implicated as a key driver. The
cloud-forest-restricted Golden toad (<i>Incilius periglenes</i>) was extinct by 1990 in a nature preserve in Costa Rica, driven by successive extreme droughts. This occurred in the absence of chytridiomycosis infection,
caused by the fungal pathogen <i>Batrachochytrium dendrobatidis</i> (BD), verified during field censuses of
golden toad populations in the process of extinction and through genetic analyses of museum specimens,
although Bd was present in other frog species in the region <i>(medium evidence, high agreement)</i> (Pounds et
al., 1999; Pounds et al., 2006; Puschendorf et al., 2006; Richards-Hrdlicka, 2013). The interaction between
expansion of chytrid fungus globally and local climate change is implicated in the extinction of a wide range
of tropical amphibians (see Section 2.4.2.7.1 Case study 2 Chytrid fungus and climate change).
The Bramble Cay Melomys (Melomys rubicola), the only mammal endemic to the Great Barrier Reef,
inhabited a small (five hectare) low-lying (<3m high) cay in the Torres Strait Islands, Australia. Recorded
having a population size of several hundred in 1978, this mammal has not been seen since 2009 and was declared extinct in 2016 (Gynther et al., 2016). SLR, documented increases in storm surge and in tropical
cyclones, driven by climate change, led to multiple inundations of the island in the 2000s. Between 1998 and
2014, herbacious vegetation, the food resource for the BC Melomys, declined by 97% in area (from 2.2 ha
down to 0.065 ha), and from 11 plant species down to two (Gynther et al., 2016; Watson, 2016; Woinarski,
2016; Woinarski et al., 2017). The island was unihabited with few non-climatic threats, providing high
confidence in attribution of extinction of the BC Melomys to climate change-driven increases in frequency
and duration of island inundation (Turner and Batianoff, 2007; Woinarski et al., 2014; Gynther et al., 2016;
Watson, 2016; Woinarski et al., 2017).
In the IUCN Red List (IUCN, 2019), 16.2% of terrestrial and freshwater species (n=3,777 species) that are listed as endangered, critically endangered or extinct in the wild (n=23,251 species) list climate change or
severe weather as one of their threats.
severe weather as one of their threads.
[START FAQ2.1 HERE]
FAQ2.1: Will species go extinct with climate change and is there anything we can do to prevent it?
Climate shapes is alwayde posing major throats to bigdinaryity and the most will enable plants and animals
Climate change is already posing major threats to biodiversity and the most vulnerable plants and animals are likely to go extinct. If climate change continues to worsen, it is expected to cause many more species to
go extinct unless we take actions to improve the resilience of natural areas, through protection, connection
and restoration. We can also help individual species that we care most about by reducing the stress they are
under from other human activities, and even helping them move to new places as their climate space shifts
and they need to shift to keep up.
Climate change has already caused some species to go extinct, and is likely to drive more species to
extinction. Species have always gone extinct in the history of our planet but human activities causing climate
change are accelerating this process. For instance, recent research predicts that one-third of all plant and
animal species could be extinct by 2070 if climate change continues as it is. Species can adapt to some extent to these rapidly changing climate patterns. We are seeing changes in behaviour, dispersal to new areas as the
climate becomes more suitable, and genetic evolution. However, these changes are small, and adaptations
are limited. Species that cannot adapt beyond their basic climate tolerances (ability to survive extremes of
temperature or rainfall) or successfully reproduce in a different climate environment from what they have
evolved in, will simply disappear. In the Arctic for example, sea ice is melting and will likely disappear in
evolved in, will simply disappear. In the Arctic for example, sea ice is melting and will likely disappear in summer time within a century. This means that the animals that have evolved to live on sea ice - polar bears
evolved in, will simply disappear. In the Arctic for example, sea ice is melting and will likely disappear in

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Fortunately, there are some things we can do to help. We can take actions to assist, protect and conserve natural ecosystems and prevent the loss of our planet's endangered wildlife, such as:

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"Assisting" species' migration: This has many names, "assisted colonisation", "assisted translocation", 1 "assisted migration", "assisted movement". In effect, it is about helping endangered species to move to a new 2 area with a good habitat for them to survive. "Passive" assisted colonisation focuses on helping species move 3 themselves, whilst the most "active" form implies picking up individuals and transporting them to a new 4 location. This is different from re-introductions that are already a normal part of conservation programs. 5 Climate-driven translocations are moving plants or animals to an area where they have never lived 6 historically, a new location that is now suitable for them due to climate change. 7

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This active form of "assisted colonisation" has been controversial, because exotic species can become invasive when they are moved between continents or oceans. For example, no one would advocate moving polar bears to Antarctica, as they would likely fest on native penguins, thus causing another conservation problem. However, moving species only a few hundred kilometers avoids most adverse outcomes, and that is often all that is needed to help a wild plant or animal cope with climate change. In extreme cases, another type of assisted adaptation is to preserve species until we get climate change under control and can 14 reintroduce them to the wild. This might include moving them into zoos or into seed or frozen embryo

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banks.

17 Extending protected zones and their connectivity: Species' ability to move to new locations and track climate 18 change are very limited – in particular, when a habitat has been turned into a crop field or a city. To help 19 them move between their natural habitats, we can increase the connectedness of protected areas or simply 20 create small patches or corridors of semi-wild nature within a largely agricultural or inhabited region, that 21 encourages wildlife to move through an area, and in which they are protected from hunting and poisons. 22 Those semi-wild protected areas can be very small, like the hedgerows between fields in England, that 23 provide both habitats for many flowers, birds, insects and corridors to move between larger protected areas. 24 Alternatively, it can just be an abandoned field that is now growing "weeds" without pesticides, hunting or 25 farming. For instance, in the United States of America, private landowners get a tax break by making their 26 land a "wildlife conservation" area using no pesticide, not cutting weeds too often, putting up brush piles and 27 bird boxes for nesting, and providing a water source. 28 29

Assisting, protecting and conserving natural ecosystems would help enhance biodiversity overall as well as 30 already endangered species. Diverse plant and animal communities are more resilient to disturbances, 31 including climate change. A healthy ecosystem also recovers more quickly from extreme events, such as 32 floods, droughts and heat waves that are a part of human-driven climate change. Healthy ecosystems are 33 critical to prevent species' extinctions from climate change, but are also important for human health and 34 well-being, providing clean, plentiful water, cleaning the air, providing recreation and holiday adventures, 35 and making people feel happier, calmer and more contented. 36 37

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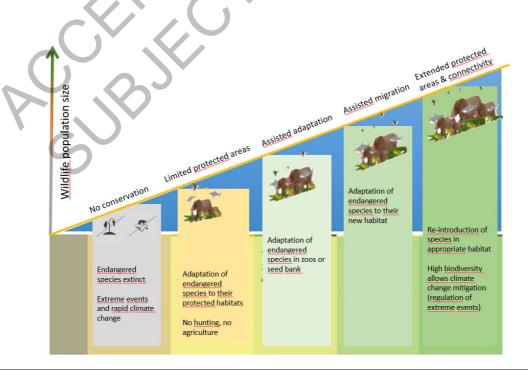


Figure FAQ2.1.1: Possible actions to assist protect and conserve natural ecosystems and prevent the loss of our planet's endangered wildlife in the face of continued climate change. (*Inspired by Natural Alliance website* © *Chris Heward/GWCT*)

[END FAQ2.1 HERE]

2.4.2.3 Observed Changes in Community Composition Driven by Climate Change

10 2.4.2.3.1 Overall patterns of community change

The most common type of community change takes the form of in situ decreases of cold-adapted species and increases of warm-adapted species (Bowler et al., 2017; Hughes et al., 2018; Kuhn and Gégout, 2019; Feeley et al., 2020). This process has lead to increases of species richness on mountaintops and decreased richness at adjacent lower elevations *(medium evidence, high agreement)* (Forister et al., 2010; Steinbauer et al., 2018). Observed shifts in community composition have consequences for species' interactions. Such indirect effects of climate change have been shown to often have greater impacts on species than direct effects of climate itself, particularly for higher level consumers (Ockendon et al., 2014).

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Like other responses, analyses indicated responses were lagging behind the change expected from regional 19 warming, and thereby accumulating 'climate debt.' Examples of climate debt, measured from community 20 composition changes, come from birds and butterflies in Europe (Devictor et al., 2012) and from lowland 21 forest herbaceous plants in France (Bertrand et al., 2011). The French study found that larger debts occurred 22 in communities with warmer baseline conditions and that some of the apparent debt stemmed from species' 23 ability to tolerate warming in situ. Geothermal streams have provided evidence about community structure 24 and ecosystem function in high temperatures. A study of 14 such habitats reported simplified the food-web 25 structures and shortened pathways of energy flux between consumers and resources (high confidence) 26

27 (O'Gorman et al., 2019).

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Prominent changes in freshwater community composition, such as increases in cyanobacteria and warm 29 tolerant zooplankton species, loss of cold water fish, gain in thermo-tolerant fish and macroinvertebrates and 30 gain in floating macrophytes, are occurring (medium evidence, high agreement, medium confidence) (Adrian 31 et al., 2016; Hossain et al., 2016; Short et al., 2016; Huisman et al., 2018; Gozlan et al., 2019). Changes in 32 relative species abundances, species composition and biodiversity due to warming trends and non-climate 33 driven changes are to be expected in lakes and rivers globally. However, thus far empirical evidence and 34 mechanistic understanding to inform modelling is too limited to draw general conclusions about the nature of 35 current and future climate change driven changes within entire food webs on a global scale (Urban et al., 36 2016). 37

39 2.4.2.3.2 Freshwater mechanistic drivers and responses

Physical changes in lakes (see Section 2.3.3) have affected primary production (see Section 2.4.4.5.2), algal 40 bloom formation and composition, zooplankton and fish size distribution and species composition (Urrutia-41 Cordero et al., 2017; Gozlan et al., 2019; Seltmann et al., 2019). Declines in abundance of cold-stenothermal 42 species (particularly Arctic charr, Salvelinus alpinus, coregonids and smelt) and increases in eurythermal 43 fish (e.g. the thermo-tolerant carp Cyprinus carpio, common bream, pike perch, roach and shad) have been 44 observed in northern temperate lakes associated with warming trends (high agreement, medium confidence) 45 (Jeppesen et al., 2012; Jeppesen et al., 2014). These changes increase predation pressure on zooplankton and 46 reduce grazing pressure on phytoplankton, which may result in higher phytoplankton biomass (De Senerpont 47 Domis et al., 2013; Jeppesen et al., 2014; Adrian et al., 2016). Reduction in lake mixing lowers the 48 concentration of nutrients in the epilimnion and may lead to higher silicon to phosphorous ratios negatively 49 affecting diatom growth (Yankova et al., 2017) or overall primary productivity (see Section 2.4.4.5.2). 50

In as study of 1,567 lakes across Europe and North America, (Kakouei, 2021) identified climate change as

⁵² the major driver of increases in phytoplankton biomass in remote areas with minimal LULCC. Greater

temperature variability can be more important than long-term temperature trends as a driver of zooplankton

biodiversity (Shurin et al. (2010). Reductions of winter severity attributed to anthropogenic climate change are increasing winter algal biomass, and motile and phototropic species at the expense of mixotrophic

species (Özkundakci et al., 2016; Hampton et al., 2017).

2 Tropical lakes are prone to loss of deep-water oxygen due to lake warming with negative consequences for

- their fisheries and their biodiversity (Lewis Jr, 2000; Van Bocxlaer et al., 2012). Many ancient tropical lakes
 (Malawi, Tanganyika, Victoria, Titicaca, Towuti and Matano) hold thousands of endemic animal species
- 5 (Vadeboncoeur et al., 2011).
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Observed climate-change effects on freshwater invertebrates are variable (Knouft and Ficklin, 2017). In 7 glacier-fed streams globally, climate change has caused community turnover and changes in abundances in 8 terms of increased generalist and decreased specialist species abundances (Lencioni, 2018; Cauvy-Fraunié 9 and Dangles, 2019). In turn, dragonflies in flowing waters, monitored during the warming period from 1988 10 through 2006 in Europe, did not show consistent changes in their distribution (Grewe et al., 2013), reviewed 11 in (Knouft and Ficklin, 2017). Long-term trends in species composition and community structure of stream 12 macroinvertebrates, specifically a general trend for decreases in species characteristic of cold, fast-flowing 13 waters and increases of thermophilic species typical of stagnant or slow-moving waters, have been attributed 14 to climate change (high agreement, high confidence) (Daufresne et al., 2007; Chessman, 2015). A study of 15 14 geothermal streams reported simplified food-web structures and shortened pathways of energy flux 16 between consumers and resources (O'Gorman et al., 2019). Macrophytes benefit from rising water 17 temperatures, but increased shading from increased phytoplankton biomass could offset this (see 2.5.3.6.2 18 for projections; Hossain et al., 2016; Short et al., 2016; Zhang et al., 2017a). 19

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22 2.4.2.3.3 Emergence of novel communities and invasive species

As climate change is increasing the movements of species into new areas, there is concern about how exotic 23 species are being impacted, either by becoming invasive or by already invasive species gaining even more 24 advantage over native species. Modeling predicts that effects of climate warming on food web structure and 25 stability favour success of invading species (Sentis et al., 2021). Both simulated warming experiments 26 (Zettlemoyer et al., 2019) and long-term observations (Losos et al., 2010) have found phenologies of exotic 27 species to respond more adaptively to warming than those of natives, and in the long-term observations the 28 success of exotics was attributed to their greater phenological responsiveness. In an expert assessment of the 29 future relative importance of different drivers of the impacts of biological invasions, climate change was 30 named as the most important driver in polar regions, the second most important in temperate regions (after 31 trade/transport) and the third most important in the tropics (after trade/transport and human 32 demography/migration) (Essl et al., 2020). 33

However, not all exotic species become invasive. As novel climate conditions develop, novel communities 34 made up of new combinations of species are emerging as populations and species adapt and shift ranges 35 differentially, not always with negative consequences (high confidence) (Dornelas et al., 2014; Evers et al., 36 2018; Teixeira and Fernandes, 2020). Novel communities differ in composition, structure, function and 37 evolutionary trajectories, as the proportion of specialists and generalists, native, introduced and range 38 shifting species changes and species interactions are altered, ultimately affecting ecosystem dynamics and 39 functioning (Lurgi et al., 2012; Hobbs et al., 2014; Heger and van Andel, 2019 Towards an Integrative). The 40 exact nature of novel communities is difficult to predict because species-level uncertainties propagate at the 41 community level due to ecological interactions (Williams and Jackson, 2007), but observations, experimental 42 mesocosms (Bastazini et al., 2021); and theoretical models (Lurgi et al., 2012; Sentis et al., 2021) provide 43 support that they will continue to emerge with climate change. 44 45

46 2.4.2.4 Observed Phenological Responses to Climate Change

47 With advances in remote sensing, quality and quantity of phenological data are rapidly increasing (Piao et 48 49 al., 2019). Since AR5, numbers of studies have increased substantially with consistent conclusions in response to warming, including advancement of spring events and lengthening of growing seasons in 50 temperate regions (through a combination of advancement of spring events and to a lesser extent, retardation 51 of autumn events) (robust evidence, high agreement) (Table 2.2, Table 2.3, Table SM2.1; Menzel et al., 52 2020). In the tropics, by contrast, precipitation changes have more strongly influenced phenology than 53 temperature changes (Cohen et al., 2018). A meta-analysis comparing observed phenological advances in 54 birds with expectations from warming of local climates concluded that the advances fell short of expectation 55 and that substantial phenological climate debt had been generated (Radchuk et al., 2019). 56 57

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Taxonomic groups have differed in their responses (Parmesan, 2007; Thackeray et al., 2010), and a few have 1 completely abstained from the general trends-for example, seabirds continue to breed with their pre-2 climate-change phenologies (Keogan et al., 2018). Newer reviews and analyses reveal differences in 3 responses among continents and across time intervals (Piao et al., 2019). Mean advance in days per decade 4 was 5.5 in China, and 3.0–4.2 in Europe but only 0.9 in North America (Piao et al., 2019). Mean values for 5 retardation of autumn leaf fall, which can be more influenced by photoperiod and less by temperature than 6 spring leaf-out, were 0.36 days per decade in Europe (Menzel et al., 2020), 2.6 days per decade in China and 7 around 3 days per decade in the USA (medium evidence, high agreement) (Piao et al., 2019). 8

9 The rapid rates of advance of spring events in the 1990s slowed down in the 2000s and stalled or even 10 reversed in some regions (Menzel et al., 2020). Wang et al. (2019) noted, from remote sensing, that during 11 the 'global warming hiatus' from 1998–2012, there were no global trends in either Spring green-up or 12 autumn colouring. Annual crops, for which timing is determined by farmers, were an exception. When 13 natural systems were advancing fast prior to 1998, farmers advanced more slowly, but during the natural 14 'hiatus', farmed crops advanced faster than wild plants and cultivated trees (Menzel et al., 2020). In a long 15 (67 year) European time series (Menzel et al., 2020), autumn leaf colouring showed delays attributed to 16 17 winter & spring warming in 57% of observations (mean delay 0.36 days per decade); spring & summer phenologies advanced in 89% of wild plants despite decreased winter chilling, with c.60% of trends 18 significant and 'strongly attributable' to winter & spring warming; and growing season length increased in 19 84% of cases (mean lengthening 0.26 days yr^{-1}) (Table 2.2). 20

21 Changes in freshwater systems are consistent with changes in terrestrial systems: earlier timing of spring 22 phytoplankton and zooplankton development and earlier spawning by fish, as well as extension of the 23 growing season are occurring (robust evidence, high agreement) (Adrian et al., 2009; De Senerpont Domis et 24 al., 2013; Adrian et al., 2016; Thackeray et al., 2016). Phenological changes in lakes have been related to 25 rising water temperatures, reductions of ice cover and prolongation of thermal stratification (increasing 26 evidence and agreement since AR5; very high confidence). Crozier and Hutchings (2014) reviewed 27 phenological changes in fish and documented that changes in the timing of migration and reproduction, age 28 at maturity, age at juvenile migration, growth, survival and fecundity were associated primarily with changes 29 in temperature. The median return time of Atlantic salmon among rivers in Newfoundland and Labrador 30 advanced by 12 to 21 days over the past decades, associated with overall warmer conditions (Dempson et al., 31 2017). 32

34 2.4.2.5 Observed Complex Phenological and Range Shift Responses

35 Early meta-analyses tested the straightforward hypotheses that warming should shift timing earlier and 36 ranges poleward. Once these trends had been established, exceptions to them became foci of study. For 37 example, some plants in northern regions of the northern hemisphere were retarding their spring flowering 38 instead of advancing it as expected with warming. These turned out to be species requiring vernalisation 39 (winter chilling) to speed spring development. For these plants, phenological changes result from combined 40 effects of advancement caused by spring warming and retardation caused by winter warming. Incorporating 41 this level of complexity into analyses revealed that a greater proportion of species were responding to 42 climate change than estimated under the simple expectation that warming should always cause advancement 43 (92% responding vs 72% from earlier analyses) (Cook et al., 2012). 44

- Animal species can show vernalisation equivalent to that in plants (Stålhandske et al., 2017). However, a semi-global meta-analysis across terrestrial animals failed to detect delaying effects of warming winters (Cohen et al., 2018). The same animal-based meta-analysis contrasted phenological changes in temperatezone animals, which are principally explained by changes of temperature, with those at lower latitudes, which follow changes of precipitation (Cohen et al., 2018).
- Vitasse et al. (2018), working with Alpine trees, found that phenological delay with increasing elevation had declined from 34 days per 1,000 m in the 1960s to 22 days per 1,000 m, greatly reducing the differences in timing between trees growing at different elevations. This reduction was greatest after warmer winters, suggesting winter warming as a principal cause of the overall trend.
- 56

45

1	Lian et al. (2020) observed that earner spring lear-out in the Northern Hemisphere is causing increases in
2	evapotranspiration that are not fully compensated by increased precipitation. The consequence is increased
3	soil moisture deficit in summer, expected to exacerbate impacts of heatwaves as well as drought stress. In
4	Arctic freshwater ecosystems, Heim et al. (2015) demonstrated the importance of seasonal cues for fish
	migration, which can be impacted by climate change due to reduced stream connectivity and fragmentation,
5	
6	earlier peak flows, and increased evapotranspiration.
7	
8	Precipitation has also been implicated in exceptions to the rule that ranges should be shifting to higher
9	elevations. In dry climates, increases of precipitation accompanying climate warming can facilitate
10	downslope range shifts (Tingley et al., 2012).
11	
12	Multiple responses can co-occur. Hällfors et al (Hällfors et al., 2021), in a study of 289 lepidoptera in
13	Finland, found 45% had either shifted their ranges northward or advanced their flight season with warming.
14	The 15% of species that did both (shifting northward by 113.1 km and advancing flight period by 2.7 days
	per decade, on average, over a 20 year period) had the largest population increases, and the 40% of species
15	
16	that showed no response had the largest population declines.
17	
18	
19	Table 2.2: Global Fingerprints of Climate Change Impacts across Wild Species. Updated from (Parmesan and Hanley,
20	2015). For each dataset, a response for an individual species or functional group was classified as (1) no response (no
21	significant change in the measured trait over time), (2) if a significant change was found, the response was classified as
22	either consistent or not consistent with expectations from local or regional climate trends. Percentages are approximate
23	and estimated for the studies as a whole. Individual analyses within the studies may differ. The specific metrics of
24	climate change analysed for associations with biological change vary somewhat across studies, but most use changes in
25	local or regional temperatures (e.g. mean monthly T or mean annual T), with some using precipitation metrics (e.g. total
26	annual rainfall). For example, a consistent response would be poleward range shifts in areas that are warming.
27	Probability (P) of getting the observed ratio of consistent: not consistent responses by chance was <10-13 for (Parmesan
28	and Yohe, 2003; Root et al., 2003; Root et al., 2005; Poloczanska et al., 2013) and was <0.001 for Rosenzweig 2008
29	(source=publication) (Parmesan and Yohe, 2003; Root et al., 2003; Root et al., 2005; Rosenzweig et al., 2008;
30	Poloczanska et al., 2013). Test were all binomial tests against p=0.5, performed by Parmesan.
31	[INSERT TABLE 2.2 HERE.]
32	
33	
34	2.4.2.6 Observed Changes to Physiology and Morphology Driven by Climate Change
35	
36	Impacts on species physiology in terrestrial and freshwater systems have been observed and attributed to
37	climate change (<i>medium confidence</i>), including changes in tolerances to high temperatures (Healy and
38	Schulte, 2012; Gunderson and Stillman, 2015; Deery et al., 2021), increased metabolic costs of living under
39	elevated temperatures (Scheffers et al., 2016) and shifts in sex ratios in species with temperature-dependent
40	sex determination (e.g. masculinisation of lizard populations (Schwanz and Janzen, 2008; Schwanz, 2016;
41	Edmands, 2021) and feminisation of turtle populations (Telemeco et al., 2009)). Skewed sex ratios can lead
42	to mate shortages, reduced population growth and adaptive potential, and increased extinction risk, because
	genetic diversity decreases as fewer individuals mate and heterozygosity is lost (Mitchell and Janzen, 2010;
43	
44	Edmands, 2021).
45	
46	Behavioural plasticity such as nest-site selection can provide a partial buffer from the effects of increasing
47	temperature, but there are environmental and physical limits to this plasticity (medium confidence)
48	(Refsnider and Janzen, 2016; Telemeco et al., 2017). Plasticity in heat tolerance (e.g. due to reversible
49	acclimation or acclimatisation) can also potentially compensate for rising temperatures (Angilletta Jr, 2009),
50	but ectotherms have relatively low acclimation in thermal tolerance and acclimation is expected to only
51	slightly reduce overheating risk in even the most plastic taxa (low confidence) (Gunderson and Stillman,
52	2015).
53	
54	Geographic variation in thermal tolerance plasticity is expected to influence species vulnerability and range
	chifts in response to alimete abange (Gundarson and Stillman 2015; Sun et al. 2021). In many estathering

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Lian et al. (2020) observed that earlier spring leaf-out in the Northern Hemisphere is causing increases in

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- shifts in response to climate change (Gunderson and Stillman, 2015; Sun et al., 2021). In many ectotherms,
- plasticity in thermal tolerance increases towards the poles, as thermal seasonality increases (Chown et al., 2004), contributing to higher vulnerability to warming in tropical organisms (*low confidence*) (Huey et al.,
- 2004), contributing to higher vulnerability to warming in tropical organisms (*low confidence*) (Huey et al.
 2009; Campos et al., 2021). Some species have evolved extreme upper thermal limits at the expense of

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plasticity, reflecting an evolutionary trade-off between these traits (Angilletta et al., 2003; Stillman, 2003). 1 The most heat-tolerant species, such as those from extreme environments, may therefore be at greater risk of 2 warming because of an inability to physiologically adjust to thermal change (low confidence) (Bozinovic et 3 al., 2011; Overgaard et al., 2014; Magozzi and Calosi, 2015). 4 5 Physiological changes have observable impacts on morphology, such as changes to body size (and length of 6 appendages) and colour changes in butterflies, dragonflies, and birds (medium confidence) (Galeotti et al., 7 2009; Karell et al., 2011), but trends are not always linear or consistent across realms, taxonomic groups or 8 geographic regions (Gotanda et al., 2015). Some morphological changes arise in response to environmental 9 changes rather than as the result of genetic adaptation or selection for an optimum body type. For example, 10 dietary changes associated with climate change have led to changes in chipmunk skull morphology (Walsh et 11 al., 2016). 12 13 Decreased body size has been suggested as a general response of species to climate change in freshwater 14 species given the temperature related constraints of metabolism with increasing body size. Reduced body 15 size in response to global warming has been documented for freshwater bacteria, plankton and fish, as well 16 as a shift towards smaller species (low confidence) (Daufresne et al., 2009; Winder et al., 2009; Jeppesen et 17 al., 2010; Crozier and Hutchings, 2014; Jeppesen et al., 2014; Farmer et al., 2015; Rasconi et al., 2015; 18Woodward et al., 2016). However, the lack of systematic empirical evidence in freshwaters and confounding 19 effects such as interactions between temperature, nutrient availability and predation limit generalisations 20 about body size effects (Pomati et al., 2020 Nutrients). 21 22 Evidence is weak for a consistent reduction in body size across taxonomic groups in terrestrial animals (low 23 confidence) (Siepielski et al., 2019). Decreased body size in warmer climates (as higher surface area to 24 volume ratios maximise heat loss) is expected based on biogeographic patterns such as Bergmann's Rule, 25 but both increases and decreases have been documented in mammals, birds, lizards and invertebrates and 26 attributed to climate change (Teplitsky and Millien, 2014; Gotanda et al., 2015; Gardner et al., 2019; Hill et 27 al., 2021). Contrasting patterns (increased body size) may be due to short-term modifications in selection 28 pressures (e.g. changes to predation and competition), variation in life histories or a result of interactions 29 with climate variables other than temperature (e.g. changes to food availability with rainfall changes) and 30 other disturbances (Yom-Tov and Yom-Tov, 2004; Gardner et al., 2019; Wilson et al., 2019) or body size 31 measurements (linear vs. volumetric dimensions) (Salewski et al., 2014). 32 33 Several lines of evidence suggest evolution of melanism in response to climate change (low confidence), 34 with colour changes associated with thermoregulation being demonstrated in butterflies (Zeuss et al., 2014; 35 MacLean et al., 2016; MacLean et al., 2019a), beetles (de Jong and Brakefield, 1998; Brakefield and de 36 Jong, 2011; Zvereva et al., 2019), dragonflies (Zeuss et al., 2014) and phasmids (Nosil et al., 2018). Such 37 changes may represent decreased phenotypic diversity and, potentially, genetic diversity (low confidence), 38 but the consequences of climate change on the genetic structure and diversity of populations have not been 39 widely assessed (Pauls et al., 2013). Simplistically, the thermal melanism hypothesis suggests that lighter 40 (higher reflectance) individuals should have increased fitness and therefore be selected for in a warmer 41 climate (Clusella-Trullas et al., 2007). However, several biotic (e.g thermoregulatory requirements, predator 42 avoidance, signalling) and abiotic (e.g. UV, moisture, interannual variability) factors interact to influence 43 changes in colour, making attribution to climate change across species and broad geographic regions difficult 44 (Kingsolver and Buckley, 2015; Stuart-Fox et al., 2017; Clusella-Trullas and Nielsen, 2020). 45 46 Interactions between morphological changes and changes to phenology may facilitate or constrain adaptation 47 to climate change (medium confidence) (Hedrick et al., 2021). For example, advancing phenology in 48 49 migratory species may impose selection on morphological traits (e.g. wing length) to increase migration speed. If advancing spring phenology results in earlier breeding, this may offset the effect of rising 50 temperatures in the breeding range and reduce the effect of increasing temperature on body size (Zimova et 51 al., 2021). A study of 52 species of North American migratory birds, based on more than 70,000 specimens, 52 showed that spring migration phenology has advanced over the past 40 years, concurrent with widespread 53 shifts in morphology (reduced body size and increased wing length), perhaps to compensate for the increased 54 metabolic cost of flight as body size decreases (Weeks et al., 2020). 55 56

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A lack of understanding of physiological constraints and mechanisms remains a barrier to predicting many of 1 the ecological effects of climate change (Bozinovic et al., 2011; Vázquez et al., 2017; González-Tokman et 2 al., 2020). Many behavioural, morphological and physiological responses are highly species and context 3 specific, making generalisations difficult (Bodensteiner et al., 2021). Recent advances in mechanistic 4 understanding (from experiments), in process-based modeling that includes microclimates and 5 developmental processes (Carter and Janzen, 2021) and in sophistication of niche models (Kearney et al., 6 2009) have improved projections, but comprehensive tests of geographic patterns and processes in thermal 7 tolerance and plasticity are still lacking, with studies limited to a few phylogenetically restricted analyses 8 showing mixed results (Gunderson and Stillman, 2015). Improved understanding of the mechanistic basis for 9 observed geographic patterns in thermal tolerance and plasticity is needed to identify species' physiological 10 limits, the potential for adaptation and the presence of evolutionary trade-offs, which will strongly influence 11 population declines, species range shifts, invasive interactions and success of conservation interventions 12 (Cooke et al., 2021; Ryan and Gunderson, 2021). 13

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2.4.2.7 Observed Impacts of Climate Change on Diseases of Wildlife and Associated Impacts on Humans

16 Assessment of changes in diseases of terrestrial and freshwater wild organisms was scarce in WGII AR4, 17 AR5, IPCC SR1.5 and IPCC SRCCL. Most emerging infectious diseases (EIDs) are zoonoses, that is, 18 transmissible between humans and animals, and climate sensitive (Woolhouse et al., 2001; Woolhouse and 19 Gowtage-Sequeria, 2005; McIntyre et al., 2017; Salyer et al., 2017). WGII AR4 found weak to moderate 20 evidence that disease vectors and their diseases had changed their distributions in concert with climate 21 change, but attribution studies were lacking. In WGII AR5 Chapter 11, geographic expansion of a few vector 22 borne diseases (VBDs) to higher latitudes and elevations had been detected and associated with regional 23 climate trends, but non-climatic drivers were not assessed well, leading to only a medium confidence in 24 attribution) (IPCC, 2014). Here we build upon previous assessments by focusing on changes in population 25 dynamics and geographic distributions of diseases of wildlife, and those of humans and domestic animals 26 that are also harbored, amplified, and transmitted by wild animal reservoir hosts and vectors. 27

28 Increased disease incidence is correlated with regional climatic changes and is expected from underlying 29 biology of relationships between temperature, precipitation, and disease ecology (robust evidence, high 30 agreement) (Norwegian Polar Institute, 2009; Tersago et al., 2009; Tabachnick, 2010; Paz, 2015; Dewage et 31 al., 2019; Deksne et al., 2020; Shocket et al., 2020; Couper et al., 2021). Whether increases in diseases in 32 wild and domestic animals correspond to increased disease risk in nearby human populations is complicated 33 by potential buffering effects of the local medical system, healthcare access, socio-economic status, 34 education, behaviours and general health of the human population (see also Chapter 7 and Cross-Chapter 35 Box ILLNESS this Chapter). 36

2.4.2.7.1 Direct effects of climate on reproduction, seasonality, growing season length and transmission of
 pathogens, vectors, and hosts

VBDs require arthropod vector hosts (e.g., insects or ticks), while other infectious diseases (e.g., fungi, 40 bacteria, and helminths) have free-living life stages and/or complex life cycles that require intermediate hosts 41 (e.g., snails), all of which have temperature-driven rates of development and replication/reproduction (robust 42 evidence, high agreement) (Mordecai et al., 2013; Liu-Helmersson et al., 2014; Moran and Alexander, 2014; 43 Bernstein, 2015; Marcogliese, 2016; Ogden and Lindsay, 2016; Mordecai et al., 2017; Short et al., 2017; 44 Caminade et al., 2019: Cavicchioli et al., 2019: Mordecaj et al., 2019: Liu et al., 2020: Rocklöv and Dubrow, 45 2020). Additionally, microbes such as bacteria thermally adapt to temperature changes through multiple 46 mechanisms, indicating that warming will not reduce antibiotic resistance (MacFadden et al., 2018; Pärnänen 47 et al., 2019; Shukla, 2019; McGough et al., 2020; Rodriguez-Verdugo et al., 2020). 48

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There is increasing evidence for a role of extreme events in disease outbreaks (Tjaden et al., 2018; Bryson et al., 2020). Heat waves have been associated with outbreaks of helminth pathogens, especially in subarctic

al., 2020). Heat waves have been associated with outbreaks of helminth pathogens, especially in subarct and Arctic areas. For example, a severe outbreak of microfilaremia, a vector-borne disease spread by

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 mosquitoes and flies, plagued reindeer in northern Europe following extreme high temperatures (Laaksonen)

mosquitoes and flies, plagued reindeer in northern Europe following extreme high temperatures (Laaksonen et al., 2010). More frequent and severe extreme events such as floods, droughts, heat waves, and storms can

- et al., 2010). More frequent and severe extreme events such as floods, droughts, heat waves, and storms can either increase or decrease outbreaks, depending upon the region and disease (*robust evidence, high*
- either increase or decrease outbreaks, depending upon the region and disease (*robust evidence, high agreement*) (Anyamba et al., 2001; Marcheggiani et al., 2010; Brown and Murray, 2013; Paz, 2015; Boyce et
- *agreement*) (Anyamba et al., 2001; Marcheggiani et al., 2010; Brown and Murray, 2013; Paz, 2015; Boyce et al., 2016; Wu et al., 2016b; Wilcox et al., 2019; Nosrat et al., 2021). Heavy precipitation events have been

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shown to increase some infectious diseases with aquatic life cycle components such as mosquito-borne,
helminth, and rodent-borne diseases (*robust evidence, high agreement*) (Anyamba et al., 2001; Zhou et al.,
2005; Wu et al., 2008; Brown and Murray, 2013; Anyamba et al., 2014; Boyce et al., 2016). Conversely,
flooding also increases flow rate and decreases parasite load and diversity in other aquatic wildlife (Hallett
and Bartholomew, 2008; Bjork and Bartholomew, 2009; Marcogliese, 2016; Marcogliese et al., 2016) and
can reduce mosquito abundance by flushing them out of the system (Paaijmans et al., 2007; Paz, 2015).

7

Droughts reduce aquatic habitat of some mosquito species while simultaneously increasing the availability of 8 stagnant standing pools of water that are ideal breeding habitats for other species, such as dengue vector 9 Aedes mosquitos (medium evidence, medium agreement) (Chareonviriyaphap et al., 2003; Chretien et al., 102007; Padmanabha et al., 2010; Trewin et al., 2013; Paz, 2015). Extreme drought has been associated with 11 an increase in bluetongue virus haemorrhagic disease in wildlife in eastern North America, though 12 mechanisms were not identified (Christensen et al., 2020). Heatwaves in some regions, especially coastal 13 regions, increased parasitism and decreased host richness and abundance leading to population crashes 14 (Larsen and Mouritsen, 2014; Mouritsen et al., 2018). Changes in temperature and precipitation, especially 15 extreme events, can alter community structure (Larsen et al., 2011) by increasing or decreasing parasites and 16 their host organisms and even altering host behavior in ways advantageous to parasites (Macnab and Barber, 17 2012). 18

18

Climate change not only affects the occurrence of pathogens and their hosts in geographic space but also the
 temporal patterns of disease transmission. Warmer winters allow greater overwinter survival of arthropod
 vectors which, coupled with lengthened tranmission seasons, drive increases in vector population sizes,
 pathogen prevalence, and hence proportion of vectors infected *(robust evidence, high agreement)* (Laaksonen et al., 2009; Molnár et al., 2013; Waits et al., 2018). For example, a parasitic nematode lung
 worm (*Umingmakstrongylus pallikuukensis*) has shortened its larval development time in half (from two

years to one year), which has increased infection rates in North American muskoxen (Norwegian Polar Institute, 2009).

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Case study 1: Climate change impacts on pathogenic helminths in Europe

Parasitic helminth worms can reduce growth and yield, or kill livestock, and infect humans and wildlife,
leading to health, agricultural and economic losses (Fairweather, 2011; Charlier et al., 2016; Charlier, 2020).
Attribution of increased helminth disease incidence and risk to climate change is stronger than for most
human diseases because of long-term records and careful analysis of other anthropogenic drivers (e.g. land
use change, agricultural/livestock intensification, and antihelminthic intervention and resistance) (van Dijk et
al., 2008; van Dijk et al., 2010; Fox et al., 2011b; Martínez-Valladares et al., 2013; Charlier et al., 2016;
Innocent et al., 2017; Mehmood et al., 2017).

- 38 In Europe, evidence from laboratory studies, long term surveillance, statistical analyses, and modelling 39 shows that multiple helminth pathogens and their host snails have extended their transmission windows and 40 have increased survival, fecundity, growth and abundances (robust evidence, high agreement). Furthermore, 41 they have expanded or shifted their ranges poleward due to increases in temperature, precipitation and 42 humidity (robust evidence, high agreement) (Lee et al., 1995; Pritchard et al., 2005; Poulin, 2006; van Dijk 43 et al., 2008; van Dijk et al., 2010; Fairweather, 2011; Fox et al., 2011b; Martínez-Valladares et al., 2013; 44 Bosco et al., 2015: Caminade et al., 2015: Caminade et al., 2019). These documented changes in climate. 45 hosts and pathogens have been linked to higher disease incidence and more frequent outbreaks in livestock 46 across Europe (high confidence) (Bosco et al., 2015). 47
- 48

49 *Case study 2: Chytrid fungus and climate change*

Infection by the chytrid fungus, *Batrachochytrium dendrobatidis* (Bd), can cause chytridiomycosis in
 amphibians. Bd is widely distributed globally and has caused catastrophic disease in amphibians, associated
 with declines of 501 species and extinctions of a further 90 species, primarily in tropical regions of the

Americas and Australia (Scheele et al., 2019; Fisher and Garner, 2020). Bd successfully travelled with high-

elevation Andean frog species as they expanded their elevational ranges upward, driven by regional

so warming, to > 5200 m (Seimon et al., 2017).

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New findings since AR5 from controlled laboratory experiments (manipulating temperature, humidity and
water availability), intensive analyses of observed patterns of infection and disease in nature, and modeling
studies have led to an emerging consensus that interactions between chytrids and amphibians are climatesensitive, and that the interaction of climate change and Bd has driven many of the observed global
amphibian declines and species' extinctions *(robust evidence, high agreement)* (Rohr and Raffel, 2010;
Puschendorf et al., 2011; Rowley and Alford, 2013; Raffel et al., 2015; Sauer et al., 2018; Cohen et al.,
2019a; Sauer et al., 2020; Turner et al., 2021).

8

9 The "thermal mismatch hypothesis" posits that vulnerability to disease should be higher at warm 10 temperatures in cool-adapted species and higher at cool temperatures in warm-adapted species and is 11 generally supported. However, the most recent studies reveal more complex mechanisms underlying 12 amphibian-disease-climate change dynamics, including variation in thermal preferences among individuals 13 in a single amphibian population *(robust evidence, high agreement)* (Zumbado-Ulate et al., 2014; Sauer et 14 al., 2018; Cohen et al., 2019b; Neely et al., 2020; Sauer et al., 2020).

15 Bd is not universally harmful—it has been recorded as endemic in frog populations that did not suffer 16 disease, where it may be commensal rather than parasitic (Puschendorf et al., 2006; Puschendorf et al., 2011; 17 Rowley and Alford, 2013). Projections of future impacts are difficult, as the virulence of Bd is variable 18across Bd populations and dependent upon the evolutionary and ecological histories, and evolutionary 19 potentials, of both the local amphibian populations and the endemic or invading Bd (robust evidence, high 20 agreement) (Retallick et al., 2004; Daskin et al., 2011; Puschendorf et al., 2011; Phillips and Puschendorf, 21 2013; Rowley and Alford, 2013; Zumbado-Ulate et al., 2014; Sapsford et al., 2015; Voyles et al., 2018; 22 Bradley et al., 2019; Fisher and Garner, 2020; McMillan et al., 2020). Further, specific local habitats might 23 serve as regional climate refugia from chytrid infection (e.g. hot and dry) (medium evidence, high 24 agreement) (Zumbado-Ulate et al., 2014; Cohen et al., 2019b; Neely et al., 2020; Turner et al., 2021). 25

26 27 28

2.4.2.7.2 Effects on geographic distribution and connectivity patterns of pathogens

As species' geographic ranges and migration patterns are modified by climate change (Section 2.4.2.1, Table 29 2.2), pathogens accompany them. Diverse vectors and associated parasites, pests, and pathogens of plants 30 and animals are being recorded at higher latitudes and elevations in conjunction with regional temperature 31 increases and precipitation changes (robust evidence, high agreement), although analysis of realized disease 32 incidence often lacks inclusion of non-climatic vs climate drivers, compromising attribution (Ollerenshaw 33 and Rowlands, 1959; Purse et al., 2005; Laaksonen et al., 2010; van Dijk et al., 2010; Alonso et al., 2011; 34 Genchi et al., 2011; Pinault and Hunter, 2011; Jaenson et al., 2012; Loiseau et al., 2012; Kweka et al., 2013; 35 Medlock et al., 2013; Dhimal et al., 2014a; Dhimal et al., 2014b seasonal; Siraj et al., 2014; Khatchikian et 36 al., 2015; Hotez, 2016a; Hotez, 2016b; Bett et al., 2017; Mallory and Boyce, 2017; Strutz, 2017; Booth, 37 2018; Dumic and Severnini, 2018; Carignan et al., 2019; Gorris et al., 2019; Le et al., 2019; Stensgaard et 38 al., 2019b snails and; Brugueras et al., 2020; Gilbert, 2021). 39

At least six major VBDs affected by climate drivers have recently emerged in Nepal and are now considered 41 endemic, with climate change implicated as a primary driver as LULCC has been assessed to have a minimal 42 influence on these diseases (high confidence) (Table SM2.1). There is increasing evidence that climate 43 warming has extended the elevational distribution of Anopheles, Culex and Aedes mosquito vectors above 44 2,000 m in Nepal (limited evidence, high agreement) (Dahal, 2008; Dhimal et al., 2014a; Dhimal et al., 45 2014b; Dhimal et al., 2015) with similar trends being recorded in neighboring Himalayan regions (medium 46 evidence, high agreement) (Phuyal et al., 2020; Dhimal et al., 2021). Host animals in novel areas may be 47 immunologically naive, and therefore more vulnerable to severe illness (Bradley et al., 2005; Hall et al., 48 49 2016).

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51 *Case study 3: Arctic and subarctic disease expansion and intensification*

High Arctic regions have warmed by more than double the global average, >2°C in most areas (see Sections
2.3.1.1.2, Figure 2.11, and Atlas 11.2.1.2, in WGI). Experimental, field ecology studies and computational
models in Arctic and subarctic regions indicate that milder winters have reduced mortality of vectors and
reservoir hosts and increased their habitat as forested taiga expands into previously treeless tundra (Table
SM2.1; Parkinson et al., 2014). Warmer temperatures and longer seasonal windows have allowed faster

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reproduction/replication, accelerated development, and increased the number of generations per year of 1 pathogens, vectors and some host animals, that in turn increase the populations of disease organisms and 2 disease transmission (Sections 2.4.2.4, 2.4.4.3.3). Ticks, mosquitos, culicoides biting midges, deer flies, 3 horseflies, and simuliid black flies that transmit a variety of pathogens are being documented in high-latitude 4 regions at higher numbers or where they have been historically absent (robust evidence, high agreement) 5 (Waits et al., 2018; Caminade et al., 2019; Gilbert, 2021). In concert with these poleward shifts of hosts and 6 vectors, pathogens, particularly tick-borne pathogens and helminth infections, have increased dramatically in 7 incidence and severity from once rare occurrences and have appeared in new regions (robust evidence, high 8 agreement, very high confidence) (Caminade et al., 2019; Gilbert, 2021). 9 10

- Zoonoses and VBDs that have been historically rare or never documented in Arctic and subarctic regions of 11 Europe, Asia, and North America, such as anthrax, cryptosporidiosis, elaphostrongylosis, filariasis (Huber et 12 al., 2020), tick-borne encephalitis, and tularemia (Evander and Ahlm, 2009; Parkinson et al., 2014; Pauchard 13 et al., 2016), are spreading poleward and increasing in incidence (robust evidence, high agreement, very high 14 confidence) (Table SM2.1; Omazic et al., 2019). Recent anthrax outbreaks and mass mortality events among 15 humans and reindeer, respectively, have been linked to abnormally hot summer temperatures that caused 16 permafrost to melt and exposed diseased animal carcasses, releasing thawed, highly infectious, Bacillus 17 anthracis spores (medium evidence, medium agreement) (Ezhova et al., 2019; Hueffer et al., 2020; Ezhova et 18al., 2021). Multiple contributing factors conspired over different time scales to compound a 2016 anthrax 19 20 outbreak occurring on the Yamal peninsula: (i) rapid permafrost thawing for 5 years preceding the outbreak; (ii) thick snow cover the year before the outbreak insulated the warmed permafrost and kept it from re-21 freezing; and (iii) anthrax vaccination rates had decreased or ceased in the region (Ezhova et al., 2019; 22 Ezhova et al., 2021). These precursors converged with an unusually dry and hot summer that: (i) melted 23 permafrost, creating an anthrax exposure hazard; (ii) increased the vector insect population; and (iii) 24
- weakened the immune systems of reindeer thus increasing their susceptibility (Waits et al., 2018; Hueffer et al., 2020).
- 27

Warmer temperatures have increased blood-feeding insect harassment of reindeer with compounding 28 consequences: (1) increased insect bite rates lead to higher parasite loads, (2) time spent by raindeer in trying 29 to escape biting flies reduces foraging while simulataneously increasing energy expenditure, (3) the 30 combination of (1) and (2) lead to poor body condition, that subsequently leads to (4) reduced winter 31 survival and fecundity (Mallory and Boyce, 2017). As temperatures warm and connectivity increases 32 between the Arctic and the rest of the world, tourism, resource extraction, and increased commercial 33 transport will create additional risks of biological invasion by infectious agents and their hosts (Pauchard et 34 al., 2016). These increases in introduction risk compounded with climate change have already begun to harm 35 indigenous peoples dependent on hunting and herding livestock (horses and reindeer) that are suffering 36 increased pathogen infection (Deksne et al., 2020; Stammler and Ivanova, 2020). 37 38

39 2.4.2.7.3 Biodiversity-disease links

Anthropogenic impacts, such as disturbances caused by climate change, can reduce biodiversity through multiple mechanisms and increase disease risk to humans *(limited evidence, low agreement)* but more research is needed to understand the underlying mechanisms (Civitello et al., 2015; Young et al., 2017b; Halliday et al., 2020; Rohr et al., 2020; Glidden et al., 2021). Known wildlife hosts of human-shared pathogens and parasites overall comprise a greater proportion of local species richness (18–72% higher) and abundance (21–144% higher) in sites under substantial human use (agricultural and urban lands) compared with nearby undisturbed habitats (Gibb et al., 2020).

- 47
- Exploitation of wildlife and degradation of natural habitats have increased opportunities for 'spill over' of pathogens from wildlife to human populations and increased emergence of zoonotic disease epidemics and pandemics *(robust evidence, high agreement)*; animal and human migrations driven by climate change have added to this increased risk *(medium evidence, medium agreement)* (see Section 2.4.2.1, Chapter 8, Cross-Chapter Box MOVING PLATE in Chapter 5; Patz et al., 2004; Cleaveland et al., 2007; Karesh et al., 2012; Altizer et al., 2013; Allen et al., 2017; Plowright et al., 2017; Olivero et al., 2017; Faust et al., 2018; Carlson
- et al., 2020; Gibb et al., 2020; Hockings et al., 2020; IPBES, 2020; Volpato et al., 2020; Glidden et al.,
- ⁵⁵ 2021). Agricultural losses and subsequent food scarcity, that is increasing due to climate change, can also
- lead to an increase in the use of bushmeat, and, hence increase risk of diseases jumping from wild animals to
- ⁵⁷ humans (*medium evidence, high agreement*) (Brashares et al., 2004; Leroy et al., 2004; Wolfe et al., 2004;

Rosen and Smith, 2010; Kurpiers et al., 2016).

2.4.2.7.4 Implications for humans of changes in diseases in wild animals

3 Changes in temperature, precipitation, humidity, and extreme events have been associated with more 4 frequent disease outbreaks, increases in disease incidence and severity, and novel disease and vector 5 emergence into new areas for wild animals, with a mechanistic understanding of the roles of these drivers 6 from experimental studies providing high confidence for the role of climate change. However, attribution of 7 how this has impacted human infectious diseases remains difficult, and definitive attribution studies are 8 lacking. The specific role of recent climate change is difficult to examine in isolation for most regions where 9 human disease incidence has also been affected by land use change (particularly agricultural and urban 10 expansion), changes in public health access and measures, socio-economic changes, increased global 11 movements of people, and changes in vector and rodent control programs, supporting medium confidence in 12 the role of climate change driving observed changes in human diseases globally. Exceptions are in areas 13 noted above (Arctic, subarctic, and high elevation regions), in which climate change fingerprints are strong 14 and/or concurrent changes in non-climatic drivers are less pronounced than in other regions (high confidence 15 for climate change attribution) (see Table SM2.1, Sections 5.5.1.3, 7.2.2.1, Cross-Chapter Box ILLNESS 16 this Chapter; Harvell et al., 2002; Norwegian Polar Institute, 2009; Tersago et al., 2009; Tabachnick, 2010; 17 Altizer et al., 2013; Garrett et al., 2013; Paz, 2015; Wu et al., 2016b; Caminade et al., 2019; Dewage et al., 18 2019; Coates and Norton, 2020; Deksne et al., 2020; Shocket et al., 2020; Couper et al., 2021; Gilbert, 19 2021). 20

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[START FAQ2.2 HERE] 23

FAO2.2: How does climate change increase the risk of diseases? 25

26 Climate change is contributing to the spread of diseases in both wildlife and humans. Increased contact 27

between wildlife and human populations increases disease risk and climate change is altering where 28

pathogens that cause diseases and the animals that carry them live. Disease risk can often be reduced by 29

improving health care and sanitation systems, training the medical community to recognize and treat 30

potential new diseases in their region, limiting human encroachment into natural areas, limiting wildlife 31

trade, and promoting sustainable and equitable socioeconomic development. 32

33 Diseases spread between humans and animals are called zoonoses. Zoonoses comprise nearly two-thirds of 34 known human infectious diseases and the majority of newly emerging infectious diseases (EIDs). COVID-19 35 is the most recent zoonosis and has killed millions of people globally while devastating economies. The risk 36 posed by EIDs has increased because of: (1) movement of wild animals and their parasites into new areas via 37 climate change, global trade, and travel; (2) human intrusion into and conversion of natural areas for 38 agriculture, livestock, industrial/raw materials extraction, and housing; (3) increased wildlife trade and 39 consumption; (4) increased human mobility resulting from global trade, war/conflicts, and migration made 40 faster and farther by fossil fuel powered travel; and (5) widespread antimicrobial use, which can promote 41 antibiotic resistant infections (Figure FAQ2.3.1). 42 43



Figure FAQ2.2.1: How diseases move from the wild into human populations. Climate change may increase diseases in nature, but whether or not this leads to an increase in disease risk for humans depends upon a range of societal, infrastructure and medical buffers that form a shield protecting humans.

6 Climate change further increases risk by altering pathogen and host animal (1) geographic ranges and 7 habitats; (2) survival, growth, and development; (3) reproduction and replication; (4) transmission and 8 exposure (5) behavior; and (6) access to immunologically naïve animals and people who lack infection 9 resistance. This can lead to novel disease emergence in new places, more frequent and larger outbreaks, and 10 longer or shifted seasons of transmission. Climate change is making it possible for many EIDs to colonize 11 historically colder areas that are becoming warmer and wetter in temperate and polar regions and in 12 mountains. Vector-borne diseases (VBDs) are diseases spread by vectors such as mosquitoes, sand flies, 13 kissing bugs, and ticks. For example, ticks that carry the virus that causes tick-borne encephalitis have 14 moved into northern subarctic regions of Asia and Europe. Viruses like dengue, chikungunya, and Japanese 15 encephalitis are emerging in Nepal in hilly and mountainous areas. Novel outbreaks of Vibrio bacteria 16 seafood poisoning are being traved to the Baltic States and Alaska where they were never documented 17 before. Many scientific studies show that infectious disease transmission and the number of individuals 18 infected depends on rainfall and temperature; climate change often makes these conditions more favourable 19 for disease transmission. 20

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Climate change can also have complicated, compounding, and contradictory effects on pathogens and vectors. Increased rainfall creates more habitat for mosquitoes that transmit diseases like malaria but too much rain washes away the habitat. Decreased rainfall also increases disease risk when people without reliable water access use containers to store water in that mosquitos, such as the vectors of dengue fever, *Aedes aegypti* and *Ae. albopictus*, use for egg laying. Hotter temperatures also increase mosquito bite rate,

1	parasite development, and viral replication: Certain species of sharts are methodiate nosis for many
2	helminth worm parasites that make humans, livestock, and wild animals sick. When it gets hot, the snails can
3	produce two to three times as many infective larvae but if it becomes too hot many pathogens and their
4	vectors cannot survive or reproduce.
5	
6	Humans also contract zoonoses directly through their skin, mucus membranes, and lungs when eating or
7	butchering animals, or coming into contact with pathogen shed in the air, urine, or faeces that contaminates
8	water, food, clothing, and other surfaces. Any activity that increases contact with wildlife, especially in high
9	biodiversity regions like the tropics and subtropics, increases disease risk. Climate change-related disease
10	emergence events are often rare but may become more frequent. Fortunately, there are ways to reduce risks
11	and protect our health, as described below.
12	
13	Habitat and biodiversity protection: Encroachment of humans into natural areas, due to expansion of
14	agriculture and livestock, timber harvests, resource extraction, and urban development has increased human
15	contact with wild animals, and creates more opportunities for disease spillover (transmission from an animal
16	to a new species, including humans). By conserving, protecting, and restoring wild habitats, we can build
17	healthier ecosystems that provide other services, such as clean air, clean and abundant water, recreation,
18	spiritual value, and well-being, as well as reduced disease spillover. If humans must go into wild areas or
19	hunt, they should take appropriate precautions such as wearing protective clothing, using insect repellant,
20	performing body checks for vectors like ticks, and washing hands and clothing well.
21	
22	Food resilience: Investing in sustainable agroecological farming will alleviate the pressure to hunt wild
23	animals and reduce the conversion of more land to agriculture/livestock use. Stopping illegal animal trading
24	and poaching and decreasing reliance on wild meats and products made from animal parts will reduce direct contact with potentially infected animals. This has the added benefit of increasing food security, nutrition,
25 26	improving soil, reducing erosion, preserving biodiversity, and mitigating climate change.
26 27	improving son, reducing crosion, preserving biodiversity, and intigating chinate change.
27	Disease prevention and response: The level of protection against infection is linked directly to the level of
28 29	development and wealth of a country. Improved education, high-quality medical and veterinary systems,
30	high food security, proper sanitation of water and waste, high housing quality, and disease surveillance and
31	alarm systems dramatically reduce disease risk and improve health. Utilizing a One Biosecurity or One
32	Health framework further improves resilience. Sharing knowledge within communities, municipalities,
33	regional, and between national health authorities globally is important to assessing, preventing and
34	responding to outbreaks and pandemics more efficiently and economically.
35	
36	Humans are facing many direct or indirect challenges because of climate change. Increasing EIDs is one of
37	our greatest challenges, due to our ever-growing interactions with wildlife and the climatic changes creating
38	new disease transmission patterns. COVID-19 is a current crisis, and follows other recent EIDs: SARS,
39	HIV/AIDS, H1N1 influenza, Ebola, Zika, and West Nile fever. EIDs have accelerated in recent decades,
40	making it clear that new societal and environmental approaches to wildlife interactions, climate change, and
41	health are urgently needed to protect our current and future well-being as a species.
42	
43	[END FAQ2.2 HERE]
44	
45	
46	2.4.2.8 Observed Evolutionary Responses to Climate Change
47	
48	Prior sections document species' tendencies to retain their climate envelopes by some combination of range
49 50	shift and phenological change. However, this tracking of climate change can be incomplete, causing species
50	or populations to experience hotter conditions than those to which they are adapted and thereby incur
51	'climate debts' (Devictor et al., 2012). The importance of population-level debt is illustrated by a study in which estimated debt values were correlated with population dynamic trands in a North American migratory.
52	which estimated debt values were correlated with population dynamic trends in a North American migratory

songbird, the Yellow Warbler, *Setophaga petechia*. Populations that were genetic outliers for their local
 climate space had larger population declines (greater debt) than populations with genotypes closer to the

climate space had larger population declines (greater debt) than populations with genotypes closer to th average values for that particular climate space. Debt values were estimated from genomic analyses

independent of the population trends, and were distributed across the species' range in a mosaic, not simply

1

parasite development, and viral replication! Certain species of snails are intermediate hosts for many

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1 2	concentrated at range margins, trends (Bay et al., 2018).	rendering the results robust to being	g confounded by broad-scale geographical
3 4 5 6 7 8	thermal tolerance and ability to adaptation to climate, populati	o perform in high ambient temperatu ons currently living at relatively coo	ncelled by genetically-based increases in ures. In species already showing local of sites should be able to evolve to adopt ence of climate changes (Singer, 2017;
9 10 11 12 13 14	warm range limits (Franks and capacity to genetically track cl (Boutin and Lane, 2014). Inde	Hoffmann, 2012). Organisms with imate change than species with long ed, observed evolutionary impacts h	to climate change in populations not at short generation times should have higher generation times, such as mammals ave been mainly documented in insects, 6) where evolutionary changes have been
15 16 17			cialisation (Bridle et al., 2014; Lancaster,
18 19 20 21 22 23	related to climate adaptation. F America have evolved to wait lengthening summers enables (Bradshaw and Holzapfel, 200 continents, 22 experienced clim	For example, pitcher-plant mosquitos for shorter day lengths before initiat them to delay overwintering until lat 1). Among 26 populations of <i>Droso</i> nate warming across two or more de	ter and add an extra generation each year <i>phila subobscura</i> studied on three ecades, and 21 of those 22 showed
24 25 26	(Balanya et al., 2006).	mosome inversions characteristic of	
27 28 29 30	would require evolving to surv not supported by experimental Whether or not they can do so	vive and reproduce outside their spector or observational evidence <i>(medium)</i> depends on the level of 'niche conservational evidence of 'niche	ability to track climate change in situ sies' historical climate envelope, which is <i>evidence, high agreement)</i> (Singer, 2017). ervatism' operating at the species level
31 32 33 34	outside of its traditional climat and Gonzalez, 2009; Bell et al	e envelope, extinction is expected in , 2019). To investigate the evolution	ined by climate finds itself completely a the absence of 'evolutionary rescue' (Bell mary potential enabling a species to survive experiments have been carried out on
35 36 37 38	Xue et al., 2019). Tests of ther transgenerational effects occur	mal performance have been complice (Sgro et al., 2016). However, the re	their evolvabilities (Castaneda et al., 2019; cated as both long-term acclimation and esults to date have been consistent: despite tantial constraints exist to the evolution of
39 40 41 42	agreement) (Hoffmann and Sg experimentally increased, the a	ro, 2011; MacLean et al., 2019b). Fo amount of genetic variance in fitness	rm range limits <i>(medium evidence, high</i> or example, as temperature was s of <i>Drosophila melanogaster</i> decreased: 2015). The hypothesis that heat stress
43 44 45 46	tolerance is evolutionarily con drawn from tropical and tempe	strained is further supported by expe erate climes were subjected to extrem it not in heat tolerances nor in tempe	eriments in which 22 Drosophila species nes of heat and cold. They differed as eratures at which optimal performances
47 48 49 50 51	extent of plasticity can vary an phenotypic values outside the evolve in response to climate of	nong populations experiencing differ prior range for the species, but plasti	to environmental change. The form and rent climates (Kelly, 2019) and generate icity itself has not yet been observed to tic changes in nature (e.g. affecting heat

tolerance) have not yet been shown to alter the boundaries of existing genetic variation for any species.
 Evolutionary rescue of entire species has not yet been observed in nature, nor is it expected based upon

experimental and theoretical studies *(medium evidence, high agreement)*.

55

56 Hybridisation between closely related species has increased in recent decades as one species shifts its range 57 boundaries and positions itself more closely to the other—hybrids between polar bears and brown bears have

1 2	been documented in northern Canada (Kelly et al., 2010). In North American rivers, hybridisation between invasive rainbow trout and native cutthroat trout has increased in frequency as the rainbow trout expanded
3	into warming waters (Muhlfeld et al., 2014). Whether climate-changed induced hybridisations can generate
4	novel climate adaptations remains to be seen.
4 5	nover enimate adaptations remains to be seen.
6	In summary, with present knowledge, evolution is not expected to be sufficient to prevent whole species'
7	extinctions if a species' climate space disappears (high confidence).
8	entitierions il a species entitate space alsappears (mgn conjuncto).
9	2.4.3 Observed Changes in Key Biomes, Ecosystems and their Services
10	2. 1.5 Observed Changes in hey biomes, Debsystems and men Services
11	2.4.3.1 Detection and Attribution for Observed Biome Shifts
12	2.1.9.1 Detection and manifold observed biome shighs
13	Attribution for biome (major vegetation form of an ecosystem) shifts is complex because of their extensive,
14	sometimes continental, spatial scale (Whittaker, 1975; Olson et al., 2001; Woodward et al., 2004); and
15	therefore, non-climatic factors strongly influence biome spatial distributions (Ellis and Ramankutty, 2008).
16	
17	The most robust attribution studies use data from individual locations with minimal confounding factors,
18	particularly recent land use change, and scale up by analysing multiple locations across a long zone between
19	biomes. As with individual species, multiple lines of evidence increase confidence (Hegerl et al., 2010;
20	Parmesan et al., 2013). Multivariate statistical analyses aid attribution studies by allowing the assessment of
21	relative weights among multiple factors, including variables related to climate change (Gonzalez et al.,
22	2012). However, drivers often have strong, significant interactions with one another, complicating
23	quantitative assessment of the strength of individual drivers (Parmesan et al., 2013). In these cases,
24	manipulative experiments are critical in assessing attribution to climate change drivers.
25	
26	Certain biomes exhibit a relatively stronger relationship to climate; for example, Arctic tundra generally has
27	a distinct ecotone with boreal conifer forest (Whittaker, 1975). In these areas, attribution of biome shifts to
28	climate change are relatively straightforward, if human land use change is minimal. However, other biomes,
29	such as many grassland systems, are not at equilibrium with climate (Bond et al., 2005). In these systems
30	their evolutionary history (Keeley et al., 2011; Strömberg, 2011; Charles-Dominique et al., 2016),
31	distribution, structure and function have been shaped by climate and natural disturbances, such as fire and
32	herbivory (Staver et al., 2011; Lehmann et al., 2014; Pausas, 2015; Bakker et al., 2016; Malhi et al., 2016).
33	Disturbance variability is an inherent characteristic of grassland systems and suitable "control" conditions
34	are seldom available in nature. Furthermore, due to the integral role of disturbance, these biomes have been
35	widely affected by long-term and widespread shifts in grazing regimes, large-scale losses of mega-
36	herbivores and fire suppression policies (Archibald et al., 2013; Malhi et al., 2016; Hempson et al., 2017). It
37	is necessary to conduct climate change attribution on a case-by-case basis for grasslands; such assessments
38	are complex as direct climate change impacts from either inherent variation within disturbance regimes or
39	directional changes in background disturbances are difficult to separate (detailed in Sections 2.4.3.2.1;
40	2.4.3.2.2; 2.4.3.5). Confidence in assessments is increased when observed trends are supported by
41	mechanistic understanding of responses identified by physiological studies, manipulative field experiments,
42	greenhouse studies and lab experiments (Table SM2.1).
43	
44	2.4.3.2 Global Patterns of Observed Biome Shifts Driven by Climate Change

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45 Observed biome shifts predominantly driven by climate change 2.4.3.2.1 46

Th IPCC Fifth Assessment Report and a meta-analysis found that vegetation at the biome level shifted 47 poleward latitudinally and upward altitudinally due to anthropogenic climate change at 19 sites in boreal, 48 temperate, and tropical ecosystems from 1700 to 2007 (Gonzalez et al., 2010a; Settele et al., 2014). In these 49 areas, temperature increased 0.4° to 1.6°C above the pre-industrial period (Gonzalez et al., 2010a; Settele et 50 al., 2014). Field research since the IPCC Fifth Assessment Report detected additional poleward and upslope 51 biome shifts over periods of 24 to 210 years at numerous sites (described below) but were not directly 52 attributed to anthropogenic climate change as the studies were not designed nor conducted properly for 53 attribution. 54

55

Many of the recently detected shifts were nevertheless consistent with climate change temperature increases 56 and observed in areas lacking agriculture, livestock grazing, timber harvesting, or other anthropogenic land 57

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uses. For example, in the Andes Mountains in Ecuador, a biome shift was detected by comparing a survey by
 Alexander von Humboldt in 1802 to a re-survey in 2012, making this the longest time span in the world for
 this type of data (Morueta-Holme et al., 2015) and 2017 (Moret et al., 2019). During 210 years, temperature
 increased 1.7°C (Morueta-Holme et al., 2015) and the upper edge of alpine grassland shifted upslope 100–
 450 m (Moret et al., 2019).

6

Other biome shifts consistent with climate change and not substantially affected by local land use include 7 northward shifts of deciduous forest into boreal conifer forest in Canada (5 km between 1970-2012, (Sittaro 8 et al., 2017) and 20 km between 1970-2014, (Boisvert-Marsh et al., 2019)) and northward shifts of 9 temperate conifer into boreal conifer forest in Canada (21 km between 1970-2015, (Boisvert-Marsh and de 10 Blois, 2021)). Research detected upslope shifts of boreal and sub-alpine conifer forest into alpine grassland 11at 143 sites on four continents (41 m, 1901–2018, (Lu et al., 2021)) and individual sites in Canada (54 m, 12 1900-2010, (Davis et al., 2020)), China (300 m, 1910-2000 (Liang et al., 2016); 33 m, 1985-2014, (Du et 13 al., 2018)), Nepal (50 m, 1860–2000, (Sigdel et al., 2018)), Russia (150 m, 1954–2006, (Gatti et al., 2019)) 14 and the United States (19 m, 1950-2016, (Smithers et al., 2018); 38 m, 1953-2015, (Terskaia et al., 2020)). 15 Other upslope cases include shifts of temperate conifer forest in Canada (Jackson et al., 2016) and the United 16 States (Lubetkin et al., 2017), temperate deciduous forest in Switzerland (Rigling et al., 2013) and temperate 17 shrubland in the United States (Donato et al., 2016). 18

- In summary, anthropogenic climate change has caused latitudinal and elevational biome shifts in at least 19 sites in boreal, temperate, and tropical ecosystems between 1700 and 2007, where temperature increased 0.4° to 1.6°C above the pre-industrial period (*robust evidence, high agreement*). Additional cases of 5 to 20 km northward and 20 to 300 m upslope biome shifts between 1860 and 2016, under approximately 0.9°C mean global temperature increase above the pre-industrial period, are consistent with climate change (*medium evidence, high agreement*).
- 26 27

19

28 2.4.3.2.2 Observed biome shifts from combined land use change and climate change

Research has detected biome shifts in areas where agriculture, fire use or suppression, livestock grazing, timber and fuelwood harvesting, or other local land use actions substantially altered vegetation, in addition to changes in climatic factors and CO₂ fertilisation. These studies were not designed or conducted in a manner to make climate change attribution possible, although vegetation changes are consistent with climate change: for example, a global review of observed changes in treelines found that 2/3 of treelines globally have shifted upslope in elevation over the past 50 years or more (Hansson, 2021, a review of).

35

Upslope and poleward forest shifts have occurred where timber harvesting or livestock grazing was 36 abandoned, allowing regeneration of trees at sites in Canada (Brice et al., 2019; Wang et al., 2020b), France 37 (Feuillet et al., 2020), Italy (Vitali et al., 2017), Spain (Ameztegui et al., 2016), the United States (Wang et 38 al., 2020b) and mountain areas across Europe (Cudlin et al., 2017). Intentional use of fire drove an upslope 39 forest shift in Peru (Bush et al., 2015) while mainly human-ignited fires drove conversion of shrubland to 40 grassland in a drought-affected area of the United States (Syphard et al., 2019b). In eastern Canada, timber 41 harvesting and wildfire drove conversion of mixed conifer-broadleaf forests to broadleaf-dominated forests 42 (Brice et al., 2020; Wang et al., 2020b). 43

44

Shrub encroachment onto savanna has occurred at numerous sites, particularly across the Southern 45 Hemisphere, mainly between 1992 and 2010 (Criado et al., 2020). Globally, overgrazing initiates shrub 46 encroachment by reducing grasses more than woody plants, while fire exclusion maintains the shrub cover 47 (D'Odorico et al., 2012; Caracciolo et al., 2016; Bestelmeyer et al., 2018). The magnitude of woody cover 48 change in savannas is not correlated to mean annual temperature change (Criado et al., 2020), however, 49 higher atmospheric CO₂ increases shrub growth in savannas (Nackley et al., 2018; Manea and Leishman, 50 2019). A global remote sensing analysis of biome changes from all causes, including agricultural and grazing 51 expansion and deforestation, estimated that 14% of pixels changed between 1981 and 2012, although this 52 approach can overestimate global changes since it uses a new biome classification system, which doubles the 53 conventional biome classifications (Higgins et al., 2016). In addition to climate change, land use change 54 causes vegetation changes at the biome level (robust evidence, high agreement). 55 56

1 2

2.4.3.3 Observed Changes in Deserts and Arid Shrublands

Divergent responses to anthropogenic climate change are occurring within and across arid regions,
depending on time period, location, detection methodology and vegetation type (see Cross-Chapter Paper 3).
Emerging shifts in ecosystem structure, functioning and biodiversity are supported by evidence from
modelled impacts of projected climate and CO₂ levels. While observed responsiveness of arid vegetation
productivity to rising atmospheric CO₂ (Fensholt et al., 2012b) may offset risks from reduced water
availability (Fang et al., 2017), climate- and CO₂ driven changes are key risks in arid regions, interacting
with habitat degradation, wildfire, and invasive species (Hurlbert et al., 2019).

10

Widespread vegetation greening, as projected in AR4, is occurring in arid shrublands (Zhang et al., 2019a; Maestre et al., 2021) as a result of increases in leaf area, woody cover and herbaceous production at desert-

13 grassland interfaces (Gonsamo et al., 2021). Plant productivity in arid regions has increased (Fensholt et al.,

14 2012b) because of improved water use efficiency associated with elevated CO₂ (Norby and Zak, 2011;

Donohue et al., 2013; Burrell et al., 2020; Gonsamo et al., 2021) *(medium evidence, high agreement)*, altered rainfall seasonality and amount (Rohde et al., 2019; Zhang et al., 2019a) *(robust evidence, high agreement)*,

rainfall seasonality and amount (Rohde et al., 2019; Zhang et al., 2019a) (robust evidence, high agreement),
 increases in temperature (Ratajczak et al., 2014; Wilcox et al., 2018)(robust evidence, high agreement) and

heavy grazing (*robust evidence, high agreement*) with relative importance differing among locations

(Donohue et al., 2013; Caracciolo et al., 2016; Archer et al., 2017; Hoffmann et al., 2019b; Rohde et al.,

20 2019). Woody plant encroachment into arid shrublands is occurring in North America (Caracciolo et al.,

21 2016; Archer et al., 2017), southern Africa (du Toit and O'Connor, 2014; Ward et al., 2014; Masubelele et

al., 2015a; Hoffman et al., 2019; Rohde et al., 2019) (*high confidence*) and Central Asia (Li et al., 2015) (*low*

confidence). In North America, sagebrush steppe changes have been attributed to increases in temperature

and earlier snowpack melt (Wuebbles et al., 2017; Mote et al., 2018; Snyder et al., 2019).

25 26

Non-native grasses are invading the sagebrush steppes (cold deserts) in North America (Chambers et al.,

27 2014) attributed to warming (Bradley et al., 2016; Hufft and Zelikova, 2016). In the eastern semi-desert

28 (Karoo) of South Africa, annual rainfall increases and a rainfall seasonality shift (du Toit and O'Connor,

29 2014) are increasing grassiness as arid grasslands expand into semi-desert shrublands (du Toit et al., 2015;

Masubelele et al., 2015b; Masubelele et al., 2015a) causing fire in areas seldom burned (Coates et al., 2016).

31

Drought, warming, and land management interactions have caused vegetation mortality (see section 2.4.4.3) and reduced vegetation cover in shrublands as projected by AR4 (Burrell et al., 2020). Increased heat and drought are causing succulent species health and abundance to decline (Musil et al., 2009; Schmiedel et al., 2012; Aragón-Gastélum et al., 2014; Koźmińska et al., 2019). Hot droughts especially reduce population resilience (*medium confidence*) (Koźmińska et al., 2019).

37 38

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2.4.3.4 Observed Changes in Mediterranean-Type Ecosystems

Since AR5, Settele et al. (2014) found that all five Mediterranean-Type Ecosystems (MTEs) of the world experienced extreme droughts within the past decade, with South Africa and California reporting the worst on record (*robust evidence, high agreement*) (Diffenbaugh et al., 2015; Williams et al., 2015a; Garreaud et al., 2017; Otto et al., 2018; Sousa et al., 2018). Climate change is causing these droughts to become more frequent and severe (*medium evidence, medium agreement*) (AghaKouchak et al., 2014, Garreaud et al 2017 The 2010-2015 megadrought, AR6 WGI Chpt 11; Otto et al., 2018).

MTEs show a range of direct responses to various forms of water deficit, but have also been affected by
increasing fire activity linked to drought (Abatzoglou and Williams, 2016), and interactions between drought
or extreme weather and fire, affecting post-fire ecosystem recovery (Slingsby et al., 2017). Responses
include shifts in functional composition (Acácio et al., 2017; Syphard et al., 2019a), decline in vegetation
health (Hope et al., 2014; Asner et al., 2016a), decline or loss of characteristic species (White et al., 2016;
Stephenson et al., 2019), shifts in composition towards more drought- or heat-adapted species and declining
diversity (also see Section 2.4.4.3; Slingsby et al., 2017.; Harrison et al., 2018).

54

55 Declines in plant health and increased mortality in MTEs associated with drought have been widely 56 documented (*robust evidence, high agreement*) (Section 2.4.4.3). Remote sensing studies show drought

documented (*robust evidence, high agreement*) (Section 2.4.4.3). Remote sensing studies show drought
 associated mortality in postfire vegetation regrowth in the Fynbos of South Africa (Slingsby et al., 2020b),

reduced canopy health in forests within MTE zones of South Africa (Hope et al., 2014), and declines in 1 canopy water content in forests of California (Asner et al., 2016a). Several studies reported climate-2 associated responses of dominant or charismatic species. High mortality in the Clanwilliam Cedar between 3 1931 to 2013 occurred at lower, hotter elevations in the Fynbos of South Africa (White et al., 2016). Drought 4 reduced growth and increased mortality of the holm oak, Quercus ilex, in the Iberian Peninsula of Spain, 5 Natalini et al. (2016). Portuguese shrublands experienced losses of many deciduous and evergreen oak 6 species, and increasing dominance of to pyrophytic xeric trees (Acácio et al., 2017). The 2012-2015 drought 7 in California caused: high canopy foliage die-back of the Giant Sequoia (Sequoiadendron giganteum) 8 (Stephenson et al., 2019), increased the dominance of oaks relative to pines resulting from increased water 9 deficit, and large-scale mortality from drought and insect pest outbreak interactions (McIntyre et al., 2015; 10 Fettig et al., 2019). 11 12

Species distribution or community composition changes have contributed to declines in diversity and/or 13 shifts towards more drought- or heat-adapted species (medium evidence, high agreement). Two conifer 14 species (Pinus longaeva, P. flexilis) shifted upslope 19 m from 1950 to 2016 in the Great Basin, USA, 15 (Smithers et al., 2018). Reduced winter precipitation caused native annual forbs to recede resulting in long-16 lasting and potentially unidirectional reductions in diversity in a Californian grassland (Harrison et al., 17 2018). More frequent extreme hot and dry weather between 1966 and 2010 caused declines in diversity 18during the post-fire regeneration phase in the Fynbos of South Africa (Slingsby et al., 2017) resulting in 19 shifts towards species with higher temperature preferences (Slingsby et al., 2017). In Italy, Del Vecchio et al. 20 (2015) observed increases in plant cover and thermophilic species in coastal foredune habitats between 1989 21 and 2012. 22

23 In southern California, USA, areas of forest and woody shrublands are shifting to grasslands, driven by a 24 combination of climate and land use factors such as increased drought, fire ignition frequency and increases 25 in nitrogen deposition (robust evidence, high agreement) (Jacobsen and Pratt, 2018; Park et al., 2018; Park 26 and Jenerette, 2019; Syphard et al., 2019b). 27

28

The effects of climate change on heat, fuel, and wildfire ignition limitations show spatial and temporal 29 variation globally (see Section 2.3.6.1), but there have been a number of observed impacts in MTEs (medium 30 evidence, high agreement). Climate change has caused increases in fuel aridity and area burned by wildfire 31 across the western United States from 1985 to 2015 (Abatzoglou and Williams, 2016). Local and global 32 climatic variability led to a 4 year decrease in the average fire return time in Fynbos, South Africa when 33 comparing fires recorded between 1951 to 1975 and 1976 to 2000 (Wilson et al., 2010). For Chile, González 34 et al. (2018) reported a significant increase in the number, size, duration and simultaneity of large fires 35 during the 2010 to 2015 "megadrought" when compared to the 1990 to 2009 baseline. 36

37

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2.4.3.5 Observed Changes in Savanna and Grasslands 38

Savannas consist of coexisting trees and grasses in the tropics and temperate regions (Archibald et al., 2019). 40 The global trend of woody encroachment reported in AR5 (Settele et al., 2014) is continuing (robust 41 evidence, high agreement, very high confidence) (see Table 2.S.1), with increases occurring in: temperate 42 savannas in North America (10-20% per decade) (Archer et al., 2017), tropical savannas in South America 43 (8% per decade), Africa (2.4% per decade) and Australia (1% per decade) (O'Connor et al., 2014; Espírito-44 Santo et al., 2016: Skowno et al., 2017: Stevens et al., 2017: McNicol et al., 2018: Venter et al., 2018: Rosan 45 et al., 2019). Additionally, forest expansion into mesic savannas reported in AR5 (Settele et al., 2014), is 46 continuing in Africa, South America and Southeastern Asia (Marimon et al., 2014; Keenan et al., 2015; 47 Baccini et al., 2017; Ondei et al., 2017; Stevens et al., 2017; Aleman et al., 2018; Rosan et al., 2019). 48 49 Extreme high rainfall anomalies also contributed to an increase in herbaceous and foliar production in the Sahel (Brandt et al., 2019; Zhang et al., 2019a). 50

51

New studies since AR5, using multiple study designs (experimental manipulations in lab and field, meta-52 analyses and modelling), attribute climate change increases in woody cover to elevated atmospheric CO_2 53

(Donohue et al., 2013; Nackley et al., 2018; Quirk et al., 2019) and increased rainfall amount and intensity 54

(robust evidence, high agreement) (Venter et al., 2018; Xu et al., 2018b; Zhang et al., 2019a). Direct 55

quantification of climate change drivers is confounded with local land use changes such as fire suppression 56

Chapter 2 FINAL DRAFT IPCC WGII Sixth Assessment Report removal of native browsers, and specifically loss of mega-herbivores in Africa (medium evidence, medium 1 agreement) (Asner et al., 2016b; Daskin et al., 2016; Stevens et al., 2016; Davies et al., 2018). The relative 2 importance of the climate- and non-climate-related causes of woody plant vary between regions, but there is 3 general agreement that climate change impacts, specifically, increasing rainfall and rising CO₂, are frequent 4 and strong contributing factors of woody cover increase (robust evidence, high agreement). 5 6 Extensive woody cover increases in non-forested biomes is reducing grazing potential (Smit and Prins, 7 2015), and changing the carbon stored per unit land area (González-Roglich et al., 2014; Puttock et al., 2014; 8 Pellegrini et al., 2016; Mureva et al., 2018) and hydrological characteristics (Honda and Durigan, 2016; 9 Schreiner-McGraw et al., 2020). Woody cover encroachment also reduces biodiversity by threatening fauna 10and flora adapted to open ecosystems (Ratajczak et al., 2012; Smit and Prins, 2015; Pellegrini et al., 2016; 11 Andersen and Steidl, 2019). 12 13 The global extent of grasslands is declining significantly because of climate change (medium confidence). In 14 temperate and boreal zones, where about half of treelines are shifting, they are overwhelmingly expanding 15 poleward and upward, with accompanying loss of montane grassland (*robust evidence, high agreement*); 16 whereas tropical treelines have been generally stable (medium evidence, medium agreement) (Harsch et al., 17 2009; Rehm & Feeley 2015; Silva et al., 2016; Andela et al., 2017; Song et al., 2018; Aide et al., 2019; 18 Gibson and Newman, 2019). The Eurasian steppes experienced a 1% increase in woody cover per decade 19 since 2000 (Liu et al., 2021) and Inner Mongolian grasslands in China experienced broad encroachment as 20 well (Chen et al., 2015). Climatic drivers of woody expansion in temperature limited grasslands, particularly 21 alpine grasslands, are most frequently attributed to warming (robust evidence, high agreement, high 22 confidence) (D'Odorico et al., 2012; Hagedorn et al., 2014), increases in water and nutrient availability from 23 thawing permafrost (medium evidence, high agreement) (Zhou et al., 2015b; Silva et al., 2016) and rising 24

CO₂ (*medium evidence, medium agreement*) (Frank et al., 2015; Aide et al., 2019). Interactions between land use changes: land abandonment, grazing management shifts, and fire suppression, and climate change are

contributing factors (Liu et al., 2021)

28 Remote sensing shows overall increasing trends in both the annual maximum NDVI and annual mean NDVI 29 in global grasslands ecosystems between 1982 and 2011 (Gao et al., 2016). Multiple lines of evidence 30 indicate that changes in grassland productivity are positively correlated with increases in mean annual 31 precipitation (Hoover et al., 2014; Brookshire and Weaver, 2015; Gang et al., 2015; Gao et al., 2016; Wilcox 32 et al., 2017; Wan et al., 2018). Increasing temperatures positively impact grassland production and biomass, 33 especially in temperature limited regions (Piao et al., 2014; Gao et al., 2016). However, grasslands in hot 34 areas are expected to decrease production with increases in temperature (limited evidence, low agreement) 35 (Gang et al., 2015). Nevertheless, grassland responses to warming and drought are being ameliorated by 36 increasing CO_2 and associated improved water use efficiency (Roy et al., 2016). For example, in a cool 37 temperate grassland experiment, warming led to a longer growing season and elevated CO₂ further extended 38 growing by conserving water, which enabled most species to remain active longer (medium evidence. 39 medium agreement) (Reyes-Fox et al., 2014). 40

42 2.4.3.6 Observed Changes in Tropical Forest

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Overall declines of tropical forest cover (Kohl et al., 2015; Liu et al., 2015; Baccini et al., 2017; Harris et al., 2021), with declines more than triple the gains (Harris et al., 2021) have been driven primarily by
deforestation and land conversion (*robust evidence, high agreement*) (Lewis et al., 2015; Curtis et al., 2018;
Espaciais, 2021). In opposition to this general trend, expansion of tropical forest cover into savannas and
grasslands has occurred in Africa, South America, and Australia (Baccini et al., 2017; Aleman et al., 2018;
Staver, 2018) (Marimon et al., 2014; Ondei et al., 2017; Stevens et al., 2017; Rosan et al., 2019).

Specific examples of climate-change driven range shifts of tropical deciduous forests upslope into alpine grasslands have been documented in the Americas (Chacón-Moreno et al., 2021; Jiménez-García et al., 2021) and in Asia (Sigdel et al., 2018). However, treeline behaviours are diverse. A study in Nepal recorded that treeline fomed by *Abies spectabilis* had been stable for more than a century, while the upper limit of large shrubs (*Rhododendron campanulatum*) had been advancing (Mainali et al., 2020). In both the Andes

⁵⁶ (Harsch et al., 2009) and Himalaya (Singh et al., 2021) most treelines have been stable, leading (Rehm &

that have occurred are probably driven by interactions between changing land-use, such as fire suppression,
and climate changes such as increased rainfall, warming and elevated CO₂ either through CO₂ fertilisation or
increases in water-use efficiency (*medium evidence, medium agreement*) (Cernusak et al., 2013; Huang et al.,
2013; Van Der Sleen et al., 2015; Yang et al., 2016).

5

Increases in productivity of tropical forests (Gatti et al., 2014; Brienen et al., 2015; Baccini et al., 2017),
Africa and SE Asia (Qie et al., 2017) have been attributed to elevated CO₂ (*robust evidence, medium agreement*) (Ballantyne et al., 2012; Brienen et al., 2015; Sitch et al., 2015; Yang et al., 2016; Mitchard,
2018). The rates of these increases have been slowing down in the central Amazon (Brienen et al., 2015; de
Meira Junior et al., 2020) and SE Asia (Qie et al., 2017). In contrast, the carbon sink (and hence rate of
biomass gain) in intact African forests was stable until 2010 and has only recently started to decline,

indicating asynchronous carbon sink saturation in Amazonia and Africa, the difference driven by rates of
 tree mortality (Hubau et al. 2020). At a global level (Hubau et al 2020) argue that the carbon sink associated

14 with intact tropical forests peaked in the 1990s and is now in decline.

15

Declines in productivity are most strongly associated with warming (Sullivan et al., 2020), reduced growth 16 rates during droughts (Bennett et al., 2015; Bonai et al., 2016; Corlett, 2016), drought related mortality 17 (Brando et al., 2014; Zhou et al., 2014; Brienen et al., 2015; Corlett, 2016; McDowell et al., 2018), fire (Liu 18et al., 2017), and cloud-induced radiation-limitation (robust evidence, high agreement) (Deb Burman et al., 19 2020). Increases in frequency and severity of droughts and shorter tree residence times due to increases in 20 growth rates caused by elevated CO₂ may be additional interactive factors increasing tree mortality (Malhi et 21 al., 2014; Brienen et al., 2015). Vulnerability to drought varies between tree species and sizes with large, 22 long-lived trees at highest risk of mortality (McDowell et al., 2018; Meakem et al., 2018). Mortality risk also 23 varies between forest types with seasonal rainforests appearing most vulnerable to drought (Corlett, 2016). 24

Lianas (long-stemmed woody vines) generally negatively impact trees, significantly reducing the growth of
heavily infested trees (Reis et al., 2020). They would benefit from climate change and disturbance (LingZi et
al., 2014; Hodgkins et al., 2018). The extent of their suitable niche can increase (Taylor and Kumar, 2016),

- al., 2014, Hodgkins et al., 2018). The extent of their suitable inche can increase (Taylor and Rumar, 2016),
 thereby decreasing forest biomass accumulation (*robust evidence, high agreement*) (van der Heijden et al.,
 2013; Fauset et al., 2015; Estrada-Villegas et al., 2020).
- 31

Climate change continues to degrade forests by reducing resilience to pests and diseases, increasing species 32 invasion, facilitating pathogen spread (Malhi et al., 2014; Deb et al., 2018) and intensifying fire risk and 33 potential die-back (Lapola et al., 2018; Marengo et al., 2018). Drought, temperature increases and forest 34 fragmentation interact to increase the prevalence of fires in tropical forests (robust evidence, high 35 agreement). Warming increases water stress in trees (Corlett, 2016) and together with forest fragmentation, 36 dramatically increases desiccation of forest canopies-resulting in deforestation that then leads to even 37 hotter and drier regional climates (Malhi et al., 2014; Lewis et al., 2015). Warming and drought increase 38 invasion of grasses into forest edges and increase fire risk (robust evidence, high agreement) (Brando et al., 39 2014; Balch et al., 2015; Lewis et al., 2015). Droughts and fires additively increase mortality and, 40 consequently, reduce canopy cover and aboveground biomass (Cross-Chapter Paper 7; Brando et al., 2014, 41 2020; Balch et al., 2015; Lewis et al., 2015). 42

43 44

2.4.3.7 Observed Changes in Boreal and Temperate Forests

45 The IPCC Fifth Assessment Report found increased tree mortality, wildfire and plant phenology changes in 46 boreal and temperate forests (Settele et al., 2014). Expanding on those conclusions, this Assessment, using 47 analyses of causal factors, attributes to anthropogenic climate change the following observed changes in 48 boreal and temperate forests in the 20th and 21st centuries: upslope and poleward biome shifts at sites in Asia, 49 Europe, and North America (Section 2.4.3.2.1); range shifts of plants (Section 2.4.2.1); earlier blooming and 50 leafing of plants (Section 2.4.2.4); poleward shifts in tree-feeding insects (Section 2.4.2.1); increases in 51 insect pest outbreaks (Section 2.4.4.3.3); increases in area burned by wildfire in western North America 52 (Section 2.4.4.2.1); increased drought-induced tree mortality in western North America (Section 2.4.4.3.1); 53 and thawing of permafrost that underlies extensive areas of boreal forest (IPCC Sixth Assessment Report, 54 Working Group I, Chapter 2, Section 2.4.3.9). Atmospheric CO₂ from anthropogenic sources has also 55

increased net primary productivity (Section 2.4.4.5.1). In summary, anthropogenic climate change has caused substantial changes to temperate and boreal forest ecosystems, including biome shifts and increases

1	in wildfire, insect pest outbreaks, and tree mortality, at a global mean surface temperature increase of 0.9° C
2	above the pre-industrial period (robust evidence, high agreement).
3	
4	Other changes detected in boreal forests and consistent with, but not formally attributed to climate change,
5	include increased wildfire in Siberia (Section 2.4.4.2.3), long-lasting smoldering belowground fires in
6	Canada and the United States (Scholten et al., 2021), tree mortality in Europe (Section 2.4.4.3.3), and post-
7	fire shifts of boreal conifer to deciduous broadleaf tree species in Alaska (Mack et al., 2021). From 1930 to
8	1960, boreal forest growth became limited more by precipitation than temperature in the Northern
9	Hemisphere (Babst et al., 2019).
10	
11	For some vegetation changes, land use and land management changes have exerted more influence than
12	climate change. These include upslope and poleward forest shifts in Europe following abandonment of
13	timber harvesting or livestock grazing (Section 2.4.3.2.2), changes in wildfire in Europe affected by fire suppression, fire prevention, and agricultural abandonment (Section 2.4.4.2.3), and forest species
14	composition changes in Scotland due to nitrogen deposition from air pollution (Hester et al., 2019). Remote
15 16	sensing suggests that the area of temperate and boreal forests increased in Asia and Europe between 1982
17	and 2016 (Song et al., 2018) and in Canada between 1984 and 2015 (Guindon et al., 2018), but forest
17	plantations and regrowth are probable drivers (Song et al., 2018).
19	planations and regiowin are producte arrivers (boing et al., 2010).
20	2.4.3.8 Observed Changes in Peatlands
21	
22	Globally, peatland ecosystems store approximately 25% (600±100 GtC) of the world's soil organic carbon
23	(Yu et al., 2010; Page et al., 2011; Hugelius et al., 2020) and 10% of the world's freshwater resources
24	(Joosten and Clarke, 2002), despite only occupying 3% of the global land area (Xu et al., 2018a). The long-
25	term role of northern peatlands in the carbon cycle was mentioned for the first time in IPCC AR4 (IPCC,
26	2007b), while SR1.5 briefly mentioned the combined effects of climate and land-use change on peatlands
27	(IPCC, 2018b). New evidence confirms that climate change, including extreme weather events (e.g.,
28	droughts; Section 8.3.1.6), permafrost degradation (Section 2.3.2.5), sea-level rise (Section 2.3.3.3), and fire
29	(Section 5.4.3.2) (Henman and Poulter, 2008; Kirwan and Mudd, 2012; Turetsky et al., 2015; Page and
30	Hooijer, 2016; Swindles et al., 2019; Hoyt et al., 2020; Hugelius et al., 2020; Jovani-Sancho et al., 2021;
31	Veraverbeke et al., 2021), superimposed on anthropogenic disturbances (for example, draining for
32	agriculture or mining; Section 5.2.1.1), has led to rapid losses of peatland carbon across the world (<i>robust</i>
33	evidence, high agreement) (Page et al., 2011; Leifeld et al., 2019; Hoyt et al., 2020; Turetsky et al., 2020;
34	Loisel et al., 2021). Other essential peatland ecosystem services, such as water storage and biodiversity, are
35	also being lost worldwide (<i>robust evidence, high agreement</i>) (Bonn et al., 2014; Martin-Ortega et al., 2014; Tiemeyer et al., 2017).
36 37	
38	The switch from carbon sink to source in peatlands globally is mainly attributable to changes in water table
38 39	depth, regardless of management or status (<i>robust evidence, high agreement</i>) (Lafleur et al., 2005; Dommain
40	et al., 2011; Lund et al., 2012; Cobb et al., 2017; Evans et al., 2021; Novita et al., 2021). Across the
41	temperate and tropical biomes, extensive drainage and deforestation have caused widespread water table
42	drawdowns and/or peat subsidence, as well as large CO ₂ emissions (<i>medium evidence, high agreement</i>).
43	Climate change is compounding these impacts (medium evidence, medium agreement). For example, in
44	Indonesia, the highest emissions from drained tropical peatlands were reported in the extremely dry year of
45	the 1997 El Niño (810-2570 TgC yr ⁻¹) (Page et al., 2002) and the 2015 fire season (380 TgC yr ⁻¹) (Field et
46	al., 2016). These prolonged dry seasons have also led to tree die-offs and fires, which are relatively new
47	phenomena in these latitudes (medium evidence, high agreement) (Cole et al., 2015; Mezbahuddin et al.,
48	2015; Fanin and van der Werf, 2017; Taufik et al., 2017; Cole et al., 2019). Low soil moisture contributes to
49	increased fire propagation (see Cross-Chapter Box 5, and Section 12.4.2.2; Dadap et al., 2019), causing
50	long-lasting fires responsible for smoke and haze pollution (robust evidence, high agreement) (Ballhorn et
51	al., 2009; Page et al., 2009; Gaveau et al., 2014; Huijnen et al., 2016; Page and Hooijer, 2016; Hu et al.,
52	2018; Vadrevu et al., 2019; Niwa et al., 2021). Increases in fires and smoke lead to habitat loss and
53	negatively impact regional faunal populations (<i>limited evidence, high agreement</i>) (Neoh et al., 2015; Erb et
54	al., 2018b; Thornton et al., 2018).
55	To here the dama test and here to do a the data of the
56 57	In large lowland tropical peatland basins that are less impacted by anthropogenic activities (i.e., Amazon and Congo river basins), the direct impact of climate change is that of a decreased carbon sink (<i>limited avidance</i>).

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Congo river basins), the direct impact of climate change is that of a decreased carbon sink (limited evidence,

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medium agreement) (Roucoux et al., 2013; Gallego-Sala et al., 2018; Wang et al., 2018a; Dargie et al., 2019; 1 Ribeiro et al., 2021). As for the temperate and boreal regions, climatic drying also tends to promote peat 2 oxidation and carbon loss to the atmosphere (medium evidence, medium agreement) (section 2.3.1.3.4) 3 (Helbig et al., 2020; Zhang et al., 2020). In Europe, increasing mean annual temperatures in the Baltic, 4 Scandinavia, and Continental Europe (Section 12.4.5.1) have led to widespread lowering of peatland water 5 tables at intact sites (Swindles et al., 2019), Sphagnum moss desiccation and die off (Bragazza, 2008; Lees et 6 al., 2019), and increased fire intensity and frequency resulting in rapid carbon loss (Davies et al., 2013; 7 Veraverbeke et al., 2021). Nevertheless, longer growing seasons and warmer, wetter climates have increased 8 carbon accumulation and promoted thick deposits regionally, as reported for some North American sites 9 (limited evidence, medium agreement) (Cai and Yu, 2011; Shiller et al., 2014; Ott and Chimner, 2016). 1011 In high-latitude peatlands, the net effect of climate change on the permafrost peatland carbon sink capacity 12 remains uncertain (Abbott et al., 2016; McGuire et al., 2018b; Laamrani et al., 2020; Loisel et al., 2021; Sim 13 et al., 2021; Väliranta et al., 2021). Increasing air temperatures have been linked to permafrost degradation 14 and altered hydrological regimes (Section 2.3.3.2, Figure 2.4a, 2.4.3.9, and Box 5.1), which have led to rapid 15 changes in plant communities and biogeochemical cycling (robust evidence, high agreement) (Liljedahl et 16 al., 2016; Swindles et al., 2016; Voigt et al., 2017; Zhang et al., 2017b; Voigt et al., 2020; Sim et al., 2021). 17 In many instances, permafrost degradation triggers thermokarst land subsidence associated with local 18 wetting (robust evidence, high agreement) (Jones et al., 2013; Borge et al., 2017; Olvmo et al., 2020; 19 Olefeldt et al., 2021). Permafrost thaw in peatland-rich landscapes can also cause local drying through 20 increased hydrological connectivity and runoff (Connon et al., 2014). In the first decades following thaw, 21 increases in methane, CO₂, and nitrous oxide emissions have been recorded from peatland sites, depending 22 on surface moisture conditions (Schuur et al., 2009; O'Donnell et al., 2012; Elberling et al., 2013; Matveev 23 et al., 2016; Euskirchen et al., 2020; Hugelius et al., 2020). Conversely, some evidence suggests increased 24 peat accumulation after thaw (Jones et al., 2013; Estop-Aragonés et al., 2018; Väliranta et al., 2021). There 25 is also a need to consider the impact of wildfire on permafrost thaw, due to its effect on soil temperature 26 regime (Gibson et al., 2018), wildfire as a}, as fire intensity and frequency have increased across the boreal 27 and Arctic biomes (limited evidence, high agreement) (Kasischke et al., 2010; Scholten et al., 2021). 28 29 Unfortunately, the CO₂ emissions from degrading peatlands is contributing to climate change in a positive 30 feedback loop (robust evidence, high agreement). In the midlatitudes, widespread anthropogenic disturbance 31 led to large historical GHG emissions and current legacy emissions of 0.15 PgC yr⁻¹ between 1990 and 2000 32 (limited evidence, high agreement) (Maljanen et al., 2010; Tiemeyer et al., 2016; Drexler et al., 2018; Qiu et 33 al., 2021). About 80 million ha of peatlands have been converted to agriculture, equivalent to 72 PgC 34 emissions between 850–2010 CE (Leifeld et al., 2019; Qiu et al., 2021). In southeast Asia, an estimated 20– 35 25 Mha of peatlands have been converted to agriculture with carbon currently being lost at a rate of 36 \sim 155 ± 30 MtC yr⁻¹ (Miettinen et al., 2016; Leifeld et al., 2019; Hoyt et al., 2020). Extensive deforestation 37 and drainage have caused widespread peat subsidence and large CO₂ emissions at a current average of $\sim 10 \pm$ 38 2 t ha⁻¹ yr⁻¹ (excluding fires, (Hoyt et al., 2020)), with values estimated from point subsidence measurements 39

being as high as 30-90 t CO₂ ha⁻¹ yr⁻¹ locally (*robust evidence, high agreement*) (Wösten et al., 1997; Matysek et al., 2018; Swails et al., 2018; Evans et al., 2019; Conchedda and Tubiello, 2020; Anshari et al., 2021). On balance, at the global scale, increases in GHG emissions from peatlands have primarily come from the compounded effects of land-use change, drought, and fire, with emissions from some thawing permafrost peatlands (*robust evidence, high agreement*).

46 2.4.3.9 Observed Changes in Polar Tundra

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Warming at high latitudes, documented in both AR4 and AR5, is leading to earlier snow and sea ice melt and 48 longer growing seasons (WGI AR6) which are continuing to alter tundra plant communities (medium 49 evidence, high agreement) (Post et al., 2009; Gauthier et al., 2013). Woody encroachment and increases in 50 vegetation productivity observed in both AR4 and AR5 are widespread and continuing. Both experiments 51 and monitoring indicate that climate warming is causing increases in shrub, grass and sedge abundance, 52 density, frequency and height, with decreases in mosses and/or lichens (robust evidence, high agreement) 53 (Myers-Smith et al., 2011; Bjorkman et al., 2018; Bjorkman et al., 2019). Shrub growth is climate-sensitive 54 and greater in years with warmer growing seasons (Myers-Smith et al., 2015). Plant species that prefer 55 warmer conditions are increasing (Elmendorf et al., 2015; Bjorkman et al., 2018), plant cover is increasing 56 and bare ground is decreasing in long-term monitoring plots (Bjorkman et al., 2019; Myers-Smith et al., 57

In addition to direct warming, indirect effects of climate change such as thawed permafrost, altered 5 hydrology and enhanced nutrient cycling (as observed in AR4 and AR5) continue and are causing 6 pronounced vegetation changes (medium evidence, medium agreement) (Schuur et al., 2009; Natali et al., 7 2012). Soil moisture status influences temperature sensitivity of plant growth and canopy heights (Myers-8 Smith et al., 2015; Ackerman et al., 2017; Bjorkman et al., 2018). In tundra ecosystems permafrost thawing 9 can decouple below-ground plant growth dynamics from above-ground dynamics, with below-ground root 10 growth continuing until soils refreeze in autumn (Cross-Chapter Paper 6; Iversen et al., 2015; Blume-Werry 11 et al., 2016; Radville et al., 2016). 12

2.4.4 **Observed Changes in Ecosystem Processes and Services** 14

2.4.4.1 Observed Browning of Rivers and Lakes 16

17 In boreal coniferous areas there has been an increase in terrestrial derived dissolved organic carbon (DOC) 18 transport into rivers and lakes, which has caused increased opacity and shift toward a brown colour 19 (browning). This process was not given much attention in AR5 even though it is a consequence of climate 20 change: hydrological intensification, greening of the Northern Hemisphere, and degradation of carbon sinks 21 in peatlands (robust evidence, high agreement) (Solomon et al., 2015; Catalán et al., 2016; Crowther et al., 22 2016; de Wit et al., 2016; Finstad et al., 2016; Creed et al., 2018; Hayden et al., 2019) factors that enhance 23 terrestrial productivity, alter vegetation communities, and affect the hydrological control on production and 24 transport of DOC (Weyhenmeyer et al., 2016). Non climate-related drivers of browning are: declining 25 atmospheric sulphur deposition, forestry practices and land-use changes (see Table 2.S.1 for detail). 26

27 Browning creates a positive feedback by absorbing photosynthetically active radiation accelerating upper 28 water (epilimnetic) warming (Solomon et al., 2015). Browning of lakes leads to shallower and more stable 29 thermoclines and thus, overall deep water cooling (Solomon et al., 2015; Williamson et al., 2015) and can 30 provoke a transition of the seasonal mixing regime from a mixed lake (polymictic) to one that is seasonally 31 stratified (Kirillin and Shatwell, 2016). 32

33 The ecological responses of browning need to be considered as concomitant effects of climate change and 34 nutrient status. Results from long-term, large-scale lake experiments were variable, showing both strong 35 synergistic effects (Urrutia-Cordero et al., 2016) and no significant effects of browning on plankton 36 community food webs (Rasconi et al., 2015). Browning has driven a shift from auto- to 37 heterotrophic/mixotrophic-based production (Wilken et al., 2013; Urrutia-Cordero et al., 2017) and supports 38

heterotrophic metabolism of the bacterial community (Zwart et al., 2016). Browning may also accelerate 39 primary production through input of nutrients associated with DOM in nutrient poor lakes and increase 40 cyanobacteria, which better cope with low light intensities (Huisman et al., 2018) and toxin levels (Urrutia-41 Cordero et al., 2016). However, the synergistic impacts of browning and climate change on aquatic 42 communities depends on regional precipitation patterns (Weyhenmeyer et al., 2016), watershed type (de Wit 43 et al., 2016), and food chain length (Hansson et al., 2013). Quantitative attribution of browning to climate 44

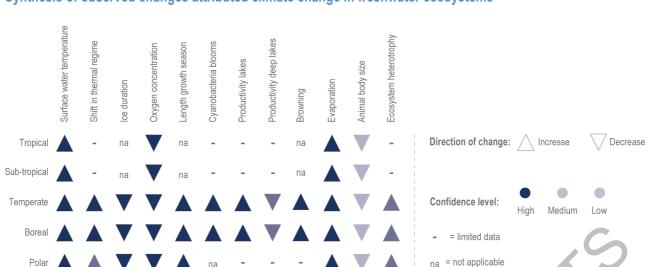
- change remains difficult (medium evidence, medium confidence). 45
- 46

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13

- In summary, new studies since AR5 have explicitly estimated the effects of warming and browning on 47 freshwaters in boreal areas with complex positive and negative repercussions on water temperature profiles 48
- (lower vs upper water) (high confidence) and primary production (medium confidence). 49
- 50





Synthesis of observed changes attributed climate change in freshwater ecosystems

Figure 2.5: Large scale observed changes in freshwater ecosystems attributed to climate change over more than four decades. For description and references, see 2.3.3, 2.4.2, and 2.5.3.6.2.

2.4.4.2 Observed Changes in Wildfire

2.4.4.2.1 Detection and attribution of observed changes in wildfire

Wildfire is a natural and essential component of many forest and other terrestrial ecosystems. Excessive 9 wildfire, however, can kill people, cause respiratory disease, destroy houses, emit carbon dioxide, and 10 damage ecosystem integrity (see Section 2.4.4.2, 2.4.4.4). Anthropogenic climate change increases wildfire 11 by exacerbating its three principal driving factors-heat, fuel, and ignition (Moritz et al., 2012; Jolly et al., 12 2015). Non-climatic factors also contribute to wildfires—in tropical areas fires are set intentionally to clear 13 forest for agricultural fields and livestock pastures (Bowman et al., 2020b). Urban areas and roads create 14 ignition hazards. Governments in many temperate zone countries implement policies to suppress fires, even 15 natural ones, producing unnatural accumulations of fuel in the form of coarse woody debris and high 16 densities of small trees (Ruffault and Mouillot, 2015; Hessburg et al., 2016; Andela et al., 2017; Balch et al., 17 2017; Lasslop and Kloster, 2017; Aragao et al., 2018, (Kelley et al., 2019). Globally, 4.2 million km² of land 18 per year burned on average from 2002 to 2016 (Giglio et al., 2018) with the highest fire frequencies in the 19 Amazon rainforest, deciduous forests and savannas in Africa, and deciduous forests in northern Australia 20 (Earl and Simmonds, 2018; Andela et al., 2019). 21

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Since the IPCC Fifth Assessment Report and the IPCC Special Report on Land, published research has 23 detected increases in the area burned by wildfire, analysed relative contributions of climate and non-climate 24 factors, and attributed burned area increases above natural levels to anthropogenic climate change in one part 25 of the world – western North America (robust evidence, high agreement) (Abatzoglou and Williams, 2016; 26 Partain et al., 2016; Kirchmeier-Young et al., 2019; Mansuy et al., 2019; Bowman et al., 2020b). Across the 27 western United States, increases in vegetation aridity due to higher temperatures from anthropogenic climate 28 change doubled burned area from 1984 to 2015 over what would have burned due to non-climate factors, 29 including unnatural fuel accumulation from fire suppression, with the burned area attributed to climate 30 change accounting for 49% (32-76%, 95% confidence interval) of cumulative burned area (Abatzoglou and 31 Williams, 2016). Anthropogenic climate change has doubled the severity of a southwest North American 32 33 drought from 2000 to 2020 that has reduced soil moisture to its lowest levels since the 1500s (Williams et al., 2020), driving half of the increase in burned area (Abatzoglou and Williams, 2016; Holden et al., 2018; 34 Williams et al., 2019). In British Columbia, Canada, the increased maximum temperatures due to 35 anthropogenic climate change increased burned area in 2017 to its highest extent in the 1950–2017 record, 36 seven to eleven times the area that would have burned without climate change (Kirchmeier-Young et al., 37 2019). In Alaska, USA, the high maximum temperatures and extremely low relative humidity due to 38 anthropogenic climate change accounted for 33–60% of the probability of wildfire in 2015, when the area 39 40 burned was the second highest in the 1940–2015 record (Partain et al., 2016). In protected areas of Canada and the United States, climate factors (temperature, precipitation, relative humidity, evapotranspiration) 41

4	In summary, neid evidence snows that antihopogenic crimate change has increased the area burned by
5	wildfire above natural levels across western North America in the period 1984–2017, at global mean surface
6	temperature increases of 0.6°C –0.9°C, increasing burned area up to 11 times in one extreme year and
7	doubling burned area over natural levels in a 32 year period (<i>high confidence</i>).
	douoning outlied area over haddal levels in a 52 year period (<i>mgn conjutence</i>).
8	2.4.4.2.2 Observed changes in wildfire globally
9	For global terrestrial area as a whole, wildfire trends vary depending on the time period of analysis. From
10	
11	1900 to 2000, global average fire frequency, based on field data, increased 0.4% but the change was not
12	statistically significant, (Mouillot and Field, 2005; Gonzalez et al., 2010b). Fire frequency increased on one-
13	third of global land, mainly from burning for agricultural clearing in Africa, Asia, and South America,
14	slightly less than the area of fire frequency decrease, mainly from fire suppression across Australia, North
15	America, and Russia (Gonzalez et al., 2010b). Analyses of the Global Fire Emissions Database document
16	that from 1996 to 2015, global burned area decreased at a rate of -0.7% y ⁻¹ (Forkel et al., 2019) but the
17	change was not statistically significant, (Giglio et al., 2013). From 1998 to 2015, global burned area
18	decreased at a rate of $-1.4 \pm 0.5\%$ y ⁻¹ (Giglio et al., 2013; Andela et al., 2017). The area of fire increases was
19	a third of the area of decreases, due to reduction of vegetation cover from agricultural expansion and
20	intensification (Andela et al., 2017) and increased precipitation (Forkel et al., 2019). Furthermore, much of
21	the decreasing trend derives from two years: 1998 with high burned area and 2013 with low burned area
22	(Forkel et al., 2019). Wildfire does not show a clear long-term trend for the world as a whole because of
23	increases and decreases in different regions (<i>medium evidence, medium agreement</i>).
23 24	mereases and decreases in anterent regions (meature evidence, meature agreement).
24 25	Where global average burned area has decreased in the past two decades, higher correlations of rates of
	change in burning to human population density, cropland area, and livestock density than to precipitation
26	
27	indicate that agricultural expansion and intensification were main causes (Andela et al., 2017). The global
28	decrease of fire frequency from 2000 to 2010 is correlated to increasing human population density (Knorr et
29	al., 2014). The fire-reduction effect of reduced vegetation cover following expansion of agriculture and
30	livestock herding can counteract the fire-increasing effect of increased heat of climate change (Lasslop and
31	Kloster, 2017; Arora and Melton, 2018; Forkel et al., 2019). The reduction of burning needed after the initial
32	clearing for agricultural expansion drives much of the decline in fire in the tropics (Andela et al., 2017; Earl
33	and Simmonds, 2018; Forkel et al., 2019). The human influence on fire ignition can be seen through the
34	decrease documented on holy days (Sundays and Fridays), traditional religious days of rest (Earl et al.,
35	2015). Overall, human land use exerts an influence on wildfire trends for global terrestrial area as a whole
36	that can be stronger than climate change (medium confidence).
37	
38	2.4.4.2.3 Observed changes in wildfire in individual regions
39	While burned area has increased in parts of Asia, Australia, Europe, and South America, published research
40	has not yet attributed the increases to anthropogenic climate change (medium evidence, high agreement).
41	
42	In the Amazon, deforestation for agricultural expansion and the degradation of forests adjacent to deforested
43	areas cause wildfire in moist humid tropical forests not adapted to fire (robust evidence, high agreement)
44	(Fonseca et al., 2017; van Marle et al., 2017; da Silva et al., 2018; da Silva et al., 2021; dos Reis et al., 2021;
45	Libonati et al., 2021). Roads facilitate deforestation, fragmenting the rainforest and increasing the dryness
46	and flammability of vegetation (Alencar et al., 2015). Extreme droughts that occur during warm phases of
47	the El Niño-Southern Oscillation (ENSO) and the Atlantic Multidecadal Oscillation combine with the
48	degradation of vegetation to cause extreme fire events (<i>robust evidence, high agreement</i>) (Fonseca et al.,
49	2017; Aragao et al., 2018; da Silva et al., 2018; Burton et al., 2020; dos Reis et al., 2021; Libonati et al.,
49 50	2017, Alagao et al., 2018, da Silva et al., 2018, Burton et al., 2020, dos Reis et al., 2021, Elbonat et al., 2021). In the State of Roraima, Brazil, distance to roads, infrastructure that enables deforestation, and ENSO
	,
51	were the two factors most explaining fire occurrence in the extreme 2015–2016 fire season (Fonseca et al., 2017) From 1072 to 2014 humad ereg increased in the American existential and the extreme second deformation of the American existence of the extreme second deformation of
52	2017). From 1973 to 2014, burned area increased in the Amazon, coinciding with increased deforestation
53	(van Marle et al., 2017). In the State of Acre, Brazil, burned area increased 36-fold from 1984 to 2016, with
54	43% burned area in agricultural and livestock settlement areas (da Silva et al., 2018). In 2019, the extreme
55	fire year 2019, 85% of the area burned in the Amazon occurred in areas deforested in 2018 (Cardil et al., 2020) Free the set of the

In summary, field evidence shows that anthropogenic climate change has increased the area burned by

local human factors (population density, roads, and built area) (Mansuy et al., 2019).

accounted for 60% of burned area from local human and natural ignitions from 1984 to 2014, outweighing

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- 56 2020). Even though relatively higher moisture in 2019 led to burning below the 2002–2019 average across
- most of South America, burning in areas of recent deforestation in the Amazon were above the 2002-2019

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average, indicating that deforestation, not meteorological conditions, triggered the 2019 fires (Kelley et al.,
 2021; Libonati et al., 2021). Furthermore, from 1981 to 2018, deforestation in the Amazon reduced moisture

- inputs to the lower atmosphere, increasing drought and fire in a self-reinforcing feedback (Xu et al., 2020).
 In the Amazon, deforestation exerts an influence on wildfire that can be stronger than climate change (*robust*)
- 5 evidence, high agreement).
- 6

In Australia, burned area increased significantly between the periods 1950–2002 and 2003–2020 in the 7 southeast state of Victoria, with the area burned in the 2019-2020 bushfires the highest on record 8 (Lindenmayer and Taylor, 2020). In addition to the deaths of dozens of people and destruction of thousands 9 10 of houses, the 2019–2020 Australia bushfires burned almost half of the area protected for conservation in Victoria and two-thirds of forests allocated for timber harvesting (Lindenmayer and Taylor, 2020), wildlife, 11 and extensive areas of habitat for threatened plant and animal species (Geary et al., 2021). Generally, past 12 timber harvesting did not lead to more severe fire canopy damage (Bowman et al., 2021). Across 13 Southeastern Australia, the fraction of vegetated area that burned increased significantly in eight of the 32 14 bioregions from 1975 to 2009 but decreased significantly in three bioregions (Bradstock et al., 2014). 15 Increases in four bioregions were correlated to increasing temperature and decreasing precipitation. 16 Decreases in burned area occurred despite increased temperature and decreased precipitation. Analyses of 17 climate across Australia from 1950 to 2017 (Dowdy, 2018; Harris and Lucas, 2019) and during periods with 18 extensive fires in 2017 in eastern Australia (Hope et al., 2019), in 2018 in Northeastern Australia (Lewis et 19 al., 2020), and in period 2019-2020 in Southeastern Australia (Abram et al., 2021; van Oldenborgh et al., 20 2021) indicate that temperature and drought extremes due to the El Niño-Southern Oscillation, Southern 21 Annular Mode, and other natural interdecadal cycles drive interannual variability of fire weather. While the 22 effects of interdecadal climate cycles on fire are superimposed on long-term climate change, the relative 23 importance of anthropogenic climate change in explaining changes in burned area in Australia remains 24 unquantified (medium evidence, high agreement). 25 26 In Africa, the rate of change of burned area for the continent as a whole ranged from a non-statistically

In Africa, the rate of change of burned area for the continent as a whole ranged from a non-statistically significant -0.45% y⁻¹ from 2002 to 2016 (Zubkova et al., 2019) to a significant -1.9% y⁻¹ from 2001 to 2016 (Wei et al., 2020). Burned area decreases coincided with areas of agricultural expansion or areas where drought reduced fuel loads (Zubkova et al., 2019; Wei et al., 2020). It is possible, however, that the 500 m spatial resolution of Modis remote sensing fire data underestimates burned area in Africa by half by missing small fires (Ramo et al., 2021). In the Serengeti-Mara savanna of east Africa, burned area showed no significant change from 2001 to 2014, although an increase in domestic livestock would tend to reduce the grass cover that fuels savanna fires (Probert et al., 2019).

- 35 In Mediterranean Europe, burned area for the region as a whole decreased from 1985 to 2011 (Turco et al., 36 2016), although burned area for Spain did not show a significant long-term increase from 1968 to 2010 37 (Moreno et al., 2014) while burned area for Portugal in 2017 was the highest in the period 1980–2017 38 (Turco et al., 2019). Increased summer maximum temperature and decreased soil moisture explained most of 39 observed burned area, suggesting a contribution of climate change, but fire suppression, fire prevention, 40 41 agricultural abandonment, and reforestation, and reduction of forest area exerted even stronger influences on burned area than climate across Mediterranean Europe (robust evidence, high agreement) (Moreno et al., 42 2014; Turco et al., 2017; Viedma et al., 2018; Turco et al., 2019). 43
- 44 In the Arctic tundra and boreal forest, where wildfire has naturally been infrequent, burned area showed 45 statistically significant increases of ~50% y⁻¹ across Siberia, Russia, from 1996 to 2015 (Ponomarev et al., 46 2016) and 2% y⁻¹ across Canada from 1959 to 2015 (Hanes et al., 2019). Wildfire burned ~ 6% of the area of 47 four extensive Arctic permafrost regions in Alaska, USA, eastern Canada, and Siberia from 1999 to 2014 48 (Nitze et al., 2018). In boreal forest in the Northwest Territories, Canada, and Alaska, USA, the area burned 49 by wildfire increased at a statistically significant rate of $6.8\% \text{ y}^{-1}$ in the period 1975–2015, (Veraverbeke et 50 al., 2017), with smouldering belowground fires that lasted through the winter covering ~1% of burned area 51 in the period 2002-2016 (Scholten et al., 2021). While burned area was correlated to temperature and 52 reduced precipitation in Siberia (Ponomarev et al., 2016; Masrur et al., 2018) and to lightning, correlated 53 with temperature and precipitation in the Northwest Territories and Alaska (Veraverbeke et al., 2017), no 54 attribution analyses have examined relative influences of climate and non-climate factors. 55 56

In Indonesia, deforestation and draining of peat swamp forests dries out the peat, providing substantial fuel for fires (Page and Hooijer, 2016). Extreme fire years in Indonesia, including 1997, 2006, and 2015, coincide

3 with extreme heat and aridity during the warm phase of the El Niño-Southern Oscillation (Field et al., 2016).

4 Fire-resistant forest in 2019 covered only 3% of peatlands and 4.5% of non-peatlands on Sumatra and

5 Kalimantan (Nikonovas et al., 2020).

In Chile, burned area in the summer of 2016–2017 was 14 times the mean for the period 1985–2016 and the
highest on record (Bowman et al., 2019). While that extreme fire year coincided with the highest daily mean
maximum temperature in the period 1979–2017 (Bowman et al., 2019), in central Chile, the area of highest
fire activity, burned area from 1976 to 2013 showed highest correlation to precipitation cycles of the El
Niño-Southern Oscillation and temperature cycles of the Antarctic Oscillation (Urrutia-Jalabert et al., 2018).

Niño-Southern Oscillation and temperature cycles of the Antarctic Oscillation (Urrutia-Jalabert et al., 2018).

Overall, burned area has increased in the Amazon, the Arctic, Australia, and parts of Africa and Asia, consistent with, but not formally attributed to anthropogenic climate change (*medium evidence, high agreement*). Deforestation, peat draining, agricultural expansion or abandonment, fire suppression, and interdecadal cycles such as the El Niño-Southern Oscillation exert a stronger influence than climate change on wildfire trends in numerous regions outside of North America (*high confidence*).

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19 2.4.4.2.4 Observed changes in fire seasons globally

IPCC AR6 Working Group 1, Chapter 12, has assessed fire weather, while this chapter assesses burned area 20 and fire frequency. The global increases in temperature of anthropogenic climate change have increased 21 aridity and drought, lengthening the fire weather season (annual period with a heat and aridity index greater 22 than half of its annual range) on one-quarter of global vegetated area and increasing average fire season 23 length by one-fifth, from 1979 to 2013 (Jolly et al., 2015). Climate change has contributed to increases in the 24 fire weather season or the probability of fire weather conditions in the Amazon (Jolly et al., 2015), Australia 25 (Dowdy, 2018; Abram et al., 2021; van Oldenborgh et al., 2021), Canada (Hanes et al., 2019), central Asia 26 (Jolly et al., 2015), East Africa (Jolly et al., 2015), and North America (Jain et al., 2017; Williams et al., 27 2019; Goss et al., 2020). In forest areas, the burned area is correlated to fuel aridity, a function of 28 temperature; in non-forest areas, the burned area is correlated to high precipitation in the previous year, 29 which can produce high grass fuel loads (Abatzoglou et al., 2018). Fire use in agriculture and livestock 30 raising or other factors have generated a second fire season on approximately one-quarter of global land 31 where fire is present, despite sub-optimal fire weather in the second fire season (Benali et al., 2017). In 32 summary, anthropogenic climate change, through a 0.9°C surface temperature increase since the pre-33 industrial period, has lengthened or increased the frequency of periods with heat and aridity that favour 34 wildfire on up to one-quarter of vegetated area, since 1979 (robust evidence, high agreement). 35

37 2.4.4.2.5 Observed changes in post-fire vegetation

Globally, fire has contributed to biome shifts (Section 2.4.3.2) and tree mortality (Section 2.4.4.2, 2.4.4.3) 38 attributed to anthropogenic climate change. Research since the IPCC Fifth Assessment Report has also found 39 vegetation changes from wildfire due to climate change. Through increased temperature and aridity, 40 anthropogenic climate change has driven post-fire changes in plant regeneration and species composition in 41 South Africa (Slingsby et al., 2017) and tree regeneration in the western United States (Davis et al., 2019b). 42 In the Fynbos vegetation of the Cape Floristic Region, South Africa, post-fire heat and drought and legacy 43 effects of exotic plant species reduced native plant species regeneration, decreasing species richness 12% 44 from 1966 to 2010 and shifting the average temperature tolerance of species upward by 0.5°C (Slingsby et 45 al., 2017). In burned areas across the western United States, the increasing heat and aridity of anthropogenic 46 climate change from 1979 to 2015 pushed low-elevation ponderosa pine (Pinus ponderosa) and Douglas-fir 47 (Pseudotsuga menziesii) forests across critical thresholds of heat and aridity that reduced post-fire tree 48 regeneration by half (Davis et al., 2019b). In the Southwestern United States of America, where 49 anthropogenic climate change has caused drought (Williams et al., 2019) and increased wildfire (Abatzoglou 50 and Williams, 2016), high-severity fires have converted some forest patches to shrublands (Barton and 51 Poulos, 2018). Field evidence shows that anthropogenic climate change and wildfire together have altered 52 vegetation species composition in the Southwestern USA and in the Cape floristic region, South Africa, 53 reducing post-fire natural regeneration and species richness of tree and other plant species, between 1966 54 and 2015, at global mean surface temperature increases of 0.3-0.9°C (medium evidence, high agreement). 55 56

[START FAQ2.3 HERE]

FAQ2.3: Is climate change increasing wildfire?

In the Amazon, Australia, North America, Siberia, and other regions, wildfires are burning wider areas than

6 *in the past. Analyses show that human-caused climate change has driven the increases in burned area in the* 7 *forests of western North America. Elsewhere, deforestation, fire suppression, agricultural burning, and*

short-term cycles like El Niño can exert a stronger influence than climate change. Many forests and

grasslands naturally require fire for ecosystem health but excessive wildfire can kill people, destroy homes,

10 and damage ecosystems.

11

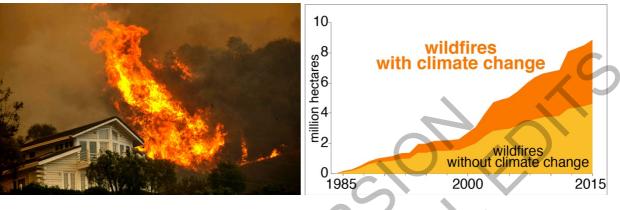


Figure FAQ2.3.1: (a) Springs Fire, May 2, 2013, Thousand Oaks, California, USA (photo by Michael Robinson Chávez, Los Angeles Times). (b) Cumulative area burned by wildfire in the western U.S., with (orange) and without (yellow) the increased heat and aridity of climate change (Abatzoglou and Williams, 2016).

Wildfire is a natural and essential part of many forest, woodland, and grassland ecosystems, killing pests,
 releasing plant seeds to sprout, thinning out small trees, and serving other functions essential for ecosystem
 health. Excessive wildfire, however, can kill people, cause breathing illnesses from the smoke, destroy
 homes (Figure FAQ2.1a), and damage ecosystems.

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Human-caused climate change increases wildfire by intensifying its principal driving factor – heat. The heat
of climate change dries out vegetation and accelerates burning. Non-climate factors also cause wildfires.
Agricultural companies, small farmers, and livestock herders in many tropical areas cut down forests and
intentionally set fires to clear fields and pastures. Cities, towns, and roads increase the number of fires that
people ignite. Governments in many countries suppress fires, even natural ones, producing unnatural
accumulations of fuel in the form of coarse woody debris and dense stands of small trees. The fuel
accumulations cause particularly severe fires that burn into tree crowns.

30 31 Evidence shows that human-caused climate change has driven increases in the area burned by wildfire in the forests of western North America. Across the western U.S., the higher temperatures of human-caused 32 climate change doubled burned area from 1984 to 2015, compared with what would have burned without 33 climate change (Figure FAQ2.1b). The additional area burned, 4.9 million hectares, is greater than the land 34 area of Switzerland. In this region, human-caused climate change has driven a drought from 2000 to 2020 35 that is the most severe since the 1500s, severely increasing the aridity of vegetation. In British Columbia, 36 Canada, the higher maximum temperatures of human-caused climate change increased burned area in 2017 37 to its widest extent in the 1950–2017 record, seven to eleven times the area that would have burned without 38 climate change. Moreover, in national parks and other protected areas of Canada and the U.S., climate 39 factors explained the majority of burned area from 1984 to 2014, with climate factors (temperature, rainfall, 40 aridity) outweighing local human factors (population density, roads, and urban area). 41 42

In other regions, wildfires are also burning wider areas and occurring more often. This is consistent with climate change but analyses have not yet shown if climate change is more important than other factors. In the Amazon, deforestation by companies, farmers, and herders who cut down and intentionally burn rainforests to expand agricultural fields and pastures causes wildfires even in relatively moister years. Drought exacerbates these fires. In Australia, much of the southeastern part of the continent has experienced extreme

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1	wildfire years, but analyses suggest th	hat El Niño, a heat phenomeno	on that cycles up and down periodically, is
2			ional burning of rainforests for oil palm
3			climate change. In Mediterranean Europe,
4	fire suppression seems to have preven		
5			some areas and contribute to major fires
6	•		ourning more often in permafrost areas
7			fluence of climate change. For the world
8			onverted from forest to farmland from
9	•		e evidence from the forests of western
10		used climate change has, on on	ne continent, clearly driven increases in
11	wildfire.		
12			
13	[END FAQ2.3 HERE]-		
14			
15			Co
16	2.4.4.3 Observed Changes in Tree	Mortality	
17			
18	2.4.4.3.1 Observed tree mortality glob		
19			ough increased aridity or drought (Section
20			ct pests (Section 2.4.4.3.3). Catastrophic
21	failure of the plant hydraulic system,		
22			eath (Anderegg et al., 2016; Adams et al.,
23	2017; Anderegg et al., 2018; Choat et	t al., 2018; Menezes-Silva et al	1., 2019; Brodribb et al., 2020).
24	Un through the IDCC Eigh A	nt Demost (Settels - + -1 - 2014)	determine and etterilentics and have been to d
25			, detection and attribution analyses had
26	found that anthropogenic climate cha		
27	industrial period and increases in arid	inty exceeding the effects of loc	cal non-climate change factors, caused

three cases of drought-induced tree mortality of up to 20% in the period 1945-2007, in western North 28 America (van Mantgem et al., 2009), the African Sahel (Gonzalez et al., 2012), and North Africa (le Polain 29 de Waroux and Lambin, 2012). Increased wildfire and pest infestations, driven by climate change, also 30 contributed to the North American tree mortality (van Mantgem et al., 2009). In addition, a meta-analysis of 31 published cases found that drought consistent with, but not formally attributed, to climate change, had 32 caused tree mortality at 88 sites in boreal, temperate and tropical ecosystems (Allen et al., 2010), with 49 33 additional cases found by the IPCC Fifth Assessment report (Settele et al., 2014). 34

- 35 Since the IPCC Fifth Assessment Report (Settele et al., 2014), global meta-analyses have found at least 15 36 (Allen et al., 2015) and 25 (Hartmann et al., 2018) additional sites of drought-induced tree mortality around 37 the world. These and other global analyses found more rapid mortality than previously (Allen et al., 2015), 38 rising background mortality (Allen et al., 2015), mortality increasing with drought severity (Greenwood et 39 al., 2017), mortality of tropical trees increasing with temperature (Locosselli et al., 2020), mortality 40 increasing with tree size for many species (Bennett et al., 2015), mortality predominantly at the dry edge of 41 species ranges (Anderegg et al., 2019a), and three-fourths of drought-induced mortality cases leading to a 42 change in the dominant species (Batllori et al., 2020). Multiple non-climate factors contribute to tree 43 mortality, including timber cutting, livestock grazing, and air pollution (Martinez-Vilalta and Lloret, 2016). 44 Globally, tropical dry forests lost, from all causes, 95,000 km², 8% of their total area, from 1982 to 2016, the 45 most extensive area of mortality of any biome (Song et al., 2018). 46
- 47 In summary, anthropogenic climate change has caused drought-induced tree mortality up to 20% in the 48 49 period 1945-2007 in western North America, the African Sahel, and North Africa, through global temperature increases of 0.3°-0.9°C above the pre-industrial period and increases in aridity, and contributed 50 to over 100 other cases of drought-induced tree mortality in Africa, Asia, Australia, Europe, and North and 51 South America (high confidence). Field observations document accelerating mortality rates, rising 52 background mortality, and post-mortality vegetation shifts (high confidence). Water stress, leading to plant 53 hydraulic failure, is the principal mechanism of drought-induced tree mortality. Timber cutting, agricultural 54 expansion, air pollution, and other non-climate factors also contribute to tree death. 55 56
- 57

2.4.4.3.2 Observed tree mortality in tropical ecosystems

1 In the Brazilian Amazon, deforestation to clear agricultural land comprises the principal cause of tree 2 mortality, reducing forest cover an average of 13,900 km² y⁻¹ from 1988 to 2020 (Espaciais, 2021). In 3 addition, an annual average temperature increase of 1.2°C from 1950 to 2018 (Marengo et al., 2018) 4 contributed to mortality in a set of 310 Amazon field plots of ~40% from 1983 to 2011 (Brienen et al., 5 2015). In another set of plots, mortality among newly recruited trees of mesic genera increased and drought-6 tolerant genera became more abundant from 1985 to 2015 (Esquivel-Muelbert et al., 2019). In other plots, 7 tree mortality did not show a statistically significant change from 1965 to 2016 but rose abruptly in severe 8 drought years, mainly during warm phases of the El Niño-Southern Oscillation (ENSO) (Aleixo et al., 2019). 9 Nearly half the area of the Amazon has experienced extremely dry conditions during ENSO warm phases, 10which can cause extensive wildfire (Section 2.4.4.2.3). Wildfire can increase tree mortality rates by >600% 11 above rates in non-burned areas, with the higher mortality persisting up to a decade after a fire (Silva et al., 12 2018; Berenguer et al., 2021). Climate change has contributed to tree mortality in Amazon rainforest 13 (medium evidence, medium agreement). 14 15 In the African Sahel field research has continued to detect tree mortality, ranging from 20% to 90% in the 16 period 1965-2018 (Kusserow, 2017; Trichon et al., 2018; Dendoncker et al., 2020), and declines in tree 17 biodiversity, with local losses of tree species up to 80% in the period 1970-2014 (Hanke et al., 2016; 18 Kusserow, 2017; Ibrahim et al., 2018; Dendoncker et al., 2020), consistent with, but not formally attributed 19 to climate change. In Algeria, mortality of Atlas cedar (Cedrus atlantica) increased from 1980 to 2006, 20

coinciding with a $\sim 1^{\circ}$ C spring temperature increase, but non-climate factors were not examined (Navarro-21 Cerrillo et al., 2019). Across southern Africa, nine of the 13 oldest known baobab trees (Adansonia digitata), 22 1100–2500 years old, have died since 2005, although the causes are unknown (Patrut et al., 2018). In South 23 Africa, savanna trees experienced an order of magnitude increase in mortality, related, but not formally 24 attributed to decreased rainfall (Case et al., 2019). In Tunisia, insect infestations related, but not formally 25 attributed to, hotter temperatures led to mortality of cork oaks (Quercus suber) (Bellahirech et al., 2019).

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2.4.4.3.3 Observed tree mortality in boreal and temperate ecosystems 28

The most extensive research on tree mortality since the IPCC Fifth Assessment Report has occurred in the 29 western United States, where anthropogenic climate change accounts for half the magnitude of a drought 30 from 2000 to 2020 that has been the most severe since the 1500s (Williams et al., 2020) and for one-tenth to 31 one-quarter of the magnitude of the 2012-2014 period of the severe 2012-2016 drought in California 32 (Williams et al., 2015a). Across the western United States, anthropogenic climate change doubled tree 33 mortality between 1955 and 2007 (van Mantgem et al., 2009). Lodgepole pine (Pinus contorta) mortality 34 increased 700% from 2000 to 2013 (Anderegg et al., 2015) and piñon pine (Pinus edulis) experienced over 35 50% mortality from 2002 to 2014 (Redmond et al., 2018). In California montane conifer forest, 36 anthropogenic climate change increased tree mortality one-quarter (Goulden and Bales, 2019). One-quarter 37 of trees died in some areas with mortality rates of ponderosa pine (Pinus ponderosa) and sugar pine (Pinus 38 *lambertiana*) increasing up to 700% of pre-drought rates (Stephenson et al., 2019; Stovall et al., 2019). 39 Substantial field evidence shows that anthropogenic climate change has caused extensive tree mortality in 40 North America (robust evidence, high agreement). 41

42

In western North America, increased infestations of bark beetles and other tree-feeding insects that benefit 43 from increased winter temperatures (IPCC AR6 WGI 3.3.1.1) and longer growing seasons (IPCC AR6 WGI 44 2.3.4.3.1) have killed drought-stressed trees (Section 2.4.2.1; Anderegg et al., 2015; Kolb et al., 2016; Lloret 45 and Kitzberger, 2018; Redmond et al., 2018; Stephens et al., 2018; Fettig et al., 2019; Restaino et al., 2019; 46 Stephenson et al., 2019). Increasing temperatures have allowed bark beetles to move further north and higher 47 in elevation, survive through the winter at sites where they would previously have died, and reproduce more 48 49 often (Raffa et al., 2008; Bentz et al., 2010; Jewett et al., 2011; Macfarlane et al., 2013; Raffa et al., 2013; Hart et al., 2017; Stephenson et al., 2019; Teshome et al., 2020; Koontz et al., 2021). Under warmer 50 conditions, some insects that were previously innocuous have become important agents of tree mortality 51 (Stephenson et al., 2019; Trugman et al., 2021). Field observations show mixed effects of bark beetle-52 induced tree mortality on subsequent fire-caused tree mortality (Andrus et al., 2016; Meigs et al., 2016; 53 Candau et al., 2018; Lucash et al., 2018; Talucci and Krawchuk, 2019; Wayman and Safford, 2021). From 54 1997 to 2018, ~5% of western U.S. forest area died from bark beetle infestations (Hicke et al., 2020). In 55 most circumstances, trees that have been weakened by drought are more vulnerable to being killed by bark 56 beetles (Anderegg et al., 2015; Kolb et al., 2016; Lloret and Kitzberger, 2018; Redmond et al., 2018; 57

1 2	Stephens et al., 2018; Fettig et al., 2019; Restaino et al., 2019; Stephenson et al., 2019; Koontz et al., 2021). Climate change has contributed to bark beetle infestations that have caused much of the tree mortality
3	in North America (<i>robust evidence, high agreement</i>) (Section 2.4.2.1).
4	
5	Across Europe, rates of tree mortality in field inventories from 2000 to 2012 were highest in Spain, Bulgaria,
6	Sweden, and Finland, positively correlated to maximum winter temperature and inversely correlated to
7	spring precipitation (Neumann et al., 2017). Tree mortality in Austria, the Czech Republic, Germany,
8	Poland, Slovakia, and Switzerland doubled from 1984 to 2016, correlated to intensified logging and
9	increased temperatures (Senf et al., 2018). Drought-related tree mortality rates from 1987 to 2016 were
10	highest in Ukraine, Moldova, southern France, and Spain (Senf et al., 2020). Climate contributed to tree
11	mortality across Europe from 1958 to 2001 (Seidl et al., 2011). In addition, insect infestations related to
12	higher temperatures (Okland et al., 2019) have caused extensive mortality of Norway spruce (Picea abies)
13	across nine European countries (Marini et al., 2017; Mezei et al., 2017). Across the Mediterranean Basin, a
14	combination of drought, wildfire, pest infestations, and livestock grazing has driven tree mortality (Penuelas
15	and Sardans, 2021). Climate change has contributed to tree mortality in Europe (<i>high agreement, medium</i>
16	<i>confidence</i>). (Section 2.4.2.1)
17	2.4.4.3.4 Tree mortality and fauna
18 19	A global meta-analysis of 631 cases of bird and mammal abundance changes in areas of tree mortality found
20	increasing abundance in a set of 186 bird species with increasing mortality and no trend in mammal
20	abundance (Fleming et al., 2021). Ground-nesting, ground foraging, tree hole nesting, and bark foraging
22	increased most, while nectar-feeding and foliage-gleaning birds declined. Invertebrates, especially ground-
23	foraging predators and detritivores, decreased.
24	
25	2.4.4.4 Observed Terrestrial Ecosystem Carbon
26	
27	2.4.4.1 Observed terrestrial ecosystem carbon globally
28	Terrestrial ecosystems contain stocks of 450 Gt (380–540 Gt) carbon in vegetation, 1700 Gt \pm 250 Gt carbon
29	in soils, and 1400 Gt \pm 200 Gt carbon in permafrost (Hugelius et al., 2014; Batjes, 2016; Jackson et al., 2017. Structure 2017, Et at al. 2018. We stal 2021) Events the statistic 2020, 4000
30	2017; Strauss et al., 2017; Erb et al., 2018a; Xu et al., 2021). Ecosystem carbon stocks, totaling 3030–4090 GtC (from lowest and highest estimates above) substantially exceed the ~900 Gt carbon in unextracted fossil
31 32	fuels (see Chapter 5 of WGI).
33	iders (see Chapter 5 of WOI).
34	Deforestation, draining of peatlands, expansion of agricultural fields, livestock pastures, and human
35	settlements, and other land use changes emitted carbon at a rate of 1.6 ± 0.7 Gt y ⁻¹ from 2010 to 2019,
36	(Friedlingstein et al., 2020), of which wildfires and peat burning emitted 0.4 ± 0.2 Gt y ⁻¹ from 1997 to 2016
37	(van der Werf et al., 2017). Anthropogenic climate change has caused a portion of these emissions through
38	increases in wildfire (Section 2.4.4.2.1) and tree mortality (Section 2.4.4.3.1) but the fraction of the total
39	remains unquantified. Land use change produced ~15% of global anthropogenic emissions, from fossil fuels
40	and land (Friedlingstein et al., 2020). Terrestrial ecosystems removed carbon from the atmosphere through
41	plant growth at a rate of -3.4 ± 0.9 Gt y ⁻¹ from 2010 to 2019 (Friedlingstein et al., 2020).
42	
43	Tropical deforestation and draining and burning of peatlands produce almost all of the carbon emissions
44	from land use change (Houghton and Nassikas, 2017; Friedlingstein et al., 2020), while forest growth

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- accounts for two-thirds of ecosystem carbon removals from the atmosphere (Pugh et al., 2019b). Global terrestrial ecosystems comprised a net sink of -1.9 ± 1.1 Gt y⁻¹ from 2010 to 2019 (Friedlingstein et al., 2020), mainly due to growth in forests (Harris et al., 2021; Xu et al., 2021), mitigating ~31% of global emissions from fossil-fuel burning and land use change (Friedlingstein et al., 2020).
- 49

In summary, terrestrial ecosystems contain 3000-4000 Gt carbon in vegetation, permafrost, and soils, three to five times the amount of carbon in unextracted fossil fuels, and 4.4 times the carbon currently in the atmosphere (*robust evidence, high agreement*). Tropical deforestation, draining and burning of peatlands and other land use changes emit 0.9-2.3 Gt y⁻¹ of carbon, ~15% of global emissions from fossil fuel and ecosystems (*robust evidence, high agreement*). Terrestrial ecosystems currently remove more carbon from

- the atmosphere, 2.5-4.3 Gt y^{-1} , than they emit, so tropical rainforests, Arctic permafrost, and other
- ⁵⁶ ecosystems provide the global ecosystem service of naturally preventing carbon from contributing to climate

57 change (*high confidence*).

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Chapter 2

2.4.4.4.2 Observed stocks in high-carbon terrestrial ecosystems

2 The ecosystem that attains the highest aboveground carbon density in the world is coast redwood (Sequoia 3 sempervirens) forest, in California, USA, with 2600 ± 100 t ha⁻¹ carbon (Van Pelt et al., 2016). The 4 ecosystem with the second highest documented carbon density in the world is mountain ash (Eucalyptus 5 regnans) forest in Victoria, Australia, with ~1900 t ha⁻¹ (Keith et al., 2009). Within the tropics, tropical 6 evergreen broadleaf forests (rainforests) in the Amazon, the Congo, and Indonesia attain the highest carbon 7 densities, reaching a maximum of 230 t ha-1 in the Amazon (Mitchard et al., 2014) and the Congo (Xu et al., 8 2017). Temperature increases reduce tropical rainforest aboveground carbon density 9.1 t ha⁻¹ per degree 9 Celsius, through reduced growth and increased mortality (Sullivan et al., 2020). 10

Tropical forests contain the largest vegetation carbon stock in the world, with 180-250 Gt of above- and 12 belowground carbon (Saatchi et al., 2011; Baccini et al., 2012; Avitabile et al., 2016). The Amazon contains 13 a carbon stock of 45–60 Gt (Baccini et al., 2012; Mitchard et al., 2014; Englund et al., 2017). 14

15

11

1

Ecosystems with high soil carbon densities include peat bogs in Ireland, with up to 3000 t ha⁻¹ (Tomlinson, 16

2005), Cuvette Centrale swamp forest peatlands in Congo, with an average of ~2200 t ha⁻¹ (Dargie et al., 17

2017), Arctic tundra, with an average of ~900 t ha⁻¹ (Tarnocai et al., 2009), and mangrove peatlands in 18Kalimantan, Indonesia, with an average of 850 ± 320 t ha⁻¹ (Murdiyarso et al., 2015). Arctic permafrost 19

- contains 1400 Gt \pm 200 Gt to 3 m depth, the largest soil carbon stock in the world (Hugelius et al., 2014). 20
- Globally, peatlands contain 470-620 Gt carbon (Page et al., 2011; Hodgkins et al., 2018) of which boreal 21
- and temperate peatlands contain 415 ± 150 Gt (Hugelius et al., 2020) and tropical peatlands 80–350 Gt (Page 22 et al., 2011; Dargie et al., 2017; Gumbricht et al., 2017; Ribeiro et al., 2021). Other analyses increase the 23
- upper estimates for boreal and temperate peatlands to 800-1200 Gt (Nichols and Peteet, 2019; Mishra et al., 24 25 2021b).
- 26

31

Tropical forests and Arctic permafrost contain the highest ecosystem carbon stocks in aboveground 27 vegetation and soil, respectively, in the world (robust evidence, high agreement). These ecosystems form 28 natural sinks that prevent the emission to the atmosphere of 1400-1800 Gt carbon that would otherwise 29 increase the magnitude of climate change (high confidence). 30

2.4.4.4.3 Biodiversity and observed terrestrial ecosystem carbon 32

High biodiversity and ecosystem carbon generally occur together, with rainforests in the Amazon, the 33 Congo, and Indonesia containing the largest aboveground vegetation carbon stocks (Saatchi et al., 2011; 34 Baccini et al., 2012; Avitabile et al., 2016) and the highest vascular plant species richness (Kreft and Jetz, 35 2007) in the world. Aboveground ecosystem carbon and animal species richness show high correlation but 36 also high spatial variability (Strassburg et al., 2010). Aboveground carbon is correlated to genus richness 37 globally (Cavanaugh et al., 2014), but to species richness only in local areas (Poorter et al., 2015; Sullivan et 38 al., 2017). Species richness generally increases vegetation productivity in the humid tropics while tree 39 abundance increases productivity in drier conditions (Madrigal-Gonzalez et al., 2020). Across the Amazon, 40 $\sim 1\%$ of tree species contain 50% of the aboveground carbon, due to abundance and maximum height (Fauset 41 et al., 2015). Aboveground carbon in tropical forest shows positive correlations to vertebrate species richness 42 (probability values not reported) (Deere et al., 2018; Di Marco et al., 2018). In logged and burned tropical 43 forest in Brazil, species richness of plants, birds, and beetles increased with carbon density up to ~100 t ha⁻¹ 44 (Ferreira et al., 2018). 45

46 National parks and other protected areas, which, in June 2021, covered 15.7% of global terrestrial area 47 (UNEP-WCMC, 2021), contain ~90 Gt carbon in vegetation and ~150 Gt carbon in soil (one-fifth and one-48 49 tenth, respectively, of global stocks) and remove carbon from the atmosphere at a rate of ~ 0.5 Gt y⁻¹ (onesixth of global removals) (Melillo et al., 2016). The most strictly protected areas contain carbon at higher 50 densities, but illegal deforestation and fires in some protected areas emit 38 ± 17 Mt y⁻¹ globally (Collins and 51 Mitchard, 2017). In the Amazon, protected areas store more than half of the aboveground vegetation carbon 52 stock of the region but account for only one-tenth of net emissions (Walker et al., 2020). Conservation of 53 high biodiversity areas, particularly in protected areas, protects ecosystem carbon, prevents emissions to the 54 atmosphere, and reduces the magnitude of climate change (high confidence). 55

2.4.4.4.4 Observed emissions and removals from high-carbon terrestrial ecosystems 1 Most global deforestation is occurring in tropical forests (Pan et al., 2011; Liu et al., 2015; Houghton and 2 Nassikas, 2017; Erb et al., 2018a; Li et al., 2018; Harris et al., 2021), primarily for clearing of agricultural 3 land (Hong et al., 2021), causing primary tropical forest to comprise a net source of carbon to the atmosphere 4 from 2001 to 2019 (emissions to the atmosphere 0.6 Gt y⁻¹, removals from the atmosphere -0.5 Gt y⁻¹, net 0.1 5 Gt y⁻¹) (Harris et al., 2021). While wildfires emitted an average of 0.4 ± 0.2 Gt y⁻¹ carbon from 1997 to 2016 6 (van der Werf et al., 2017), individual fire seasons can emit the same magnitude, such as the 0.4 Gt carbon 7 from the Amazon fires of 2007 (Aragao et al., 2018), 0.5 Gt carbon from the Amazon fires of 2015-2016 8 (Berenguer et al., 2021) and 0.2 Gt from the Australia fires of 2019-2020 (Shiraishi and Hirata, 2021). So, 9 wildfires account for up to one-third of annual average ecosystem carbon emissions, while major fire seasons 10 can emit up to two-thirds of global ecosystem carbon emissions (medium evidence, medium agreement). 11 12 Primary boreal and temperate forests also comprised net sources in the period 2001-2019, but, when 13 including all tree age classes, boreal, temperate, and tropical forests were net sinks, as growth exceeded 14 permanent forest cover losses (Harris et al., 2021), though boreal and temperate forests are much stronger 15 sinks (Pan et al., 2011; Liu et al., 2015; Houghton and Nassikas, 2017). Estimates of carbon removals from 16 remote sensing may provide more accurate estimates of boreal forest carbon balances than earth system 17 models, which overestimate regrowth after forest and timber (Wang et al., 2021a). Mortality of boreal forest 18in British Columbia from mountain pine beetle infestations converted 374 000 km² from a net carbon sink to 19 a net carbon source (Kurz et al., 2008). Modeling suggests that a potential increase in water-use efficiency 20 and regrowth could offset the losses in part of the forest mortality area (Giles-Hansen et al., 2021). 21 22 The Amazon as a whole was a net carbon emitter from 2003 to 2008 (Exbrayat and Williams, 2015; Yang et 23 al., 2018b), primarily due to expansion of agricultural and livestock areas, which caused over two-thirds of 24 deforestation from 1990 to 2005 (De Sy et al., 2015; De Sy et al., 2019). Four sites in the Amazon also 25 showed net carbon emissions from 2010 to 2018, from deforestation and fire (Gatti et al., 2021). In the 26 Amazon, deforestation emitted 0.17 ± 0.05 Gt y⁻¹ carbon from 2001 to 2015 (Silva Junior et al., 2020) while 27 fires emitted 0.12 ± 0.14 Gt y⁻¹ carbon from 2003 to 2015 (Aragao et al., 2018). An analysis of the Amazon 28 carbon loss from deforestation and degradation estimated a loss of 0.5 Gt y⁻¹ from 2010 to 2019, with 29 degradation accounting for three-fourths (Qin et al., 2021). Intact old-growth Amazon rainforest has been a 30 net carbon sink (Hubau et al., 2020) but may have become a net carbon source from 2010 to 2019 (Qin et al., 31 2021). 32 33 In Indonesia and Malaysia, draining and burning of peat swamp forests for oil palm plantations emitted 60 -34 260 Mt y⁻¹ carbon from 1990 to 2015, converting peatlands in that period from a carbon sink to a source 35 (Miettinen et al., 2017; Wijedasa et al., 2018; Cooper et al., 2020). Deforestation of mangrove forests 36 emitted 10-30% of deforestation emissions in Indonesia from 1980 to 2005 (Donato et al., 2011; Murdiyarso 37 et al., 2015), even though mangroves comprised only 3% of Indonesia primary forest area in 2000 (Margono 38 et al., 2014; Murdivarso et al., 2015). 39 40 In North America, wildfire emitted 0.1 ± 0.02 Gt y⁻¹ of carbon from 1990 to 2012, but regrowth was slightly 41 greater to produce a net sink (Chen et al., 2017). In California, USA, two-thirds of the 70 Mt carbon 42

emissions from natural ecosystems from 2001 to 2010 came from the 6% of the area that burned (Gonzalez
et al., 2015). Anthropogenic climate change caused up to half of the burned area (Section 2.4.4.2.1).

In the Arctic, anthropogenic climate change has thawed permafrost (Guo et al., 2020), leading to carbon emissions of 1.7 ± 0.8 Gt y⁻¹ in the winter from 2003 to 2017 (Natali et al., 2019). Wildfires in Arctic tundra in Alaska from ~1930 to 2010 caused up to 0.5 m of permafrost thaw (Brown et al., 2015), exposing peatland carbon (Brown et al., 2015; Gibson et al., 2018), including soil carbon deposits up to 1600 years old (Walker et al., 2019).

Tropical deforestation, draining and burning of peatlands, and thawing of Arctic permafrost due to climate change have caused those ecosystems to emit more carbon to the atmosphere than they naturally remove through vegetation growth (*high confidence*).

- 56 2.4.4.5 Observed Changes in Primary Productivity
- 57

- 2.4.4.5.1 Observed changes in terrestrial primary productivity
- 1 The difference between photosynthesis by plants (gross primary productivity [GPP]) and plant energy use 2 through respiration is the net growth of plants (net primary productivity [NPP]), which removes CO₂ from 3 the atmosphere and mitigates emissions from deforestation and other land use changes (Section 2.4.4.4). 4 Global terrestrial NPP has exceeded land use emissions since the early 2000s, making terrestrial ecosystems 5 a net carbon sink (Friedlingstein et al., 2020). 6
- 7

Global terrestrial NPP increased 6% from 1982 to 1999, through increased temperature and increased solar 8 radiation in the Amazon from decreased cloud cover (Nemani et al., 2003), then decreased 1% from 2000 to 9 2009, because of extensive droughts in the southern hemisphere (Zhao and Running, 2010). From 1999 to 10 2015, increased aridity caused extensive declines in the Normalized Difference Vegetation Index (NDVI) 11 globally, particularly semi-arid ecosystems (Huang et al., 2016), indicating widespread decreases in NPP 12 (Yuan et al., 2019). 13

14

Global terrestrial GPP increased 2% from 1951 to 2010 and continued increasing at least through 2016, with 15 increased atmospheric CO₂ showing a greater influence than natural factors (Li et al., 2017; Fernandez-16 Martinez et al., 2019; Liu et al., 2019a; Cai and Prentice, 2020; Melnikova and Sasai, 2020). Global forest 17 area increased 7% from 1982 to 2016, mainly from forest plantations and regrowth in boreal and temperate 18forests in Asia and Europe (Song et al., 2018), while regrowth in secondary forests > 20 years old, mainly in 19 boreal, temperate, and sub-tropical regions, generated a net removal of 7.7 Gt y⁻¹ CO₂ from the atmosphere 20 from 2001 to 2019 (Harris et al., 2021). Vegetation growth that exceeds the modelled CO₂ fertilisation, gaps 21 in field data, and incomplete knowledge of plant mortality and soil carbon responses introduce uncertainties 22 into quantifying the magnitude of CO₂ fertilisation (Walker et al., 2021). A combination of CO₂ fertilisation 23 of global vegetation and secondary forest regrowth has increased global vegetation productivity (medium 24 evidence, medium agreement). 25

26

The relative increase in GPP per unit of atmospheric CO₂ increase declined from 1982 to 2015, indicating a 27 weakening of any CO₂ fertilisation effect (Wang et al., 2020c). Increased growth from CO₂ fertilisation has 28 begun to shorten the life span of trees due to a trade-off between growth rate and longevity, based on 29 analyses of tree rings of 110 species around the world (Brienen et al., 2020). Furthermore, water availability 30 controls the magnitude of NPP (Beer et al., 2010; Jung et al., 2017; Yu et al., 2017), including water from 31 precipitation (Beer et al., 2010), soil moisture (Stocker et al., 2019), groundwater storage (Humphrey et al., 32 2018; Madani et al., 2020a), and atmospheric vapour (Novick et al., 2016; Madani et al., 2020b). Drought 33 stress reduced NPP across tropical forests from 2000 to 2015 (Zhang et al., 2019b) and GPP in the tropics 34 from 1982 to 2016 (Madani et al., 2020b). Drought stress has also reduced GPP in some semi-arid and arid 35 lands (Huang et al., 2016; Liu et al., 2019a). In addition, nitrogen and phosphorus constrain CO₂ fertilisation 36 (Terrer et al., 2019), though phosphorus limitation of tropical tree growth is species-specific (Alvarez-Clare 37 et al., 2013; Thompson et al., 2019). NPP has decreased during some time periods and in some regions 38 where drought stress has exerted a greater influence than increased atmospheric CO₂ (medium evidence, high 39 agreement). 40

41

Observed changes in freshwater ecosystem productivity 2.4.4.5.2 42

Temperature affects primary productivity through moderating phytoplankton growth rates, ice cover, thermal 43 stratification and growing season length (Rühland et al., 2015; Richardson et al., 2018). Global warming has 44 reinforced eutrophication, especially cyanobacteria blooms (Wagner and Adrian, 2009; Kosten et al., 2012; 45 O'Neil et al., 2012; De Senerpont Domis et al., 2013; Adrian et al., 2016; Visser et al., 2016; Huisman et al., 46 2018) (very high confidence). Conversely, warming can reduce cyanobacteria in hypertrophic lakes 47 (Richardson et al., 2019). Freshwater cyanobacteria may benefit directly from elevated CO₂ concentrations 48

(Visser et al., 2016; Ji et al., 2017; Huisman et al., 2018; Richardson et al., 2019). 49

50

Macrophyte growth in freshwaters is likely to increase with rising water temperatures, atmospheric CO₂ and 51 precipitation (robust evidence, high agreement) (Dhir, 2015; Hossain et al., 2016; Short et al., 2016; 52

Reitsema et al., 2018). Nonetheless, primary productivity in rivers is variable and unpredictable (Bernhardt 53

et al., 2018) because seasonal variations in temperature and light are uncorrelated, frequent high flow events 54

- reduce biomass of autotrophs and droughts can strand and desiccate autotrophs. 55
- 56

Chapter 2 FINAL DRAFT In large, nutrient-poor lakes warming-induced prolonged thermal stratification can reduce primary 1 production (medium evidence) (Kraemer et al., 2017). Warming may reduce phytoplankton concentrations 2 when temperature-induced increases in consumption of phytoplankton outpace increases in phytoplankton 3 production (De Senerpont Domis et al., 2013). These decreases in productivity may be under-recognised 4 responses to climate change. 5 6 Summary: Evidence is *robust* for increase in primary production with warming trends, but increases or 7

declines of algae cannot entirely be attributed to climate change; they are lake specific and modulated 8 through weather conditions, lake morphology, salinity, land-use and restoration, and biotic interactions 9 (medium evidence, medium confidence) (O'Beirne et al., 2017; Velthuis et al., 2017; Rusak, 2018; Ho et al., 10 2019). 11

12 13

2.4.5 **Conclusions on Observed Impacts**

14 The consistency of patterns of biological change with expectations from regional or global warming 15 processes, coupled with an understanding of underlying processes, the coherence of these patterns at both 16 regional and global scales, all form multiple lines of evidence (Parmesan et al., 2013) that it is very likely 17 that observed range shifts and phenological changes in individual species can be attributed to regional and 18global climate changes (very high confidence) (Section 2.4.2, Table 2.2; Table 2.3; Table SM2.1; Parmesan 19 et al., 2013). 20

21 Global and regional meta-analyses of diverse systems, habitats and taxonomic groupings document that 22 approximately half of all species with long-term records have shifted their ranges poleward and/or upward in 23 elevation and ~2/3 have advanced their timing of spring events (phenology) (Section 2.4.2, Table 2.2; 24 Parmesan and Hanley, 2015; Parmesan, 2019). Changes in abundance tend to match predictions from climate 25 warming, with warm-adapted species significantly out-performing cold-adapted species in warming habitats 26 (Feeley et al., 2020) and the composition of local communities becoming more 'thermophilised' i.e., 27 experiencing 'increase in relative abundance of heat-loving or heat-tolerant species' (Section 2.4.2.3; Cline et

- 28 al., 2013; Feeley et al., 2020). 29
- 30 New studies since AR5, with more sophisticated analyses designed to capture complex responses, indicate 31 that past estimates of the proportion of species impacted by recent climate change have been underestimates 32 due to their unspoken assumptions that local or regional warming should lead solely to poleward/upward 33 range shifts and advancements of spring timing (Duffy et al., 2019). More complex analyses have 34 documented cases of winter warming driving delayed spring timing of northern temperate species due to 35 chilling requirements, and increased precipitation driving species' range shifts downslope in elevation, and 36 eastward and westward in arid regions (high confidence). Further new studies have shown that phenological 37 changes have, in some cases, successfully compensated for local climate change and reduced degree of range 38 shifts (medium confidence). Limited number of studies of this type make it difficult to estimate the generality 39 of these effects globally (Section 2.4.2.5, Table 2.2). 40
- 41

Responses in freshwater species are consistent with responses in terrestrial species, including poleward and 42 upward ranges shifts, earlier timing of spring plankton development, earlier spawning in fish, and extension 43 of the growing season. Observed changes in freshwater species are strongly related to anthropogenic climate 44 change driven changes in the physical environment (e.g. increased water temperature, reduced ice cover, 45 reduced mixing in lakes, loss of oxygen, reduction in river connectivity). While evidence is high for an 46 increase in primary production in nutrient rich lakes with warming trends (high confidence), increasing or 47 declining algal formations are lake specific and modulated through variability in weather conditions, lake 48 49 morphology, changes in salinity, stoichiometry, land-use and restoration measures, and food web interactions. In boreal coniferous forest, there has been an increase in terrestrial derived dissolved organic 50 matter transported into rivers and lakes as a consequence of climate change (that has induced increases in 51 run-off and greening of the northern hemisphere), as well as to changes in forestry practices. This has caused 52 waters to become brown resulting in an acceleration of upper water warming and an overall cooling of deep 53 water (high confidence). Browning may accelerate primary production through input of nutrients associated 54 with DOM in nutrient poor lakes and increase cyanobacteria growth, which better cope with low light 55 intensities (medium confidence) (Sections 2.4.2.1, 2.4.2.2, 2.4.2.3, 2.4.2.4). 56 57

 declines in human health and wel (medium confidence). Clear attribution comparison of observed changes ecosystems. Table 2.3: Confidence in detecting a climate change. Lines of evidence for support stated confidence in attribution. 	button of these impacts is often not possible, but in biodiversity / ecosystem health and known is and attributing observed changes in terrestrial and for attribution of observed changes to climate change ion of key statements on observed biological change present lines of evidence. This is a summary table Experiment Image: Complex statistic analysis Models Complex statistic analysis Region Period Lines of evidence	services from those particula reshwater species and systems t e and increased CO2 are used to es to climate change and that is fully detailed in Table vations				
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declines in human health and wel (medium confidence). Clear attrib comparison of observed changes ecosystems.	in biodiversity / ecosystem health and known	services from those particula				
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declines in human health and wel	(medium confidence). Clear attribution of these impacts is often not possible, but inference can be made by					
	ll-being: e.g. tourism, recreation, food, liveliho					
	and project to continue, in biodiversity and eco					
America, western Asia, and East						
	agricultural expansion or abandonment, fire suppression, and inter-decadal cycles, strongly influence fire occurrence. Areas with the greatest increases in fire season length include the Amazon, western North					
	consistent with anthropogenic climate change.					
	a (high confidence). In parts of Africa, Asia, Au					
	m confidence). Increases in burnt area have bee					
	re, aridity, and drought have increased the leng					
	frost at numerous sites, dried some areas, and in <i>high confidence</i>) (Sections 2.4.4.4, 2.5.3.4).	increased fire, causing net				
A A	he carbon emissions from land use change. In t					
	(high confidence). Tropical deforestation, drai					
	and store globally critical stocks of carbon but					
-						
	ed through community interactions (high confid					
	rincipal mechanism of drought-induced tree mo					
1 0 0	d field evidence has detected tree mortality fro bical forests around the world (high confidence)					
	attributed increasing tree mortality at sites in A					
	creased wildfire and insect pests in many locat					
•	contributing to increased tree mortality directly	e ·				
	regime <i>(medium confidence)</i> (Section 2.4.3).					
(grasslands, arid regions and tundra) is <i>likely</i> being driven by climate change and increased CO ₂ in concert						
shifts in the deciduous/boreal for	change <i>(high confidence)</i> . These new studies help fill prior geographic and habitat gaps, for example documenting upward shifts in the forest/alpine tundra ecotone in the Andes, Tibet and Nepal, and northward					
documenting upward shifts in the shifts in the deciduous/boreal for		driven by climate change, and support prior studies that attributed such shifts to anthropogenic climate change (high confidence). These new studies help fill prior geographic and habitat gaps, for example				
change (<i>high confidence</i>). These documenting upward shifts in the shifts in the deciduous/boreal for	new studies help fill prior geographic and habi					
poleward and upslope, that are co driven by climate change, and su change (<i>high confidence</i>). These documenting upward shifts in the shifts in the deciduous/boreal for	new studies help fill prior geographic and habi	ed precipitation patterns anthropogenic climate				

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FINAL DRAFT	Cha	apter 2	IPCC WGII Six	th Assessment Report
About half of all species where land use change has been minimal have shifted their ranges, with 80–90% of movements being in the direction expected from regional warming trends – i.e. poleward and upward.	Global	Range 20 - 260 years		robust evidence high agreement very high confidence
Downslope elevational shifts and east-west shifts (shown for trees and birds) have been associated with regional increases in precipitation where precipitation has been shown to be the principal driver of a range boundary	USA	~ 40-60 years		limited evidence, high agreement, medium confidence
About two- thirds of all species with long- term (>20 years) records have shifted the timing of spring events in directions expected from regional winter and spring warming.	Global	Varies by study. Range: 20–400 years		robust evidence high agreement very high confidence
Winter chilling-depending species have delayed or not changed spring events despite spring warming countered by winter warming. When these species are taken into account, it is estimated that 92% of species in these studies have responded to regional warming trends	Northern Europe and USA	Varies by study. Range = 26–46 years		medium evidence high agreement high confidence
Anthropogenic climate change, acting through increased heat and aridity at global mean surface temperature increases of 0.6- 0.9°C, has increased the area burned by wildfire over natural levels, increasing burned area up to 11 times in one extreme year and doubling over natural levels in a 32-year period	western north America	1984–2017		robust evidence high agreement high confidence
Anthropogenic climate change has caused drought-induced tree mortality of up to 20% in three regions, through global mean surface temperature increases of 0.3-0.9°C above the pre-industrial period and increases in aridity, more than non-climate change factors	North America and Africa	ca. 1945–2007	P 200	medium evidence high agreement medium confidence
Anthropogenic climate change has caused latitudinal and elevational vegetation biome shifts in at least 19 sites in boreal, temperate, and tropical ecosystems, between 1700 and 2007, through local temperature increases of 0.4 to 1.6°C above the pre-industrial period more than non- climate change factors	Global	1500–2007		robust evidence high agreement high confidence
Anthropogenic climate change and wildfire together have altered vegetation species composition in at least two regions, reducing post-fire natural regeneration and species richness of tree and other plant species, at global mean surface temperature increases of 0.3-0.9°C	western North America, Africa	1966–2015		medium evidence high agreement medium confidence

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Beetles & moths shifting poleward and upward has brought new pest species into some forests; warming winters and longer growing season has increased destructive outbreaks of beetles and moths in temperate and boreal forests	North America, Europe	Varies by study	X & 	medium-high confidence
Exotic species are responding differently from native species in both abundance changes and phenological changes, but not in a consistent fashion	North America		~	low/medium evidence low agreement
The most cold-adapted species are generally declining in population abundances and contracting their ranges poleward and upward: (e.g. sea-ice dependent, mountain-top restricted, upper headwaters, coldest lakes)	Arctic, Himalayas , Antarctic, Alps		X &	medium confidence
Diseases of both wildlife and humans have emerged into new areas they have not been in historically	Global	past 20–100 years		medium confidence
Warming has amplified the trophic state lakes are already in. Eutrophic lakes have become more productive while oligotrophic lakes tend to become more nutrient limited	Global	Past 20–50 years	À 2 () 💱	robust evidence high agreement high/medium confidence
Woody encroachment into open (grassland, desert) systems has occurred, with climate change as one of the drivers, along with changes in grazing and other land uses	Global			medium confidence
In boreal, coniferous areas changes in forestry practices and climate change have caused an increase in terrestrial derived dissolved organic matter (DOM) transport into rivers and lakes leading to their browning	Boreal	Past decades	Constants	robust evidence high agreement high confidence
Climate change induced warming leads to shifts in thermal regime of lakes	Global	Past decades		robust evidence high agreement high confidence
Climate change causes gains and losses in freshwater water level	Global	Past decades		limited evidence low confidence
Greenhouse gas emissions from freshwater ecosystems are equivalent to around 20% of global burning fossil-fuel CO ₂ emission	Global	Past decades		medium evidence medium agreement medium/low confidence
In lakes weather extremes in wind, temperature, precipitation and loss of ice foremost affect the thermal regime with repercussions on water temperature, transparency, oxygen and nutrient dynamics, affecting ecosystem functionality	North America, Europe	Varies by study		medium/limited evidence high agreement medium/low confidence

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Climate change induced warming leads to shifts in thermal regime of rivers and streams; lowland rivers show a stronger thermal response than high-altitude, cold- water receiving streams	North America, Europe	Past decades	*	robust evidence medium agreement high confidence	
Loss of biodiversity in streams can be directly attributed to climate change through increased water temperatures, hydrological changes such as increased peak discharges, flow alteration and droughts	Global	Past decades		high agreement very high confidence	
Climate change is causing range shifts of freshwater fish	North America Europe	Past decades	A	high agreement, high confidence	

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2.5 Projected Impacts and Risk for Species, Communities, Biomes, Key Ecosystems and their Services

5 Under the Risk Assessment Framework that was introduced in AR5 (2014), risk means the probability of 6 harmful consequences resulting from climate change. It results from the interaction of vulnerability, 7 exposure, and hazard (see Chapter 1) and can be represented as the probability of occurrence of hazardous 8 events or trends multiplied by the impacts if these events or trends occur (IPCC, 2014). The Framework 9 defines vulnerability as a pre-existing condition, incorporating the extent to which species or ecosystems are 10 susceptible to, or unable to cope with, adverse effects of climate change. Vulnerable species have limited 11 adaptive capacity stemming from physiological and behavioural constraints, limited dispersal abilities and 12 restricted resource requirements or capacities for distributional and genetic changes (Foden et al., 2019) 13 (Foden et al., 2013; Cizauskas et al., 2017). Traits that render entire ecosystems vulnerable are harder to 14 define, but it is clear that vulnerabilities are high in the coldest habitats, in those with limited geographic 15 ranges such as low-lying islands, and in specialized, restricted habitats such as serpentine outcrops in 16 California (Anacker and Harrison, 2012) and dry meadows in Fennoscandia and Tibet (Yang et al., 2018a). 17 Ecosystem vulnerability can depend critically on the fates of plants that function as 'foundation species,' 18 providing community biomass aboveground and below, structuring habitat for fauna and providing 19 ecosystem services such as erosion control (Camac et al., 2021). 20

22 2.5.1 Projected Changes at Species and Community Levels

24 2.5.1.1 Assessment of Models and Sources of Uncertainties

Methods for projecting impacts of climate change on biodiversity can be classified into three types: 1) statistical models such as species distribution models (Elith and Leathwick, 2009); 2) mechanistic or process-based models (Chuine and Régnière, 2017) and 3) trait-based models (Pacifici et al., 2015). It is only recently that models have been developed looking at smaller levels of warming. such as 1.5°C (Hoegh-Guldberg et al., 2018a; Warren et al., 2018).

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- Species distribution models (SDMs) or niche-based models assess potential geographic areas of suitable 32 climate for the species in current conditions and then project them into future conditions (Trisurat, 2018; 33 Vieira et al., 2018). There are limitations in all models and it is critical that modellers understand the 34 35 assumptions, proper parameterization, and limitations of each model technique, including differences among climate models, emission scenarios or representative concentration pathways, and baselines (Araujo et al., 36 2019). Several systems automate development of SDMs, including R-packages (e.g., (Beaumont et al., 2016; 37 Hallgren et al., 2016), development of other model types (Foden et al., 2019) and aid in use of climate model 38 data (Suggitt et al., 2017), including allowing for connectivity constraints (Peterson et al., 2013). Buisson et 39 al. (2010) found most variation in model outputs stem from differences in design, followed by GCMs. 40
- 41
- Mechanistic approaches, also known as process-based models, project species' responses to climate changes
 by explicitly incorporating known biological processes, thresholds and interactions (Morin and Thuiller,

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1 2009; Maino et al., 2016). Mechanistic models are able to accommodate a broad range of climate change

impact mechanisms and include species-specific characteristics such as dispersal distances, longevity,
 fecundity, genetic evolution, phenotypic plasticity. However, sufficient knowledge is available for only a

fee undry, generic evolution, phenotypic plasticity. However, sufficient knowledge is available for only a
 few well-studied species. Species traits have been used to more broadly estimate potential climate change
 impacts (Foden et al., 2013; Cizauskas et al., 2017).

5 6

Most models are at large scales (20 km–50 km), and so cannot capture micro-climatic refugia generated by diversities of slope aspect, elevation or shade (Suggitt et al., 2015; Suggitt et al., 2018). In analyzing records of 430 climate-threatened and range-declining species in England, (Suggitt et al., 2015; Suggitt et al., 2018) showed that topographic diversity reduced population declines most strongly in areas experiencing the most local warming and in species most sensitive to warming. In these circumstances topographic diversity reduced population extinction risk by 22% for plants and by 9% for insects.

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None of the modelling techniques are predictions of the future; they are projections of possible futures. To date, few studies have validated model performance against observations, and many of these have been on islands, reducing ability to generalize (Fordham et al., 2018). Species' models should be considered as hypotheses of what a future world might look like if the climate projections came to pass. Suggestions have been made on how to start bringing more biotic interactions into species distribution models (Early and Keith, 2019), but limited basic ecological understandings of interactions, along with limits on computation and funding, constrain how far and how fast these modelling techniques can advance.

22 2.5.1.2 Risk Assessment and non-modelling approaches

23 In order to add realism and reliability to risk assessments at the species and community levels, non-24 modelling approaches based on known biological traits or processes, as well as on expert opinion (Camac et 25 al., 2021), are used to temper model outputs with ground-based validation. Trait-based assessment 26 approaches use species' biological characteristics as predictors of sensitivity, adaptive capacity and 27 extinction risk due to climate change. Climate exposure can be estimated using GIS-based modelling, 28 statistical programs or expert judgment (Chin et al., 2010). These trait-based approaches are widely applied 29 to predict responses of biodiversity to climate change because they do not require modelling expertise nor 30 detailed distibutional data (Pacifici et al., 2015; Willis et al., 2015). Most of these methods have not been 31 independently validated and do not allow direct comparison of vulnerability and risk among taxonomic 32 groups. 33

34 Some studies have combined two or three approaches for climate change risk assessment of biodiversity in 35 order to capture the advantages of each and avoid their limitations. Warren et al. (2013) used combinations 36 of SDMs and trait-based approaches to estimate the proportions of species losing their climatically suitable 37 ranges under the various future scenarios of climate and dispersal rate. Similarly, spatial projections of 38 climate change exposure were combined with traits to assess vulnerability of sub-Saharan amphibians 39 (Garcia et al., 2014). Laurance et al. (2012) combined 31 functional groups of species and 21 potential 40 drivers of environmental change to assess both the ecological integrity and threats for tropical protected 41 areas on a global scale. Keith et al. (2014) used a combination of three approaches (SDMs-trait-mechanistic) 42 to determine how long before extinction a species would become eligible for listing as threatened based on 43 the IUCN Red List criteria. 44

45 46 2.5.1.3 Risk of Species' Extinctions

48 2.5.1.3.1 Overview

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49 This assessment of current findings is of studies across a range of taxa and modelling techniques. Extinction risk estimates whether or not a particular species may be at risk of extinction over the coming decades if 50 climatic trends continue and usually does not take into account other human-induced stressors (e.g. invasive 51 species or pollution). It is not a prediction that a species will definitely go extinct, because even complete 52 loss of a species' range is projected, the scale of the model cannot estimate persistence in micro-climatic 53 refugia (Suggitt et al., 2015; Suggitt et al., 2018). Individuals and populations can survive after conditions 54 for successful reproduction are gone, leading to a lagged decline, called 'extinction debt' (Alexander et al., 55 2018). Therefore, range loss is an established criterion for assessing endangerment status and risk of 56 extinction. As a species range becomes smaller and occupied habitats become more isolated, the likelihood 57

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1 2 3	2 climatically suitable space and additional stressors (especially LULCC of critical habitat) the	
4		
5 6	5 The IUCN Red List Criteria (IUCN, 2019) classifies a species as 'critically endangered' if it 6 range loss of \geq 80%, with a resulting likelihood of extinction of \geq 50% in the near term (10-	100 years,
7 8 9	$\geq 50\%$, with a resulting likelihood of extinction of $\geq 20\%$ in the near term (10-100 years).	ange loss of
10		
11		area: (Lehner et
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15	risk, through changes in water temperatures, stream flow, loss of cold water habitat, increase	d variability of
16	precipitation, and increased disease risk from warming temperatures (high evidence, high ag	reement, high
17	<i>confidence</i>) (Knouft and Ficklin, 2017; Pletterbauer et al., 2018; Reid et al., 2019; Jaric et al	., 2019), adding
18	to stress from overexploitation and LULCC (Craig et al., 2017; IPBES, 2019 Global assessm	ient report).
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32		Tizu una Dzu)
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34		Woodward et al.,
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37	2019). In southern Europe, aquatic insects (Ephemeroptera, Plecoptera, and Trichoptera) are	endangered by
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41		onflies,
42	42 amphibians and turtles).	
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49 50		70 at mgn risk
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- changes, climate change, nutrient emissions, a global hydrological model and a global map of water bodies.
 It projects that changes in both water quality (eutrophication) and quantity (flow) will generate negative
- relations in freshwater ecosystems between persistence of species originally present in each community and a constellation of stressors, including harmful algal blooms. Under 4°C rise by 2050, mean abundance of

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species is projected to decline by 70% in running water and 80% in standing water (medium evidence, high agreement, medium confidence) (Janse et al., 2015 a).

2.5.1.3.3 Global projections of extinction risk 4

In prior reports, risk assessed from the literature was generally based on estimates of overall range 5 contractions with climate change. In AR4, extinction risk was carefully quantified: 'There is medium 6 confidence that approximately 20-30% of species assessed so far are likely to be at increased risk of 7 extinction if increases in global average warming exceed 1.5–2.5°C (relative to 1980 to 1999). As global 8 average temperature increase exceeds about 3.5°C, model projections suggest significant extinctions (40-9 70% of species assessed) around the globe.' These estimates approximately correspond to 50% reductions in 10 range size (IPCC, 2007a). AR5 stated 'a large fraction of terrestrial and freshwater species face increased 11 extinction risk under projected climate change during and beyond the 21st century, especially as climate 12 change interacts with other pressures ... (high confidence)' (Field et al., 2014). A series of multi-species and 13 global analyses have been published since AR5, using both statistical models and trait-based approaches. 14

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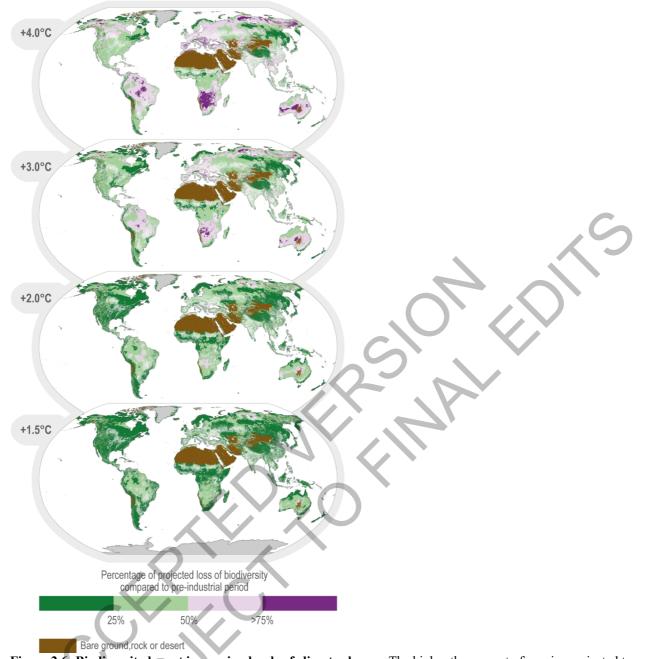
In this Chapter, risk to species, with implications for ecosystems, is assessed using three different 16

- approaches. First is an assessment of the geographic distributions of species' losses at different levels of 17
- GAST warming, termed 'biodiversity loss, measured as the proportion of species within a given location 18
- becoming classified as 'endangered' or 'critically endangered' (sensu IUCN). This measure provides estimates 19
- of which sites are at most risk of losing substantial numbers of species locally, leading to degradation of that 20
- ecosystems' ability to function. Second is an assessment of risk of proportions of species' becoming extinct 21
- globally at different levels of GAST warming, measured using the IUCN criteria for 'critically endangered', 22
- and termed 'species' extinction risk'. This measure is closest to assessing the complete loss of a species in the 23 wild, and can be used to compare to past (paleo) extinction rates. Third is an assessment of proportions of 24
- species becoming rare or endangered globally (not just locally), and is the foundation for the Burning 25
- Embers on biodiversity risk in Figure 2.11. These three approaches provide complementary information of 26
- the overall risks to biodiversity and ecosystem integrity under different warming levels. 27
- 28

Biodiversity risk, estimated as the proportion of species in a given area projected to become endangered 29

- (sensu IUCN), is projected to affect a greater number of regions with increasing warming, with about one-30
- third of land area risking loss of >50% of species currently inhabiting those ecosystems. Species' losses are 31
- projected to be worst in the northern South America, southern Africa, most of Australia, and northern high 32
- 33 latitudes (medium confidence) (Figure 2.6).
- 34

Biodiversity loss at increasing levels of climate change



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Figure 2.6: Biodiversity loss at increasing levels of climate change. The higher the percent of species projected to lose suitable climate in a given area, the higher the risk to ecosystem integrity, functioning and resilience to climate 3 change. Warming levels are based on global levels (GSAT) above pre-industrial temperatures. Colour shading represent 4 proportion of species for which the climate is projected to become unsuitable within a given pixel across their current 5 distributions at a given GSAT warming level, based on the data underpinning (Warren et al., 2018) (modelled 6 n=119,813 species globally, with no dispersal, averaged over 21 CMIP5 climate models). Areas shaded in green are 7 above the 50% biodiversity loss threshold, meaning that <50% of species in that area are projected to go locally extinct. 8 Areas shaded in pink and purple represent significant risk of biodiversity loss (areas where climates become unsuitable, 9 rendering them locally extinct, for >50% and >75% of species, respectively). The maps of species richness remaining 10 have been overlaid with a landcover layer (2015) from the European Space Agency Climate Change Initiative. This 11 landcover layer leaves habitats classified by the ESA as natural as being transparent, cities as black, water as blue, 12 permanent snow/ice as white and bare/rock as dark brown. Areas with a landcover identified as agriculture are 5% 13 transparent, such that potential species richness remaining if the land had not been converted to agricultural shows as 14 pale shading of the legend colours (very pale pink or very pale green). These paler areas represent biodiversity loss due 15 to habitat destruction, but with a potential to be restored, with green shading having potential for restoration to higher 16 species richness than pink and purple shadings. 17

- Risk of species' becoming extinct globally was estimated as the probability of loss of suitable climate space 1 rendering it critically endangered (sensu IUCN). It is *likely* that the percentage of species at high risk of 2 extinction (median and maximum estimates) will be 9% (max 14%) at 1.5°C, 10% (max 18%) at 2°C, 12% 3 (max 29%) at 3.0°C, 13% (max 39%) at 4°C and 15% (max 48%) at 5°C (Figure 2.7). Among groups 4 containing largest numbers of species at high risk of extinctions for mid-levels of projected warming (3.2°C) 5 are: invertebrates (15%), specifically pollinators (12%), amphibians (11%, but salamanders are at 24%) and 6 flowering plants (10%) (Figure 2.8a). All groups fare substantially better at 2°C, with extinction 7 projections reducing to <3% for all groups, except salamanders at 7% (medium confidence) (Figure 2.8a). 8 Even the lowest estimates of species' extinctions (9%) are 1000x natural background rates (Section 2.5.4; De 9 Vos et al., 2015). Projected species' extinctions at future global warming levels are in accord with 10 projections from AR4, assessed on much larger numbers of species with much greater geographic coverage 11
- and a broader range of climate models. (Figure 2.7; Figure 2.8)
- 13 14
- 70 quantiles 0.95 60 50 0.83 % extinction AR4 WG2 40 extinction verv risk assessment 30 likely* likely 20 0.50 10 0.17 0.05 0 0 4 5 6 GSAT increase above pre-industrial (°C)
- 15 Figure 2.7: Synthesis of modelled climate-driven extinction risk studies. The relationship between modelled 16 projections of extinction (expressed as a proportion of species at risk of extinction assessed in individual studies) and 17 GSAT increase above the pre-industrial average. Data (global sample size n = 178 modelled estimates) were sourced 18 from a number of sources, including digitization of data points in Figure 2 in the synthetic analysis of (Urban, 2015), 19 note: unweighted for sample size, n = 126; Table 4.1 of the AR4 WG2 chapter 2 (Fischlin et al., 2007), n = 40; 20 (Hannah et al., 2020) n = 6); and (Warren et al., 2018) n = 6). The Quantile regression (which is robust to the non-21 22 normal distribution of the response variable, and less sensitive to data outliers) was used to fit quantile estimates for levels relevant to informing "likely" (between the 0.17 and 0.83 quantiles, shaded in orange) and "very likely" ranges 23 (between the 0.05 and 0.95 quantiles, shaded in green) relating extinction risk to GSAT increase (Quantile regression 24 implemented using the Barrodale and Roberts algorithm in XLSTAT). The roughly equivalent estimate of this risk as 25 expressed in AR4 (Fischlin et al., 2007) is indicated by the dotted block indicating the medium confidence statement 26 "Approximately 20 to 30% of plant and animal species assessed so far (in an unbiased sample) are likely to be at 27 increasingly high risk of extinction as global mean temperatures exceed a warming of 2 to 3°C above preindustrial 28 levels (medium confidence)." This box is open on the right-side because AR4 estimates stipulated temperatures at or 29 exceeding given levels. 30

- 33 Projections of extinction risk by taxa are presented both for risk of becoming critically endangered (losing
- $\geq 80\%$ of suitable climate habitat, Figure 2.8a) and endangered (losing $\geq 50\%$ of suitable climate habitat,
- Figure 2.8b). The percentages of species projected at risk of becoming endangered (or worse) was 49% for
- insects, 44% for plants, and 26% for vertebrates at \sim 3°C global rise in temperature (Warren et al., 2018).
- Those estimates dropped considerably at lower levels of warming, down to 18%, 16%, and 8% at 2°C; and 6%, 8% and 4% at 1.5° C (Warren et al., 2018); thus not entirely dis-similar to the numbers in AR4 (Figure
- 2.7). Figure 2.8 shows the benefits of dispersal in offsetting extinction risk in birds, mammals, butterflies,

moths and dragonflies. While dispersal may benefit individual species, it poses additional risks to

communities and ecosystems as interactions between species are changed or eliminated.

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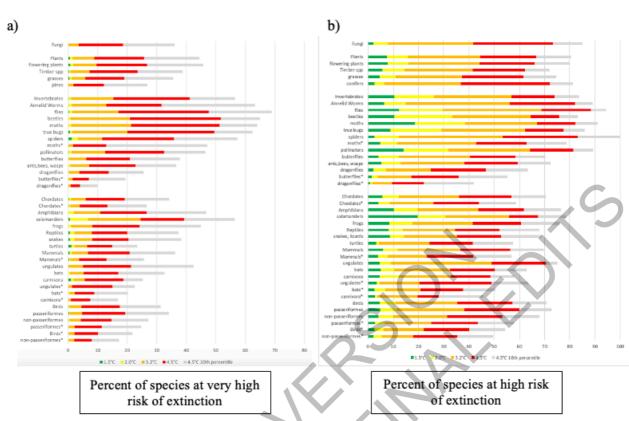


Figure 2.8: Percent of species of different groups classified as being under risk of extinction. a) Percent of the species 6 group listed projected to be at very high risk of extinction, corresponding to the IUCN Red List criteria for a species 7 classified as "critically endangered" (version 3.1) through losing >80% of its climatically suitable range area. b) Percent 8 of the species group listed projected to be at high risk of extinction, corresponding to the IUCN Red List criteria for a 9 species classified as "endangered" (version 3.1) through losing >50% of its climatically suitable range area. For a) and 10 b), values calculated from the underlying data underpinning (Warren et al., 2018). Values for each temperature are the 11 mean values across 21 CMIP5 models. The grey band represents the high-end of extinction risk from the 10th 12 percentile of the climate models to show the maximum range of values while the low end (90th percentile, 1.5°C) is not 13 shown as it is too small to appear on the plots. Taxa marked with * represent potential benefits from adaptation, 14 specifically dispersal at realistic rates (Warren et al., 2018); those with no * have dispersal rates that are essentially not 15 detected in the spatial resolution of the models (20 km). See (Warren et al., 2018) for caveats and more details. Sample 16 size for each group is as follows: Invertebrates (33949), Annelid Worms (155), Butterflies (1684), Moths (6910), 17 Dragonflies (599), Pollinators (1755), Spiders (2212), Beetles (7630), True Bugs (1728), Bees/Ants/Wasps (5914), 18 Flies (4809), Plants (72399), Flowering Plants (52310), Conifers (340), Timber spp (1328), Grasses (3389), Fungi 19 20 (16187), Vertebrates (12642), Mammals (1769), Carnivores (107), Ungulates (80), Bats (500), Birds (7968), Passeriformes (4744), Non-passeriformes (3224), Amphibians (1055), Frogs (887), Salamanders (163), Reptiles (1850), 21 Snakes (1741), Turtles (94). 22

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For local biodiversity loss, at 1.58°C (median estimate), >10% of species are projected to become
endangered (sensu IUCN); at 2.07°C (median) >20% of species are projected to become endangered,
representing high and very high biodiversity risk, respectively *(medium confidence)* (see Section 2.5.4;
Figure 2.11; Table 2.5, Table 2.S.4).

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Using data from geological time scales, Song et al. (2021) predicted that a warming of 5.2 °C above preindustrial would result in mass extinction comparable to that of the five mass extinctions over the past 540

- My, on the order of 70–85% of species going extinct, in the absence of non-climatic stressor. Mathes et al.
- 33 (2021) found evidence in the geological record that short-term rapid warming, on top of long-term warming
- trends, increases extinction risk by up to 40% over that expected from the long-term trend alone, with a

Most of the large-scale studies that have been performed are for losses based on climate alone (Figures 2.6, 2.7, 2.8). However, climate is rarely the only stressor affecting species survival. Habitat loss is currently the largest driver of range loss and extinction risk for most species (IUCN). Communities in different regions are becoming more similar to each other as species tolerant of human activities prosper and spread, with many rare and endemic species already having been driven extinct, primarily by LULCC (Pimm et al., 2006). Thus, it will likely be the interaction of climate change and habitat conversion (often also being driven by climate change) that will ultimately determine the risk and ability to survive of many species.

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10 2.5.1.4 Changing Risks of Diseases

11 Multiple studies predict increases in disease incidence or geographic and phenological changes of pathogens, 12 vectors, and reservoir host species due to climate change with or without other non-climatic variables 13 (González et al., 2010; Moo-Llanes et al., 2013; Roy-Dufresne et al., 2013; Liu-Helmersson et al., 2014; 14 Laporta et al., 2015; Ryan et al., 2015; Haydock et al., 2016; Hoover and Barker, 2016; Prist et al., 2017; 15 Blum and Hotez, 2018; Dumic and Severnini, 2018; Hundessa et al., 2018; Ryan et al., 2019; Ryan et al., 16 2021). However, models predicting changes in infectious disease risk are complex and sometimes produce 17 conflicting results and lack consensus (Caminade et al., 2014; Giesen et al., 2020). For example, malaria is 18 projected to increase in some regions of Africa, Asia, and South America by the end of the 21st century if 19 public health interventions are not sufficient, but malaria is also forecasted to decrease in some of the higher 20 risk areas (Peterson, 2009; Caminade et al., 2014; Ryan et al., 2015; Khormi and Kumar, 2016; Leedale et 21 al., 2016; Murdock et al., 2016; Endo and Eltahir, 2020; Mordecai et al., 2020). 22

While malaria risk is predicted to decrease in some lowland tropical areas as temperatures become too hot for vector or parasite development, other, warmer-adapted diseases like dengue and Zika, transmitted by *Aedes aegypti*, are predicted to increase (Ryan et al., 2019; Ryan et al., 2021). In more temperate regions, arboviruses and other vector-borne diseases with wider thermal breadths, such as West Nile fever, Ross River fever, and Lyme disease, are predicted to increase with climate warming (Ogden et al., 2008; Leighton et al., 2012; Shocket et al., 2018; Shocket et al., 2020; Couper et al., 2021), and drought can exacerbate these effects of temperature (Paull et al., 2017).

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A global analysis of 7346 wildlife populations and 2021 host-parasite combinations found that organisms 32 adapted to cool and mild climates are likely to experience increased risks of outbreaks with climate warming 33 while warm-adapted organisms may experience lower disease risk, providing further support for predictions 34 that climate change will increase infectious disease transmission in higher latitude regions across a 35 taxonomically diverse array of pathogens (robust evidence, high confidence) (Cohen et al., 2020). A study 36 examining the future risk of arboviruses (chikungunya, dengue, yellow fever, and Zika viruses) spread by 37 Aedes aegypti and Ae. albopictus projected increased disease risk due to interactions of multiple variables, 38 including increased human connectivity, urbanisation and climate change (Kraemer et al., 2019), although 39 vector species' ranges broaden only slightly (Campbell et al., 2015). 40

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In sum, climate change is expected to expand and redistribute the burden of vector-borne and other environmentally-transmitted diseases by shifting many regions toward the thermal optima of vector-borne disease transmission for multiple parasites, increasing transmission, while pushing temperatures above optima and toward upper thermal limits for other vectors and pathogens, decreasing transmission (Mordecai et al., 2019; Mordecai et al., 2020). These effects are mediated by other human impacts such as land use change, mobility, socio-economic conditions, and vector and pathogen control measures (Parham et al., 2015; Tjaden et al., 2018).

50 2.5.2 Projected Changes at Level of Biomes and Whole Ecosystems

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2.5.2.1 Global Overview, Assessment of Ecosystem-level Models, and Sources of Uncertainties

Shifts in terrestrial biome and changes in ecosystem processes in response to climate change are most frequently projected with dynamic global vegetation models (DGVMs), or land-surface models that form part of Earth System Models, which use gridded climate variables, atmospheric CO₂ concentration and information on soil properties as input variables. Since AR5, most of DGVMs have been upgraded to capture

Chapter 2 FINAL DRAFT IPCC WGII Sixth Assessment Report carbon-nitrogen cycle interactions (e.g. Le Quéré et al., 2018), many also include a representation of 1 wildfire, and fire-vegetation interactions (Rabin et al., 2017), and a small number now also accounts for land 2 management (such as wood removal from forests, crop fertilisation harvest of irrigation (Arneth et al., 2017). 3 Other forms of disturbance, such as tree mortality in response to, for example, episodic weather extremes or 4 insect pest outbreaks, are relatively poorly represented, or not at all, although they demonstrably impact 5 calculated carbon cycling (Pugh et al., 2019a). Simulated biome shifts are generally in agreement in 6 projecting broad patterns at the global scales, but vary greatly regarding the simulated trends in historical and 7 future carbon uptake or losses, both regionally and globally (WGI, Chapter 5 AR6; Chang et al., 2017). 8 9 Similar to other models, models to project large-scale changes in vegetation and ecosystem processes have to 10 deal with structural uncertainty (associated with the choice and the representation of processes in models), 11 input-data uncertainty (associated with variability in initial conditions and parameter values) and error 12 propagation (associated with coupling models) (Rounsevell et al., 2019). The IPBES methodological 13 assessment report on scenarios and models of biodiversity and ecosystem services provides a comprehensive 14 overview over the relevant issues (Ferrier et al., 2016). 15 16 In order to assess the model's performance, most models have been individually evaluated against a range of 17 observations. Moreover, in the annual updates of the global carbon budgets a model has to meet a small set 18of basis criteria to have its output included (Le Quéré et al., 2018). More systematic benchmarking 19 approaches have also been proposed that utilise a range of different data sets (Kelley et al., 2013; Chang et 20 al., 2017), in order to assess multiple simulated processes. These methods in principle allow to assign quality 21 scores to models based on their overall performance (Kelley et al., 2013). So far, this scoring does not yet 22 allow a clear quality ranking of models since the individual DGVMs tend to score well for some variables 23 and badly for others. A recent comparison of global fire-vegetation model outputs was also able to clearly 24 identify outliers when using a formalised benchmarking and scoring approach (Hantson et al., 2020). 25 However, benchmarking does not address sources of uncertainty and it would be advisable to perform 26 "perturbed-physics" experiments, in which multiple model parameters are varied in parallel more frequently, 27 as a means to test parameter-value uncertainty (Wramneby et al., 2008; Booth et al., 2012; Lienert and Joos, 28 2018). 29 30 Species diversity impact ecosystem functioning and hence ecosystem services (Hooper et al., 2012; Mokany 31 et al., 2016). So far, however, integrated modelling of ecosystem processes and biodiversity across multiple 32

Species diversity impact ecosystem functioning and hence ecosystem services (Hooper et al., 2012; Mokany et al., 2016). So far, however, integrated modelling of ecosystem processes and biodiversity across multiple trophic levels and food webs is in its infancy (Harfoot et al., 2014). Whether or not enhanced integration of state, function, and functional diversity across multiple trophic levels in models will markedly alter projections of how ecosystems respond to climate change thus remains an open research question.

Beyond simulating dynamically biome shifts and carbon cycling, which are important aspects of climate regulation, DGVMs can also provide information on a number of variables closely linked to other ecosystem services, such as water availability, air quality or food provisioning (Krause et al., 2017; Rabin et al., 2020). However, they are not intended to provide a comprehensive assessment of ecosystem services. For these, other approaches applied, but to date these are mostly applied on regional scales and are only weakly dynamic (Ferrier et al., 2016).

44 2.5.2.2 Projected Changes Globally at the Biome Level

45 Climate change and the associated change in atmospheric CO₂ levels already exacerbate other human-caused 46 impacts on structure and composition of land and freshwater ecosystems, such as land-use change, nitrogen 47 deposition, of pollution. The relative importance of these drivers for ecosystems over the coming decades 48 49 will likely differ between biomes, but climate change and atmospheric CO₂ will be pervasive unless we manage to rapidly limit fossil-fuel emissions and warming (high confidence) (Pereira et al., 2010; Warren, 50 2011; Ostberg et al., 2013; Davies-Barnard et al., 2015; Pecl et al., 2017; Ostberg et al., 2018). Global 51 vegetation and Earth system models agree on climate-change driven shifts of biome boundaries potentially 52 of hundreds of km over this century, combined with several substantial alterations that take place within 53 biomes (e.g., changes in phenology, canopy structure and functional diversity, etc.). Large discrepancies 54 exist between models and between scenarios regarding the region and the speed of change (Gonzalez et al., 55 2010b; Pereira et al., 2010; Pecl et al., 2017), but robust understanding is emerging in that the degree of 56 impact increases in high emission and warming scenarios (high confidence) (Figure 2.9). 57

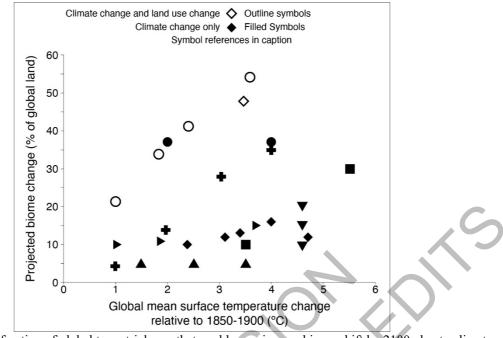


Figure 2.9: Projected fraction of global terrestrial area that could experience a biome shift by 2100, due to climate
change (filled symbols) or a combination of climate change and land use change (outline symbols), from publications in
Supplementary Table 2.S.3 (Projected vulnerabilities and risks of ecosystems to biome shifts). Circle filled (Bergengren
et al., 2011), square filled (Alo and Wang, 2008), diamond filled (Gonzalez et al., 2010b), triangle up filled (Scholze et al., 2006), triangle down filled (Sitch et al., 2008), triangle on side filled (Li et al., 2018), cross filled (Warszawski et al., 2013), circle outline (Ostberg et al., 2018), diamond outline (Eigenbrod et al., 2015).

10 11

Substantial changes in vegetation structure and ecosystem processes are already happening (see section 2.4). 12 Many of these observations have already been projected to take place as early as at least IPCC AR3 13 (Rosenzweig et al., 2007), and can they now be increasingly tested for their robustness with observational 14 evidence. These multiple changes in response to warming (and changes in precipitation and increasing 15 atmospheric CO₂ levels that go hand-in-hand with the warming) are further expected for already relatively 16 small additional temperature increases, in particular in cold (boreal, tundra) regions, as well as in dry regions 17 (high confidence): alterations of 2-47% of the areal extent of terrestrial ecosystems in scenarios of $<2^{\circ}$ C 18 warming above pre-industrial have been projected, increasing drastically with higher-warming scenarios 19 (Warren, 2011; Wårlind et al., 2014). More recent work, applying also probabilistic methods confirm the risk 20 of drastic changes in vegetation cover (e.g. forest to non-forest or vice versa) at the end of the 21st century 21 even for ca. 2°C warming scenarios, especially in tundra, and in tropical forest and savannah regions, with 22 more subtle changes (within a given biome types) likely to occur in all regions (Ostberg et al., 2013; Ostberg 23 et al., 2018). Model studies have found 5–20% of terrestrial ecosystems affected by warming ca. 2–3°C, 24 increasing to above one-third at a warming of 4-5°C (Ostberg et al., 2013; Warszawski et al., 2013). 25 26

In general, vegetation types are projected to be moving into their 'neighbouring' climates, depending on 27 whether temperature or precipitation is expected to be the predominant factor, and how vegetation interacts 28 with the increasing CO₂ levels in the atmosphere (Wårlind et al., 2014; Scheiter et al., 2015; Schimel et al., 29 2015; Huntzinger et al., 2017). For instance, boreal or temperate forest vegetation is simulated to migrate 30 polewards, closed tropical (moist) forest is expected to transition towards dry tropical forest types, while 31 climate-driven degradation might expand arid vegetation cover (Sections 2.5.4.2.1-2.5.4.2.5) However, 32 'novel ecosystems', that is, communities with no current or historical equivalent because of the novel 33 combinations of abiotic conditions under climate change, are expected to be increasingly common in the 34 future (medium confidence) although the regions where these novel ecosystems might emerge are still 35 disputed (Reu et al., 2014; Radeloff et al., 2015; Ordonez et al., 2016). The possibility of these novel 36 ecosystems and the communities that live within them poses challenges to current modelling of ecosystem 37 38 shifts, and will require new approaches to conservation that are designed to adapt to rapid changes in species composition and ensuing conservation challenges. 39

2.5.2.3 Risk to Arid Regions

3 Shifts in arid system structure and functioning that have been observed to date (section 2.4.3.3) are projected 4 to continue and include widespread woody plant encroachment, notably in savanna systems in Africa, 5 Australia and South America, and attributed to interacting land use change, climate change, and CO₂ 6 fertilisation effects (Fensholt et al., 2012a; Fang et al., 2017; Stevens et al., 2017). Arid Mongolian Steppe 7 grassland did not respond to experimentally elevated CO_2 (Song et al., 2019). Woody encroachment is 8 projected to continue or not reverse in North American drylands (Caracciolo et al., 2016), and in southern 9 African arid ecosystems (Moncrieff et al., 2014b). Dryland woody encroachment may increase carbon 10 stocks, depending on emissions scenario (Martens et al., 2021), but reduce soil water and biodiversity of 11 grassland-dependent species diversity (Archer et al., 2017). Warm season (C4) grass expansion into arid 12 shrublands risks sudden ecosystem transformation due to introduced wildfire (Bradley et al., 2016), a risk 13 anticipated for grass-invaded desert ecosystems of Australia and south-western United States (Horn and St. 14 Clair, 2017). Novel fire regimes in grassy shrublands have enhanced grass cover locally in southern African 15 Nama-Karoo (du Toit et al., 2015). 16

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Range retractions are projected for endemic plants in southern Africa (Young et al., 2016) and dry 18 woodlands in Morocco (Alba-Sánchez et al., 2015). Increasing thermal stress is projected to increase woody 19 plant mortality in Sonoran Desert ecosystems (Munson et al., 2016), and facilitate perennial grass 20 replacement by xeric shrubs in the south-western USA (Bestelmeyer et al., 2018). Ecological effects may 21 occur rapidly when extreme events compound long-term trends (Hoover et al., 2015), but evolve more 22 slowly as opportunity costs accumulate due to warming (Cross-Chapter Paper 3; Cunningham et al., 2021). 23

25 2.5.2.4 Risk to Mediterranean-type Systems

All Mediterranean Type Ecosystems (MTEs) show high confidence in projected increases in the intensity 27 and frequency of hot extremes and decreases in the intensity and frequency of cold extremes and medium 28 confidence in increasing ecological drought due to increased evapotranspiration (all regions) and reduced 29 rainfall (excluding California, where model agreement is low) (see Chapter 11 of WG1). Projections also 30 show a robust increase in the intensity and frequency of heavy precipitation in the event of 2° C warming or 31 more for MTEs in South Africa, the Mediterranean Basin and California, but are less clear for Australia and 32 Chile (see Chapter 11 of WG1). 33

34

MTEs are characterised by the distinctive seasonal timing of precipitation and temperature and disruption of 35 this regime is likely to be critical for their maintenance. Unfortunately, projections of changes in rainfall 36 seasonality have received less attention and are far more uncertain than many other aspects of climate 37 change (Pascale et al., 2016; Breinl et al., 2020), limiting our ability to predict the ecological consequences 38 of climate change in MTEs. Responses to experimental manipulation of rainfall seasonality show potential 39 for shifts in plant functional composition and diversity loss, but results vary with soil type (van Blerk et al., 40 2021). 41

42 Unfortunately, global and regional scale dynamic vegetation models show poor performance for large areas 43 of MTEs, because they do not characterise shrub and CAM-photosynthetic plant functional types well 44 (Moncrieff et al., 2015). Furthermore, the grain of these models are too coarse for quantifying impacts to 45 many vegetation formations which are patchy or of limited extent (e.g. forests). There is high confidence that 46 observations of high mortality in trees and other growth forms, reduced reproductive and recruitment 47 success, range shifts, community shifts towards more thermophilic species, and type conversions are set to 48 49 continue, either due to direct climate impacts through drought and other extreme weather events, or their interaction with factors like fire and pathogens (Sections 2.4.3.6; 2.4.4.2; 2.4.4.3; 2.5.4.2.4; 2.5.5.2). 50

51

Fire is a key driver across most MTEs due to summer-dry conditions. Climate projections for the MTEs 52 translate into high confidence that periods of low fuel moisture will become more severe and prolonged and 53 that episodes of extreme fire weather will become more frequent and severe (see Chapter 11, Section 8.3 54 WGI). This will lead to the birth of novel fire regimes in MTEs characterised by an increase in the 55 probability of greater burned area and extreme wildfire events (e.g. megafires), with associated loss of 56 57 human life and property, and long-term impacts on ecosystems and accelerating the possible loss of

1 2	resilience and capacity to recover (Abatzoglou and Williams, 2016; González et al., 2018; Boer et al., 2020; Moreira et al., 2020; Nolan et al., 2020; Duane et al., 2021; Gallagher et al., 2021).
3 4	Fire is virtually certain to have additional impacts through compound events (AR6 WGI Chapter 11.8).
4 5	Extreme postfire weather is extremely likely to continue to impact diversity (Slingsby et al., 2017), retard
6	vegetation regrowth (Slingsby et al., 2020a) and accelerate vegetation shifts (Batllori et al., 2019). Any
7	increases in the intensity and frequency of heavy precipitation are highly likely to compromise soil stability
8	in recently burnt areas (Morán-Ordóñez et al., 2020). The impacts of fire often depend on interactions with
9	non-climatic factors such as habitat fragmentation (Slingsby et al., 2020b), management (Steel et al., 2015)
10	or the spread of flammable exotic plantation forestry and invasive species (Kraaij et al., 2018; McWethy et
11	al., 2018). Managing these factors provides opportunities for adaptation and mitigation (Moreira et al.,
12	2020).
13	
14	Human adaptation and mitigation responses to climate change may create additional threats to MTEs. MTEs
15	have dry summers by definition, posing a challenge for the year-round supply of water to growing human
16	populations and agriculture. With recent major droughts in all MTEs (Section 2.4.3.6), there is increasing
17	reliance on groundwater for bulk water supply (Kaiser and Macleod, 2018). The majority of groundwater systems have exceeded or are rapidly approaching their environmental flow limits (de Graaf et al., 2019),
18 19	threatening human populations and ecosystems that depend on these systems for their persistence through
20	unfavourable climatic conditions (McLaughlin et al., 2017 Plants). Similarly, much of the MTEs are open
20	shrublands and grasslands and proposed extensive tree-planting to sequester atmospheric CO ₂ could result in
22	loss of biodiversity and threaten water security (Doblas-Miranda et al., 2017; Bond et al., 2019).
23	
24	2.5.2.5 Risk to Grasslands and Savannas
25	
26	Worldwide, woody cover is increasing in savannas (Buitenwerf et al., 2012; Donohue et al., 2013; Stevens et
27	al., 2017), as a result interactions of elevated CO_2 combined with altered fire and herbivory impacts (i.e.
28	from land-use change; see Section 2.4.3.5; CCP3.2; Venter et al., 2018; Wu et al., 2021). In some regions,
29	altered climate may also contribute (CCP3.2). Elevated CO ₂ benefits plants with C3 photosynthesis (often woody plants), more than C4 species (Moncrieff et al., 2014a; Scheiter et al., 2015; Knorr et al., 2016a).
30 31	Increases in woody vegetation in grassy ecosystems could provide some carbon increase (<i>medium</i>)
31	<i>confidence</i>) (Zhou et al., 2017; Mureva et al., 2018), but is expected to decrease biodiversity (Smit and Prins,
33	2015; Abreu et al., 2017; Andersen and Steidl, 2019), decrease water availability (Honda and Durigan, 2016;
34	Stafford et al., 2017) and alter ecosystem services like grazing and wood provision (<i>high confidence</i>)
35	(Anadon et al., 2014).
36	
37	The relative importance of climate, disturbance (e.g. fire/herbivory) and plant feedbacks in shaping present
38	and future savanna distribution vary between continents (Lehmann et al., 2014), which makes projections of
39	changing biome extent challenging (Moncrieff et al., 2016). It has been shown that simulation studies that do
40	not account of CO ₂ interactions and only consider climate change impacts do not realistically capture the
41	future distribution of savannas (<i>high confidence</i>) (Higgins and Scheiter, 2012; Moncrieff et al., 2016;
42	Scheiter et al., 2020). Due to the continued strong effect of CO_2 on tree (and shrub) to grass ratios in future, models suggest a loss of savanna extent and conversion into closed canopy forest/thicket and an expansion
43 44	of savanna-type vegetation into arid grasslands (Wårlind et al., 2014; Moncrieff et al., 2016). In arid
44 45	savannas and their interface to grasslands, survival of woody vegetation (which may be stimulated to grow
45 46	by increasing CO_2) will depend 1 on their capacity to survive potentially more severe and frequent droughts
47	(Sankaran and Staver, 2019). Across a range of models, for RCP4.5 future climate change and CO ₂
48	concentrations, savanna expanse declined by around 50% (converting to closed canopy systems) by 2070 in
49	Africa and South America, 25% in Asia with small changes in Australia (Moncrieff et al., 2016; Kumar et
50	al., 2021). Future fire spread is expected to be reduced with increased woody-dominance (Scheiter et al.,
51	2015; Knorr et al., 2016b; Scheiter et al., 2020), feeding back to further increase tree to grass ratios (high
52	confidence).
53	
	$-1 \cdot 1 + 4 \cdot$

Chapter 2

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Like the tropical forest biome, savannas are at large risk, given the projected climate changes in combination with land-use change (see Cross Chapter Paper 3). About 50% of Brazilian Cerrado has been transformed to agriculture and pastures (Lehman and Parr 2016) and African savannas have been proposed to follow a

agriculture and pastures (Lehman and Parr, 2016), and African savannas have been proposed to follow a
 similar tropical agricultural revolution pathway in order to enhance agronomical prosperity (Ryan et al.,

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7

Chapter 2

2016). In fact, indirect climate change impacts arising from mitigation efforts on land may be particularly perilous to savannas: extensive tree-planting to restore ecosystems and remove CO₂ from the atmosphere, as pledged, for example, under the African Forest Restoration Initiative, could lead to carbon losses, loss of biodiversity and damage ecosystem's water balance if trees are planted in what naturally are grasslands or savannas (Box 2.2, FAQ2.6; Bond et al., 2019).

2.5.2.6 Risk to Tropical Forest

8 Key factors affecting the future distribution of tropical humid and dry forests are amounts and seasonalities 9 of precipitation, increased temperatures, prolonged droughts and droughted-moderated fires (robust 10evidence, high agreement) (Bonai et al., 2016; Corlett, 2016; Lyra et al., 2017; Anderson et al., 2018; da 11 Silva et al., 2018; Fontes et al., 2018; O'Connell et al., 2018; Aguirre-Gutiérrez et al., 2019; Bartlett et al., 12 2019; Brando et al., 2019; Stan and Sanchez-Azofeifa, 2019). Probability of severe drought is projected to 13 quadruple in natural areas in Brazil with above 2°C warming (Barbosa and Lakshmi Kumar, 2016; Marengo 14 et al., 2020). Most multi-model studies assuming rapid economic growth/business-as-usual scenarios (A2, 15 A1B, RCP8.5) show an increase in future woody biomass and areas of woody cover towards the end of the 16 21st century in the temperate regions (Boit et al., 2016; Nabuurs et al., 2017) and in tropical forests in East 17 Africa (Ross et al., 2021) but decrease in the remaining tropical regions (Anadón et al., 2014; Boit et al., 182016; Lyra et al., 2017; Nabuurs et al., 2017; Maia et al., 2020). Terrestrial species are predicted to shift to 19 cooler temperatures and higher elevations (Pecl et al., 2017). Tropical species are more susceptible to 20 climate warming than temperate species (Rehm and Feeley, 2016; Sentinella et al., 2020). This susceptibility 21 will be exacerbated by road-building increasing ease of access into forests (Brinck et al., 2017; Taubert et al., 22 2018; Bovendorp et al., 2019; Senior et al., 2019). Furthermore, most tropical cloud forest species are unable 23 to invade grasslands and this will increase risk of extinctions in tropical cloud forests (Rehm and Feeley, 24 25 2015).

Sea level rise as the result of climate change is likely to influence mangroves in all regions, with greater
impact on North and Central America, Asia, Australia, and East Africa than West Africa and South America
(*robust evidence, high agreement*) (Alongi, 2015; Ward et al., 2016). On a small scale, mangroves are
potentially moving landward (Di Nitto et al., 2014), while on a large scale they will continue to expand
poleward (Alongi, 2015).

32 Most simulations predict a significant geographical shifts of transition areas between tropical forests and 33 savanna in the tropical and subtropical Americas and Himalayas (Anadón et al., 2014) (Rashid et al., 2015). 34 Forest die-back, as postulated for the Amazon region, does not occur in the majority of simulations (Malhi et 35 al., 2009; Poulter et al., 2010; Rammig et al., 2010; Higgins and Scheiter, 2012; Huntingford et al., 2013; 36 Davies-Barnard et al., 2015; Sakschewski et al., 2016; Wu et al., 2016a). Model projections of future 37 biodiversity in tropical forests are rare. Arguably, species are most vulnerable to climate change effects in 38 higher altitudes or at the dry end of tropical forest occurrence (medium evidence, medium agreement) 39 (Krupnick, 2013; Nobre et al., 2016; Trisurat, 2018). Tropical lowlands are expected to lose plant species as 40 temperatures rise above species' heat tolerance but could also generate novel communities of heat tolerant 41 species (robust evidence, high agreement) (Colwell et al., 2008; Trisurat et al., 2009; Trisurat et al., 2011; 42 Krupnick, 2013; Zomer et al., 2014a; Zomer et al., 2014b; Sullivan et al., 2020; Pomoim et al., 2021). 43 44

Statistical models that correlate data on species abundance with information on human pressures (such as 45 land-use change (Srichaichana et al., 2019), population density (Leclère et al., 2020) hunting (Mockrin et al., 46 2011) found for tropical and sub-tropical forests that birds, invertebrates, mammals and reptiles show a 47 decline in their probability of presence with declining forest cover, which is particularly pronounced in forest 48 49 specialists or narrow-ranged species (Newbold et al., 2014). Different soil fauna groups showed different responses in abundance and diversity to climate change conditions (Coyle et al., 2017; Facey et al., 2017) but 50 these changes can impact decomposition rates and biogeochemical cycles (medium evidence, low 51 agreement). 52

- 53
- ⁵⁴ Invasive plant species are predicted to expand upward by 500-1,500m in the Western Himalaya (Thapa et al.,
- 2018), and by 6-35% per year from the current extent in South America (*robust evidence, high agreement*)
 (Bhattarai and Cronin, 2014). Global assessment (Wang et al., 2017) also revealed that ecoregions of high
- (Bhattarai and Cronin, 2014). Global assessment (Wang et al., 2017) also revealed that ecoregions of high
 elevation tropical forests and sub-tropical coniferous forests have high risk of invasive plant expansion in the

low CO₂ emission scenarios, with negative impacts on ecosystem functioning and local livelihoods (Shrestha 1 et al., 2019). 2

3 The impact of unsustainable land use on tropical forests continues in all regions (see Cross-Chapter Paper 7). 4 Projected climate changes will not only cause impacts on biodiversity but also on the livelihoods of affected 5

people (robust evidence, high agreement). Increased drought drives crop failures that cause local 6

communities to expand agricultural area by further clearing native forests (Desbureaux and Damania, 2018). 7

Climate change is projected to enlarge the area of suitability for booming tree crops such as oil palm, acacia, 8

Eucalyptus, and rubber (Koninck et al., 2011; Cramb et al., 2015; Nath, 2016; Hurni et al., 2017; Li et al., 9

2017; Varkkey et al., 2018). An increase of 8% in area of rubber plantations in Yunnan province, China, 10

between 2002-2010 to higher altitude due to decreased environmental limits, potentially increases pressure 11

on remaining biodiversity both within and outside of protected areas (Zomer et al., 2014a). As a 12 consequence, the suitable area for mammals is projected to be reduced by 47.7% (RCP 2.6) and 67.7% 13

(RCP8.5) by 2070, with large variability depending on the different species (See also Cross-Chapter Paper 7; 14 Brodie, 2016).

15

16

2.5.2.7 Risk to Boreal and Temperate Forests 17

18 As in the Arctic, warming substantially exceeding the global average has already been observed for the 19 northern parts of the temperate and boreal forest zone (Gauthier et al., 2015), and is projected to continue 20 (see Chapter 4 of WGI, see Cross-Chapter Paper 6). As a consequence, boreal tree species are expected to 21 move northwards (or in mountain regions: upwards) into regions dominated by tundra, unless constrained by 22 edaphic features, and temperate species are projected to grow in regions currently occupied by southern 23 boreal forest (high confidence). In both biomes, deciduous trees are simulated to increasingly grow in 24 regions currently dominated by conifers (Wårlind et al., 2014; Boulanger et al., 2017). These simulation 25 results have been supported by observational examples. In Eastern Siberia, fire disturbance of larch-26 dominated forest was followed by recovery to birch-dominated forest (Stuenzi and Schaepman-Strub, 2020). 27 In Alberta Lodgepole Pine (Pinus contorta) lost its dominant status after attacks by Mountain Pine Beetles 28 (Dendroctonus ponderosae) caused the canopy to switch to non-pine conifers and broadleaved trees 29 (Axelson et al., 2018). In contrast to the examples above, some boreal forests have proven resilient to 30 disturbances, including to recent unprecedented spates of insect attacks (Campbell et al., 2019a; Prendin et 31 al., 2020). 32

33

Reforestation, either natural or anthropogenic, leads to summer cooling and winter warming of the ground, 34 while forest thinning or removal by fire has the reverse effects and deepens the upper layer free of 35 permafrost (Stuenzi et al., 2021a). Interactions between permafrost and vegetation are important. For 36 example, trees in East Siberian taiga obtained water mostly from rain in wet summers and mostly from 37 permafrost meltwater in dry summers (Sugimoto et al., 2002), suggesting that these forests will be 38 particularly vulnerable to combination of drought with retraction further underground of permafrost under 39 climate warming. 40

41

2.5.2.8 Risk to Peatland Systems

42 43

The overall effect of climate change on the extent of northern peatlands is still debated (limited evidence, low 44 agreement). It is expected that climate change will drive high-latitude peatland expansion poleward of their 45 present distribution due to warming, permafrost degradation, and glacier retreat, which could provide new 46 land and conditions favourable for peat development (limited evidence, medium agreement) (Zhang et al., 47 2017b), as seen during the last deglacial warming (robust evidence, high agreement) (MacDonald et al., 48 2006; Jones and Yu, 2010; Ratcliffe et al., 2018). Peatland area loss (shrinking) near the southern limit of 49 their current distribution or in areas where the climate becomes unsuitable is also expected (medium 50 evidence, medium agreement) (Section 2.3.4.3.2; Finkelstein and Cowling, 2011 temperature, and; Gallego-51 Sala and Colin Prentice, 2013; Schneider et al., 2016; Müller and Joos, 2020), though these peatlands could 52 persist if moisture is maintained via peatlands' self-regulating capacity. In Western Canada, a study suggests 53 that peatlands may persist until 2100, even though the climate will be less suitable (Schneider et al., 2016). 54 Simulations suggest that climate change driven increases in temperature and atmospheric CO₂ could drive 55 reductions in the northern peatland area up to 18% (SSP1-2.6), 41% (SSP2-4.5), and 61% (SSP5-8.5) by 56 2300 (Müller and Joos, 2020). This is in contrast with findings of northern peatland persistence and 57

1 2	expansion under RCP2.6 and RCP6.0 scenarios during 1861–2099 by another modelling study (Qiu et al., 2020). In the tropics, the only available study suggests peatland area will increase until 2300, mainly due to
3	the CO ₂ fertilisation effect (Müller and Joos, 2020).
4	
5	The combination of climate and land-use change represents a substantial risk to peatland carbon stocks, but
6	full assessment is impeded because peatlands are yet to be included in Earth System models (<i>limited</i>
7	evidence, high agreement) (Loisel et al., 2021). It is expected that the carbon balance of peatlands globally
8	will switch from sink to source in the near future (2020-2100), mainly because tropical peatland emissions,
9	together with those from climate change -driven permafrost thaw, will likely surpass the carbon gain
10	expected from climate change-driven enhanced plant productivity in the northern high latitudes (Gallego-
11	Sala et al., 2018; Chaudhary et al., 2020; Turetsky et al., 2020; Loisel et al., 2021), mainly because of
12	groundwater drawdown (robust evidence, medium agreement) (Hirano et al., 2014; Brouns et al., 2015; Cobb
13	et al., 2017; Itoh et al., 2017; Evans et al., 2021). The overall northern peatland carbon sink has been
14	simulated to persist for at least 300 years under RCP2.6, but not under RCP8.5 (Qiu et al., 2020).
15	
16	Increases in fire extent, severity, and duration are expected in all peatland regions in the future due to
17	temperature increases (Section 4.3.1.1), changes in precipitation patterns (section 4.3.1.2), and increases in
18	ignition sources (such as lightning) (Section 5.4.3.2), with associated rapid carbon losses to the atmosphere
19	(medium evidence, high agreement) (Dadap et al., 2019; Chen et al., 2021a; Nelson et al., 2021). For
20	example, drought has been linked to fires in SE Asian peatlands (Field et al., 2009) and there are predicted
21	decreases in mean summer precipitation (10-30%) for high and low RCPs, particularly over the Indonesian
22	region by mid and late twenty-first century (Section 12.4.2.2; Tangang et al., 2020; Taufik et al., 2020).
23	During wet years, the fire probability in Indonesian peatlands also significantly increases (+15-40 %) when
24	July-October temperatures surpass 0.5°C anomalies compared to a 1995-2015 baseline (Fernandes et al.,
25	2017). Overall, current evidence suggests that peat carbon losses via fire have the potential to be equal to, or
26	greater than, losses due to human peatland drainage and disturbance (limited evidence, high agreement)
27	(Turetsky et al., 2015).
<i>∠</i> /	(1 urciony of al., 2013).
27	
	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon
28 29 30	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon (<i>medium evidence, low agreement</i>) (Estop-Aragonés et al., 2018; Hugelius et al., 2020; Loisel et al., 2021;
28 29	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon (<i>medium evidence, low agreement</i>) (Estop-Aragonés et al., 2018; Hugelius et al., 2020; Loisel et al., 2021; Väliranta et al., 2021). In some permafrost peatlands, prolonged and warmer growing seasons due to climate
28 29 30	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon (<i>medium evidence, low agreement</i>) (Estop-Aragonés et al., 2018; Hugelius et al., 2020; Loisel et al., 2021; Väliranta et al., 2021). In some permafrost peatlands, prolonged and warmer growing seasons due to climate change (section 2.3.4.3.1), along with increases in nitrogen deposition since 1850 (Lamarque et al., 2013),
28 29 30 31 32 33	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon (<i>medium evidence, low agreement</i>) (Estop-Aragonés et al., 2018; Hugelius et al., 2020; Loisel et al., 2021; Väliranta et al., 2021). In some permafrost peatlands, prolonged and warmer growing seasons due to climate change (section 2.3.4.3.1), along with increases in nitrogen deposition since 1850 (Lamarque et al., 2013), are promoting plant primary productivity. Other studies indicate increased nitrogen-mediated sequestration
28 29 30 31 32	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon (<i>medium evidence, low agreement</i>) (Estop-Aragonés et al., 2018; Hugelius et al., 2020; Loisel et al., 2021; Väliranta et al., 2021). In some permafrost peatlands, prolonged and warmer growing seasons due to climate change (section 2.3.4.3.1), along with increases in nitrogen deposition since 1850 (Lamarque et al., 2013), are promoting plant primary productivity. Other studies indicate increased nitrogen-mediated sequestration could be exceeded by increased decomposition due to climate-change-driven warming and fire (<i>medium</i>)
28 29 30 31 32 33	In permafrost peatlands, studies differ, with some projecting net loss or and others net gain of carbon (<i>medium evidence, low agreement</i>) (Estop-Aragonés et al., 2018; Hugelius et al., 2020; Loisel et al., 2021; Väliranta et al., 2021). In some permafrost peatlands, prolonged and warmer growing seasons due to climate change (section 2.3.4.3.1), along with increases in nitrogen deposition since 1850 (Lamarque et al., 2013), are promoting plant primary productivity. Other studies indicate increased nitrogen-mediated sequestration could be exceeded by increased decomposition due to climate-change-driven warming and fire (<i>medium evidence, low agreement</i>) (Natali et al., 2012; Vonk et al., 2015; Keuper et al., 2017; Burd et al., 2018;
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FINAL DRAFT

2.5.2.9 Risk to Polar Tundra Ecosystems 53

54 For boreal-tundra systems, AR5 projected transformation of species composition, land cover and permafrost 55 extent, decreasing albedo and increasing greenhouse gases emission (medium confidence). The Special 56 Report on Global Warming of 1.5°C classified tundra and boreal forests as particularly vulnerable to 57

feedbacks to the climate system.

degradation and encroachment of woody shrubs (*high confidence*). The Special Report on Oceans and
 Cryosphere (SROCC) projected climate-related changes to arctic hydrology, wildfire and abrupt thaw, (*high confidence*) and broad disappearance of arctic near-surface permafrost in this century, with important
 consequences for global climate (*very high confidence*). Chapter 2 of AR6 has focused on new key findings
 about observed and projected changes in tundra vegetation and related hydrology, with implications for

6 7

Due to the rapid warming in high northern latitudes, Arctic tundra is one of the terrestrial biomes where 8 climate change impacts are already clearly visible (Settele et al., 2014; Uboni et al., 2016). Climate models 9 project that warming for the Arctic is likely to continue at more than double the global rate. Compared to the 10 period 1995–2014, mean annual surface air temperatures in Arctic tundra are projected to increase by 7.9°-11 10°C by the end of the century for scenarios of high greenhouse gas emissions (RCP 7.0 and 8.5). For 12 scenarios of low greenhouse gas emissions (RCP 1.9 and 2.6), the projected increase is 2.6°–3.2°C (see 13 Chapter 4 pf WGI). The Arctic is also projected to have among the largest increases in precipitation globally, 14 although there is high uncertainty in these projections. In contrast to climate change, land use change is 15 projected to be very low in Arctic tundra systems (van Asselen and Verburg, 2013). 16

17

Models of vegetation response to climate project acceleration in coming decades of observed increases in 18 shrub dominance and boreal forest encroachment that have been driven by recent warming (Settele et al., 19 2014), leading to a shrinking of the area of tundra globally (medium confidence) (Mod and Luoto, 2016; 20 Gang et al., 2017). Simulating changes in tundra vegetation is complicated by permafrost dynamics (e.g. 21 formation of thaw ponds), changes in precipitation, or low nutrient availability, which may promote 22 abundance of graminoids (van der Kolk et al., 2016). The changes in vegetation, when combined with 23 warming and increased precipitation effects on soil thawing and carbon cycling, are projected to modify 24 greenhouse gas emissions and have biophysical feedbacks to regional and global climate. Large uncertainty 25 in modelled carbon cycle changes arises from differences between the vegetation models (Nishina et al., 26 2015; Ito et al., 2016). In addition, climate change is expected to strongly interact with other factors, such as 27 fire, to further increase uncertainty in projections of tundra ecosystem function (Jiang et al., 2017). 28

29 30 2.5.2.10 Committed Impacts of Climate Change on Terrestrial Ecosystems and Implications of Overshoot

Projections point to potentially large changes of canopy structure and composition within and across the
terrestrial biomes in response to climate change and changes in atmospheric CO₂. These changes will
contribute to altered ecosystem carbon uptake and losses, biophysical climate feedbacks (Sections 2.3.2;
2.4.4; 2.5.3.2; 2.5.3.3, 2.5.3.4, 2.5.3.5, Figure 2.10, Table 2.4), and multiple other ecosystem services
(Sections 2.5.3, 2.5.4), as well as for biodiversity (Sections 2.4.2, 2.4.3, 2.4.4, 2.4.5, 2.5.1.3, 2.5.1.4, 2.5.2,
Figure Box 2.1.1, Table Box 2.1.1, Table SM2.4). Until now, most studies project changes over next decades
until the end of this century.

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However, there is an increasing body of literature that has found continued, longer-term responses of 40 ecosystems to climate change, so-called 'committed changes,' that arise from lags that exist in many systems. 41 Many processes in ecosystems take more than a few decades to quasi-equilibrate to environmental changes. 42 Therefore, trends of changing vegetation cover identified in simulations of transient warming continue to 43 show up in simulations that hold climate change at low levels of warming (medium confidence) (Boulton et 44 al., 2017; Pugh et al., 2018; Scheiter et al., 2020). Such changes, which could tip ecosystems into an 45 alternative state, could also be triggered by a 'warming overshoot' - if global warming were to exceed a 46 certain threshold, even if mean temperatures afterwards decline again (Albrich et al., 2020a). 47 48

49 For instance, even if warming achieved by 2100 remained constant after 2100, such committed responses

50 continue to occur. These include: (1) continued Amazon forest loss (Boulton et al., 2017), consistent with 51 results in (Pugh et al., 2018) that found continued tropical forest cover loss across a range of models and

results in (Pugh et al., 2018) that found continued tropical forest cover loss across a range of models and simulation set-ups, and (2) across Africa, an increased shift towards woody C3 vegetation was found in

simulation set-ups, and (2) across Africa, an increased shift towards woody C3 vegetation was found in equilibrium state, the overall response depending on the atmospheric CO₂ concentration (Scheiter et al.,

equilibrium state, the overall response depending on the atmospheric CO₂ concentration (Scheiter et al.,
 2020). In Pugh et al. (2018), the opposite was found for boreal forest cover, which showed a strong

- committed increase. The committed changes in vegetation composition correspond to large committed
- changes in terrestrial carbon uptake and losses (Boulton et al., 2017; Pugh et al., 2018; Scheiter et al., 2020),
- and would plausibly also appear in other ecosystem functioning and services. These studies point to the

Chapter 2

importance of having not only a multi-decadal but also a multi-century perspective when exploring the
 impacts of political decisions on climate change mitigation taken now. Even if climate-warming targets are
 met, published evidence so far suggests that fundamental changes in some ecosystems are *likely* as these
 correspond to well-understood ecosystem physiological responses that trigger long-term changes in
 composition.

[START BOX 2.1 HERE]

Box 2.1: Assessing Past Projections of Ecosystem Change Against Observations

To assess future climate change impacts on ecosystems we use models to project their future distribution. 12 Comparing the trends in the observed changes against the projections can help assess the strength of the 13 model projections. In this box, we compare observed trends of changes in ecosystem structure to projections 14 highlighted in previous IPCC reports (specifically AR3 (IPCC, 2001), AR4 (Fischlin et al., 2007) and AR5 15 (Settele et al., 2014). We use this to assess how well the projections are matching up with observed changes. 16 The map represents studies documenting observed changes in common plant functional groups (e.g. trees, 17 18 grasses, shrubs). Studies, documenting changes in plant functional groups, were collated from published papers in natural and semi-natural areas. Studies were included if climate change, or interactions between 19 climate change and land use showed a causal link to the observed change. Studies were excluded if the 20 changes only from landscape/land use transformation (e.g. deforestation). In each paper, we recorded the 21 geographical location, the type of functional change and noted the causes. Observed changes are plotted onto 22 a biome map derived from the WWF ecoregions database (Olson et al., 2001). Trends in changing plant 23 functional types are good indicators of potential biome shifts and are used to assess how observations match 24 up with projections. 25

26 27

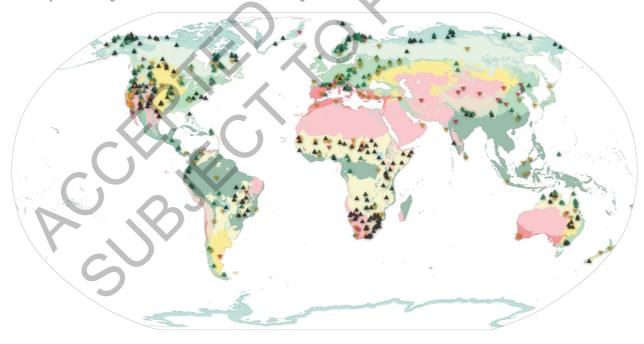
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Shifts in distribution of plant functional types

caused by climate change or combination of land use & climate change



Plant functional type changes

- ► Forest cover change
- Forest cover gain
- A Herbaceous cover gain
- ▲ Shrub / woodland cover gain
- Forest / woodland decline
 Herbaceous cover loss
- 28

Tropical grasslands / savannas / shrublands

Terrestrial biomes

Tropical broadleaf forests

Tropical coniferous forests

Temperate conifer forests

Boreal forests

Temperate broadleaf forests

Temperate grasslands / savannas / shrublands

Flooded grasslands

Montane grasslands

Mediterranean type ecosystems Deserts & xeric shrublands

Tundra

Figure Box 2.1.1: Observed changes in the distribution of plant functional types that are caused by climate change or combination of land use and climate change. Shifts in plant functional types are indicative of shift in biome function and structure

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Table Box 2.1.1: Comparison of projections on biome change from the Third, Fourth and Fifth assessment report (IPCC, 2001; Fischlin et al., 2007; Settele et al., 2014) with observed changes in ecosystems as assessed in this current report (see section 2.4 and Fig Box 2.1.1). Observed changes marked in bold show good agreement with past projections, those in red show mismatch with observations and projections.

Biome	AR3	AR4	AR5	Observed trends 1990–2021
Mediterranean Type ecosystems	Increased disturbance by fire and warming will cause a loss of unique habitats	Loss of 65% of area due to warming. Increased fire frequencies will favor resprouting plants. An increase in grass dominance. Forest expansion within MTE's systems due to elevated CO ₂ .	Range contractions of all species	Increased in water deficit and fire activity (sections 2.4.3.6, 2.4.4.2) causing declines in diversity, tree mortality (Fig Box 2.1.1) with resprouting trees worst affected. Increasing dominance of grasses (often alien). Increasing dominance of deciduous over evergreen species (Fig Box 2.1.1).
Tundra	Tree and shrub encroachment into tundra.	Increased woody plant growth due to longer and warmer growing seasons and replacement of dwarf tundra by shrub tundra Poleward expansion of tundra into polar desert and encroachment of coniferous trees into tundra	Continued woody expansion in tundra regions with reduced surface albedo due to less snow and more woody cover	Increase in woody shrub cover in tundra and expansion of boreal forest into tundra. (Fig Box 2.1.1, 2.4.3.4)
Boreal forest	Reduced productivity due to weather related disturbances (e.g. increase fire risk). Deciduous broadleaf tree encroachment into boreal forest	Extensive boreal tree spread into tundra. Boreal forest dieback within boreal zone and contraction of boreal forest at southern ecotone with continental grasslands		Expansion into Tundra and upslope tree line advance (Section 2.4.3.8 and Fig Box 2.1.1). Increased mortality due to drought, fire, beetle infestations (Sections 2.4.3.8, 2.4.42.1, 2.4.4.3.1)
Tropical forest	Increasing CO ₂ concentration would increase net primary productivity	Increases in forest productivity and biomass through increased CO ₂ with localised decreases in the Amazon. Shift in forest species composition. Expansion of forest area into mesic savanna.	Shift in the climate envelope of moist tropical forests but forests are less likely to undergo major retractions or expansions than suggested in AR4	Expansion of tropical forest into savannas in Africa, Asia, South America (Section 2.4.3.7, Fig Box 2.1.1) Forest biomass increases (though slowing). (Section 2.4.4.4) Forest degradation from drought, warming x fire and shorter residence time of trees. (Section 2.4.3.7) Shift in species composition towards species with more arid adapted trait (Section 2.4.3.7)
Temperate forest	Forest decline and increased mortality	Increase in tree mortality from drought related declines. A general increase of deciduous at the expense of evergreen vegetation is predicted at all latitudes		Map indicates a shift towards deciduous species in W N America (Fig Box 2.1.1) Tree death due to interactions with drought x pest outbreaks x fire (2.4.3.8, 2.4.42.1., 2.4.4.3.1)

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Grasslands and savannas	Increasing CO ₂ concentration will increase net primary productivity	Increased tree dominance in savannas and grasslands (from elevated CO ₂). With C3 plants benefiting more than C4 plants	Rising CO ₂ will increase the likelihood woodier states (but the transition will vary in different environments)	Greening and encroachment across tropical and temperate savannas in Africa, Asia, Australia and N America (Section 2.4.3.5) Expansion of trees into grasslands and advancement of tree lines Signs of increased C4 grass productivity in drought conditions. Increased C3 grass productivity (Section 2.4.3.5).
Desert/ arid shrublands		An increase in desert vegetation productivity was projected in southern Africa, the Sahel, central Australia, the Arabian Peninsula and parts of central Asia due to a positive impact of rising atmospheric CO ₂ .		Greening (increased LAI, woody cover) and increased herbaceous production is occurring at desert grassland interfaces (Chapter – Cross-Chapter Paper 3).

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Assessment: There is a high agreement between observations and projections of tree death in temperate and 3 boreal forests, with current projections (AR6) indicating this trend will continue (Section 2.5.4). Forest death 4 is most widely recorded in central Europe and Westerns North America (Fig Box 2.1.1). There is also very 5 high agreement between observations and projections of woody encroachment in savannas, grasslands and 6 tundra, with projections also indicating that this trend is likely to continue (Section 2.5.4). Observations of 7 desert greening show good agreement with earlier projections. Patterns of desertification are also occurring 8 although the geographical match between projections and observations shows moderate agreement, likely 9 due to the strong role of land use in this process. Projections of tropical forest expansion into mesic savannas 10 and boreal forest expansion into tundra also shows agreement with observations. 11

Projections on the future of Mediterranean shrublands, deserts, xeric shrublands and temperate grassy systems are limited making assessment of this relationship less clear. It is also unclear, due to limited observations, how widespread a shift from deciduous forest species to evergreen forest species is. Some observations suggest this is occurring although it is not clear how widespread this change is and if the geographical pattern is as projected.

[END BOX 2.1 HERE]

2.5.3 Risk Assessment of Ecosystems and Related Services

2.5.3.1 Risks in Protected Areas

National parks and other protected areas, which, in June 2021, covered 15.7% of global terrestrial area (UNEP-WCMC, 2021), conserve higher biodiversity than adjacent unprotected areas (Gray et al., 2016), and protect one-fifth of global vegetation carbon stocks and one-tenth of global soil carbon stocks (Section 2.4.4.4). This section assesses climate change specifically in protected areas. Even though it is in a part of the chapter on projected risks, this section includes both observed exposure and projected risks to gather the chapter information on protected areas in one place.

33 2.5.3.1.1 Observed exposure of protected areas

Deforestation, agricultural expansion, overgrazing, and urbanisation exposed to intense human pressure onethird of global protected area (6 million km²) in 2009, a 6% increase from 1993 (Venter et al., 2016; Jones et al., 2018). The observed change in exposure to climate change has not yet been quantified for protected areas globally but research has analysed spatial patterns and magnitudes of observed changes for the 360 000 km²

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1 2 3 4 5 6 7 8 9 10	1895 to 2010, mean annual temperatur century ⁻¹ , double the rate of the U.S. a area, compared with 4% for the U.S. a Arctic, at high elevations, and the aric weather station measurements in and temperature increased at a rate of 0.02	re of the U.S. national park as a whole, and precipitation as a whole, due to a high fra a Southwestern USA. (Gonz near six South African Nation $24 \pm 0.003^{\circ}$ C y ⁻¹ from 1960 for d area has been exposed to c	decreased on 12% of the national park ction of U.S. national park area in the alez et al., 2018). In addition, analyses of
11 12 13 14 15 16 17 18 19	from individual protected areas that conditions that are currently not prese protected area (Hoffmann et al., 2019 disappearance to 40% and the exposu the highest projected exposure in subt	3.5°C temperature increase loomprise half of global prote nt within an individual prot b). A lower emissions scena re to novel climates to 41% ropical projected areas (Hot litions from protected areas	by 2070, current climate could disappear cted area and novel climates (climate ected area) could expose half of global ario of ~1.5°C could reduce the climate (Hoffmann et al., 2019b). Models project ffmann and Beierkuhnlein, 2020). Projected is most extensive in Africa, Oceania, and
20 21 22 23 24 25 26 27 28 29 30 31 32 33	temperate conifer forests and grasslam change scenario of ~3.5°C temperatur tropical forests, 1.6 million km ² in 20 areas, while climate currently present million km ² , 12% of current total area combined could expose 2% of the hur analyses under RCP8.5 also project su Bolivia, Chile, and Peru (Fuentes-Cas Holsinger et al., 2019), China (Zomer	ds, and tundra (Hoffmann e e increase by 2100 could ex 00, to climate that would be in humid tropical forest pro by 2050 (Tabor et al., 2018 nid tropical forest protected ibstantial disappearance of e tillo et al., 2020), Canada, M et al., 2015), Europe (Nila l expose an extensive part o	novel to humid tropical forest protected tected areas could disappear from 0.6 8). High deforestation and climate change area (Tabor et al., 2018). Regional current climate from protected areas in Mexico, and the U.S. (Batllori et al., 2017; et al., 2019), and Indonesia (Scriven et al., f global protected area to disappearing and
 33 34 35 36 37 38 39 40 41 42 43 	a climate change scenario of 4°C temp trees (Adansonia perrieri, A. suarezer network (Vieilledent et al., 2013). Ott suitable climate from protected areas in Finland (Virkkala et al., 2013), bird woodlands in Germany (Steinacker et tropical dry forests in Mexico (Prieto-	berature increase by 2100, space asis) in Madagascar could short species and vegetation ty include Atlantic Forest amp ls and trees in Canada and M al., 2019), butterflies and n Torres et al., 2016). Project ase risks to the survival of sp	vpes at risk of partial disappearance of hibians in Brazil (Lemes et al., 2014), birds Mexico (Stralberg et al., 2020), bog nammals in Egypt (Leach et al., 2013), and ed disappearance of suitable climate pecies and vegetation types of conservation

Protected rivers, lakes, and other freshwater protected areas require inter-catchment connectivity to maintain species and population movements (Bush et al., 2014a; Hermoso et al., 2016; Thieme et al., 2016), but dams and other barriers interrupt connectivity (Grill et al., 2019). Climate change could also reduce freshwater connectivity (Section 2.3.3.3). Globally, over two-thirds of river reaches (by length) lack protected areas in their upstream catchments and nine-tenths of river reaches (by length) do not achieve full integrated protection (Abell et al., 2017).

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Terrestrial and freshwater protected areas can also serve as climate change refugia, locations where suitable conditions may persist for the species into the future (e.g. Section 2.6.5.6). In Canada, Mexico, and the U.S., only a fraction of protected area is located in potential climate change refugia under a 4°C temperature increase, estimated at 4% (Michalak et al., 2018) to 7% (Batllori et al., 2017). Potential refugia from biome shifts due to climate change under temperature increases of 1.8-3.4°C cover <1% of the U.S. national park area (Gonzalez et al., 2010b), a fraction that reduces to near zero when climate change is combined with

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1 2 3 4 5 6 7	habitat fragmentation due to land use change (Eigenbrod et al., 2015). Protected areas in boreal ecosystems could serve as refugia for species shifting north in Canada (Berteaux et al., 2018) and Finland (Lehikoinen et al., 2019). Invasive species, habitat loss, and other disturbances in protected areas could be lower than in unprotected areas across Europe (Gallardo et al., 2017) and specifically in Spain (Regos et al., 2016) and in Sri Lanka (Kariyawasam et al., 2020). Protected areas conserve refugia from climate change under a temperature increase of 4°C, important for biodiversity conservation but limited to <10% of current protected area (<i>medium confidence</i>).
8 9	2.5.3.2 Risks to Ecosystems and Services from Wildfire
10	2.5.5.2 Risks to Deosystems and Services from Haughe
11	2.5.3.2.1 Future projections of wildfire globally
12	Continued climate change under high emissions scenarios that increase global temperature ~4°C by 2100
13	could increase global burned area 50% (Knorr et al., 2016b) to 70% (Kloster and Lasslop, 2017) and global
14	mean fire frequency ~30% (Gonzalez et al., 2010b), with increases on one-third (Gonzalez et al., 2010b) to
15	two-thirds (Moritz et al., 2012) of global land and decreases on one-fifth (Gonzalez et al., 2010b; Moritz et
16	al., 2012). Lower emissions that would limit the global temperature increase to <2°C would reduce projected increases of global burned area to 30% (Lange et al., 2020) to 35% (Kloster and Lasslop, 2017) and
17 18	projected increases of fire frequency to ~20% (Gonzalez et al., 2010b; Huang et al., 2015). Continued
18	climate change could further lengthen fire weather seasons (IPCC AR6 WGI Chapter 12). Models that
20	combine projected climate change with potential agricultural expansion project decreases in total burned area
21	(Huang et al., 2015; Knorr et al., 2016b; Park et al., 2021). The area of projected increases in burned area
22	and fire frequency due solely to continued climate change is higher for the world as a whole than the area of
23	projected decreases (medium evidence, medium agreement).
24	
25	Increased wildfire due to continued climate change increases risks of tree mortality (Sections 2.5.2.6, 2.5.2.7,
26	2.5.3.2), biome shifts (Section 2.5.2.2), and carbon emissions (Sections 2.5.2.10, 2.5.3.4). Wildfire and
27	biome shifts under projected climate change of 4° C above the pre-industrial period, combined with international trade and transport, cause high risks of invasive species across one-sixth of global area,
28 29	including extensive high-biodiversity regions (Early et al., 2016).
30	including extensive high blodiversity regions (Eurif) et ul., 2010).
31	Wildfire risks to people include death and destruction of homes, respiratory illnesses from smoke (Ford et
32	al., 2018; Machado-Silva et al., 2020), post-fire flooding from areas exposed by vegetation loss, and
33	degraded water quality through increases in sediment flows (Dahm et al., 2015) and chemical precursors of
34	carcinogenic trihalomethanes when water is later chlorinated for drinking (Section 2.5.3.7; Uzun et al., 2020)
35	. Under RCP8.5 and shared socio-economic pathway SSP3 (high population growth, slow urbanisation), the
36	number of people living in fire-prone areas could increase by three-quarters, to 720 million people in 2100,
37	in a projected global population of 12.4 billion people (Knorr et al., 2016b). Lower emissions under RCP4.5
38	could reduce the number of people at risk by 70 million people. In these projections, human population
39 40	growth increases human exposure to wildfires more than increase in burned area (Knorr et al., 2016c). A global temperature increase <2°C could increase global population exposure to wildfire by ~30% (Lange et
40 41	al., 2020). Increased wildfire under continued climate change increases probabilities of human exposure to
42	fire and risks to public health (<i>medium evidence</i> , <i>high agreement</i>).
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44 2.5.3.2.2 Future projections of wildfire in high-risk areas

Regions identified at high risk of increased burned area, fire frequency, or fire weather by multiple global 45 analyses include: Amazon (Gonzalez et al., 2010b; Huang et al., 2015; Knorr et al., 2016c; Burton et al., 46 2018; Abatzoglou et al., 2019), Mediterranean Europe (Gonzalez et al., 2010b; Burton et al., 2018; 47 Abatzoglou et al., 2019), Arctic tundra (Moritz et al., 2012; Flannigan et al., 2013), western Australia 48 (Gonzalez et al., 2010b; Burton et al., 2018; Abatzoglou et al., 2019), western United States (Gonzalez et al., 49 2010b; Moritz et al., 2012; Knorr et al., 2016c). Higher-resolution spatial projections indicate high risks of 50 increased wildfire in the Amazon, Australia, boreal ecosystems, Mediterranean Europe, and the United 51 States under climate change (medium evidence, medium agreement). 52

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In the Amazon, climate change under RCP8.5, combined with high deforestation, could double the area of high fire probability (Fonseca et al., 2019), double burned area by 2050 (Brando et al., 2020) increase burned area 400–2800% by 2100 (Le Page et al., 2017), and increase fire intensity 90% (De Faria et al., 2017).

Lower greenhouse gas emissions (RCP4.5) and reduced deforestation could reduce fire risk to a one-fifth

Chapter 2 FINAL DRAFT IPCC WGII Sixth Assessment Report increase in the area of high fire probability (Fonseca et al., 2019) and a 100-500% increase in burned area by 1 2100 (Le Page et al., 2017). Moreover, increased fire, deforestation, and drought, acting through vegetation-2 atmosphere feedbacks, increase risks of extensive forest dieback and potential biome shifts of up to half of 3 Amazon rainforest to grassland, a tipping point that could release an amount of carbon that would 4 substantially increase global emissions (Oyama and Nobre, 2003; Sampaio et al., 2007; Lenton et al., 2008; 5 Nepstad et al., 2008; Malhi et al., 2009; Settele et al., 2014; Lyra et al., 2016; Zemp et al., 2017a; Zemp et 6 al., 2017b; Brando et al., 2020). Continued climate change, combined with deforestation, increases risks of 7 wildfire and extensive forest dieback in the Amazon rainforest (robust evidence, high agreement). 8 9 In Australia, climate change under RCP8.5 increases risks of pyroconvective fire by 20 to 40 days in 10 rangelands of Western Australia, South Australia, and the Northern Territory (Dowdy et al., 2019). 11 Pyroconvective fire conditions could reach more frequently into the more populated areas of New South 12 Wales, particularly at the start of austral summer (Di Virgilio et al., 2019). General circulation models do not 13 agree, however, on projected areas of fire increase in New South Wales (Clarke and Evans, 2019). Increases 14 in heat and potential increases in wildfire threaten the existence of temperature montane rainforest in 15 Tasmania, Australia (Mariani et al., 2019). 16 17 In Mediterranean Europe, climate change of 3°C could double or triple burned area, while keeping the 18 temperature increase to 1.5°C could limit burned area increase to 40-50% (Turco et al., 2018). Under 19 RCP8.5, the frequency of heat-induced fire weather could increase 30% (Ruffault et al., 2020). Severe fire 20 followed by drought could cause biome shifts of forest to non-forest (Batllori et al., 2019) and tree mortality 21 >50% (Dupire et al., 2019). 22 23 In Arctic tundra, boreal forests, northern peatlands, including permafrost areas, climate change under 24 scenarios of 4°C temperature increase could triple burned area in Canada (Boulanger et al., 2014), double the 25 number of fires in Finland (Lehtonen et al., 2016), increase lightning-driven burned area 30 to 250% 26 (Veraverbeke et al., 2017; Chen et al., 2021a), push half of the area of tundra and boreal forest in Alaska 27 above the burning threshold temperature, and double burned area in Alaska (Young et al., 2017a). Thawing 28 of Arctic permafrost from a projected temperature of 4°C and resulting wildfire could release 11-200 Gt 29 carbon that could substantially exacerbate climate change (Section 2.5.2.9). 30 31 In the United States, climate change under RCP8.5 could increase burned area 60-80% by 2049 (Buotte et 32 al., 2019) and the number of fires with an area $>50 \text{ km}^2$ by 300–400% by 2070 (Barbero et al., 2015). In 33 montane forests in the U.S., climate change under RCP8.5 increases the risk of fire-facilitated conversion of 34 \sim 7% of forest to non-forest by 2050 (Parks et al., 2019). In California, climate change under a scenario of 35 4°C temperature increase could double fire frequency in some areas (Mann et al., 2016), but emissions 36 reductions that limit the temperature increase to ~2°C could keep fire frequency from increasing (Westerling 37 et al., 2011). Carbon dioxide fertilisation and increased temperature under climate change could increase 38 invasive grasses and wildfire in desert ecosystems of the Southwestern United States, where wildfire has 39 historically been absent or infrequent, and increase mortality of the sparse tree cover (Horn and St. Clair, 40 2017; Klinger and Brooks, 2017; Syphard et al., 2017; Moloney et al., 2019; Sweet et al., 2019). 41 42 In summary, under a high emissions scenario that increases global temperature 4°C by 2100, climate change 43 could increase global burned area 50-70% and global mean fire frequency $\sim 30\%$ with increases on one-third 44 to two-thirds of global land and decreases on one-fifth of global land (medium confidence). Lower emissions 45 that would limit the global temperature increase to $<2^{\circ}$ C would reduce projected increases of burned area to 46 ~35% and projected increases of fire frequency to ~20% (medium confidence). Increased wildfire, combined 47 with erosion due to deforestation, could degrade water supplies (high confidence). For ecosystems with 48 historically low fire frequencies, a projected 4°C global temperature increases risks of fire, contributing to 49 potential tree mortality and conversion of over half of Amazon rainforest to grassland and thawing of Arctic 50 permafrost that could release 11-200 Gt carbon that could substantially exacerbate climate change (medium 51 confidence). 52 53 2.5.3.3 Risks to Ecosystems and Services from Tree Mortality 54

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8 9 Under continued climate change, increased temperature, aridity, drought, wildfire (Section 2.5.3.2), and insect infestations (Section 2.4.4.3.3) will tend to increase tree mortality across wide parts of the world (McDowell et al., 2020). Boreal and temperate forest loss to fire, wind, and bark beetles could cause more negative than positive effects for most ecosystem services, including carbon storage to regulate climate change (Sections 2.4.4.3, 2.5.2.6, 2.5.2.7, 2.5.3.4), water supply for people (Section 2.5.3.6.1), timber production (Chapter 5), and hazard protection (Thom and Seidl, 2016). In addition, deforestation in tropical and temperate forests can increase local temperatures 0.3° to 2°C (Hesslerová et al., 2018; Lejeune et al., 2018; Zeppetello et al., 2020) and this effect can extend up to 50 km (Cohn et al., 2019).

In Amazon rainforests, the relatively lower buffering capacity for plant moisture during drought increases 10 the risk of tree mortality and, combined with increased heat from climate change and fire from deforestation, 11 the possibility of a tipping point of extensive forest dieback and a biome shift to grassland (Oyama and 12 Nobre, 2003; Sampaio et al., 2007; Lenton et al., 2008; Nepstad et al., 2008; Malhi et al., 2009; Salazar and 13 Nobre, 2010; Settele et al., 2014; Lyra et al., 2016; Zemp et al., 2017b; Brando et al., 2020). This could 14 occur at a 4-5°C temperature increase above the pre-industrial period (Salazar and Nobre, 2010). Under 15 RCP8.5, half of Amazon tropical evergreen forest could shift to grassland through drought-induced tree 16 mortality and wildfire, but lower emissions (RCP4.5) could limit the loss to ~5% (Lyra et al., 2016). 17 Precipitation declines from reduced evapotranspiration inputs after forest loss could cause additional 18 Amazon forest loss of one-quarter to one-third (Zemp et al., 2017a). Similarly, in Guinean tropical deciduous 19 forest in Africa, climate change under RCP8.5 could increase mortality 700% by 2100 or 400% under lower 20 emissions (RCP4.5; Claevs et al., 2019). These projections indicate risks of climate change-induced tree 21 mortality reducing tropical forest areas in Africa and South America up to half under a 4°C increase above 22 the pre-industrial period, but a lower projection of a 2°C increase could limit the projected increases in tree 23 mortality (robust evidence, high agreement). 24

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Temperate and boreal forests possess greater diversity of physiological traits related to plant hydraulics, so 26 they are more buffered against drought than tropical forests (Anderegg et al., 2018). Nevertheless, in 27 temperate forests, drought-induced tree mortality under RCP8.5 could cause the loss of half of northern 28 hemisphere conifer forest area by 2100 (McDowell et al., 2016). In the western United States, one-tenth of 29 forest area is highly vulnerable to drought-induced mortality under RCP8.5 by 2050 (Buotte et al., 2019). In 30 California, increased evapotranspiration in Sierra Nevada conifer forests increases the potential fraction of 31 the area at risk of tree mortality 15-20% per degree Celsius (Goulden and Bales, 2019). In Alaska, fire-32 induced tree mortality from climate change under RCP8.5 could reduce the extent of spruce forest (Picea 33 sp.) 8–44% by 2100 (Pastick et al., 2017). Under RCP8.5, tree mortality from drought, wildfire, and bark 34 beetles could reduce timber productivity of boreal forests in Canada by 2100 below current levels (Boucher 35 et al., 2018; Chaste et al., 2019; Brecka et al., 2020). In Tasmania, projected increases in wildfire (Fox-36 Hughes et al., 2014) increase risks of mortality in mesic vegetation (Harris et al., 2018b) and threaten the 37 disappearance of the long-lived endemic pencil pine (Athrotaxis cupressoides) (Holz et al., 2015; Worth et 38 al., 2016) and temperate montane rainforest (Mariani et al., 2019). These projections indicate risks of climate 39 change-induced tree mortality reducing some temperate forest areas by half under emissions scenarios of 40 2.5-4°C above the pre-industrial period (medium evidence, high agreement). 41

43 2.5.3.4 Risk to Terrestrial Ecosystem Carbon Stocks

44 Globally, increasing atmospheric CO₂ enhances the terrestrial sink but temperature increases constrain it. 45 reflecting biological process understanding, highlighted in previous IPCC reports (high confidence). 46 Analyses of atmospheric inversion model output and spatial climate data indicate a sensitivity of net 47 ecosystem productivity to CO₂ fertilisation of 3.1 ± 0.1 to 8.1 ± 0.3 Gt per 100 ppm CO₂ (~1°C increase) and 48 49 a sensitivity to temperature of -0.5 ± 0.2 , to -1.1 ± 0.1 Gt per degree Celsius (Fernandez-Martinez et al., 2019). The future of the global land carbon sink (Section 2.4.4.4) nevertheless remains highly uncertain 50 because (i) of regionally complex interactions of climate change and changes in atmospheric CO_2 with 51 vegetation, soil and aquatic processes, (ii) episodic events such as heat-waves or droughts (and related 52 impacts through mortality, wildfire or insects, pests and diseases, (Section 2.5.5.2, 2.5.5.3) so far are only 53 incompletely captured in carbon cycle models, and (iii) legacy effects from historic land-use change and 54 environmental changes are incompletely captured but likely to decline in future, and (iv) lateral carbon 55 transport processes such as export of inland waters or erosion are incompletely understood and modelled (56 AR6 WGI Chapter 5; Pugh et al., 2019a; Friedlingstein et al., 2020; Krause et al., 2020). Enhanced carbon 57

losses from terrestrial systems further limit the available carbon budget for global warming staying below
 1.5°C (Rogelj et al., 2018). Analyses of satellite remote sensing and ground-based observations has indicated

that between 1982 and 2015 the CO_2 fertilisation effect has already declined, implying a negative climate

system feedback (Wang et al., 2020c). Peatlands, permafrost regions and tropical ecosystems are particularly
 vulnerable due to their large carbon stocks in combination with over-proportional warming, increases in

vulnerable due to their large carbon stocks in combination with over-proportional warming, increases in
 heatwaves and droughts and/or a complex interplay of climate change and increasing atmospheric CO₂

7 (Section 2.5.2.8, 2.5.2.9, 2.5.3.2).

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Model projections suggest under all warming scenarios a reduction of permafrost extent and potentially large 9 carbon losses (AR6 WGI Chapter 5). Already a mean temperature increase of 2°C could reduce the total 10 permafrost area extent by ca 5-20% by 2100 (Comyn-Platt et al., 2018; Yokohata et al., 2020). Associated 11 CO₂ losses of order of 15 up to nearly 70 GtC by 2100 have been projected across a number of modelling 12 studies (Schneider von Deimling et al., 2015; Comyn-Platt et al., 2018; Yokohata et al., 2020). Limiting the 13 global temperature increase to 1.5°C, compared to 2°C could reduce projected permafrost CO₂ losses by 14 2100 by (median) 24.2 GtC (calculated for 3m depth (Comyn-Platt et al., 2018). Losses are possibly 15 underestimated in those studies that consider only upper permafrost layers. Likewise, the actual committed 16 carbon loss may well be larger (e.g., eventually a loss of ca. 40% of today's permafrost area extent if climate 17 is stabilised at 2°C above pre-industrial levels) due to the long time-scale of warming in deep permafrost 18 layers (Chadburn et al., 2017). It is unknown at which level of global warming abrupt permafrost collapse 19 estimated to enhance CO₂ emission by 40% in 2300 in a high emissions scenario, compared to gradual thaw 20 emissions (Turetsky et al., 2020) would have to be considered an important additional risk. Large 21 uncertainties arise also from interactions with changes in surface hydrology and/or northward migrating 22 woody vegetation as climate warms, which could dampen or even reverse projected net carbon losses in 23 some regions (McGuire et al., 2018a; Mekonnen et al., 2018; Pugh et al., 2018) so overall there is low 24 confidence on how carbon-permafrost interactions will affect future carbon cycle and climate, although net 25 carbon losses and thus positive (amplifying feedbacks) are *likely* (Sections 2.5.2.10, 2.5.3.5; Shukla et al., 26 2019). See also AR6 WGI (Chapter 5) for discussion of impacts of higher emission and warming scenarios. 27

28 Peatland carbon is estimated as ca. 550-1000 GtC in northern latitudes (many of these peatlands would be 29 found in permafrost regions) (Turetsky et al., 2015; Nichols and Peteet, 2019) and > 100 GtC in tropical 30 regions (Turetsky et al., 2015; Dargie et al., 2017). Both for northern mid- and high-latitude and for tropical 31 peatlands a shift from contemporary CO₂ sinks to sources were simulated in high warming scenarios (Wang 32 et al., 2018a; Qiu et al., 2020). Due to the lack of large-scale modelling studies, the confidence on climate 33 change impacts on peat carbon uptake and emissions is low. The largest risk to tropical peatlands is expected 34 to arise from drainage and conversion to forestry or agriculture, outpacing impacts of climate change (Page 35 and Baird, 2016; Leifeld et al., 2019; Cooper et al., 2020) although the magnitude of possible carbon losses 36 are uncertain and depend strongly on socio-economic scenarios. (Sections 2.4.3.8, 2.4.4.2; 2.4.4.2, 2.5.2.8) 37 38

For tropical and sub-tropical regions the interplay of atmospheric CO₂ with precipitation and temperature 39 becomes of particular importance for future carbon uptake, since in warm and dry environments, elevated 40 CO₂ fosters plants with C3 photosynthesis and enhances their water use efficiency relative to C4 species 41 (Moncrieff et al., 2014a; Midgley and Bond, 2015; Knorr et al., 2016a). As a consequence, enhanced woody 42 cover is expected to occur in future especially in mesic savannas, while in xeric savannas an increase in 43 woody cover would occur in regions with enhanced precipitation (Criado et al., 2020). Even though semi-44 arid regions have dominated the recent decades' global trend in land CO₂ uptake (Ahlström et al., 2015), so 45 far most studies that investigated future climate change impacts on savanna ecosystems have concentrated on 46 changes in areal extent (2.5.2.5), rather than on carbon cycling, with *medium confidence* on increasing 47 woody:grass ratios (Moncrieff et al., 2014a; Midgley and Bond, 2015; Moncrieff et al., 2016; Criado et al., 48 49 2020). Increases in woody vegetation in what is now grass-dominated would possibly come with a carbon benefit, for instance a broad range of future climate and CO₂ changes were found to enhance vegetation C 50 storage in Australian savannas (Scheiter et al., 2015). Results from a number of field experiments indicate 51 however, that impacts on total ecosystem carbon storage may be smaller, due to a loss in belowground 52 carbon (Coetsee et al., 2013; Wigley et al., 2020). Nunez et al., 2021) critique existing incentives to promote 53 invasion of non-native trees into treeless areas as a means of carbon sequestration, raising doubts about the 54 effects on fire, albedo, biodiversity and water yield (see Box 2.2). 55

Substantial climate-change driven impacts on tropical tree cover and vegetation type are projected in all 1 studies, irrespective of whether or not the degree amounts to a forest "dieback" (Sections 2.4.3.6, 2.4.4.3; 2 2.5.2.6, 2.5.3.3; AR6 WGI Chapter 5) (Davies-Barnard et al., 2015; Wu et al., 2016a; Zemp et al., 2017a). 3 Accordingly, models also suggest a continuation of tropical forests acting as carbon sinks (Huntingford et 4 al., 2013; Mercado et al., 2018). A recent study, combining field plot data with statistical models, (Hubau et 5 al., 2020) indicates that in the Amazonian and possibly also in the African forest the carbon sink in 6 aboveground biomass has already declined over the three decades to 2015. This trend is distinct in the 7 Amazon, whereas data from Africa suggest a possible decline after 2010. the authors estimate the vegetation 8 carbon sink in 2030-2040 to decline to zero (-0.5–0-46) Pg C a⁻¹ in the Amazon and 0.26 (0.04–0.47, a loss 9 of 14% compared to present) Pg C a⁻¹ in Africa. Their results suggest that CO₂ fertilisation is over time 10 outweighed by impacts of higher temperatures and drought, enhancing tree mortality and diminishing 11 growth. The degree of thermal resilience of tropical forest still remains uncertain, however (Sullivan et al., 12 2020). 13 14 The lack of simulation studies that seek to quantify all important interacting factors (CO₂, drought and fire) 15 for future carbon-cycling in savannas and tropical forests, and the apparent disagreement between trends 16 projected in models compared to data-driven estimates results in low confidence regarding the direction or 17 magnitude of carbon flux and pool size changes. Similar to tropical peatlands, given projected human 18 population growth and socio-economic changes. the continued conversion of forests and savannas into 19 agricultural or pasture systems very likely poses a significant risk of rapid carbon loss which will amplify 20 climate change induced risks substantially (high confidence) (Sections 2.5.2.10, 2.5.3.5; Aragao et al., 2014; 21 Searchinger et al., 2015; Aleman et al., 2016; Nobre et al., 2016). 22 23 The impacts of climate-induced altered animal composition and trophic cascades on global land ecosystem 24 carbon cycling are as yet unquantified (Schmitz et al., 2018) even though climate change is expected to lead 25 to shifts in consumer-resource interactions that also contribute to losses of top-predators or top herbivores 26 (Sections 2.4.2.2, 2.5.1.3, 2.5.4; Lurgi et al., 2012; Damien and Tougeron, 2019). Cascading trophic effects 27 triggered by top predators or the largest herbivores propagate through food webs and reverberate through to 28 the functioning of whole ecosystems, altering notably productivity, carbon and nutrient turnover and net 29 carbon storage (medium confidence) (Wilmers and Schmitz, 2016; Sobral et al., 2017; Stoner et al., 2018). 30 Across different field experiments, the ecosystem consequences of the presence or absence of herbivores and 31 carnivores have been found to be quantitatively as large as the effects of other environmental change drivers 32 such as warming, enhanced CO₂, fire or variable nitrogen deposition (medium confidence) (Hooper et al., 33 2012; Smith et al., 2015). Some local and regional-scale modelling experiments have begun to explore 34 animal impacts on vegetation dynamics and carbon and nutrient cycling (Pachzelt et al., 2015; Dangal et al., 35 2017; Berzaghi et al., 2019). Given that turnover rate is a chief factor that determines future land ecosystem 36 carbon dynamics and hence carbon-climate feedbacks (Friend et al., 2014). To improve projections, it is 37 imperative to better quantify the broader role of carnivores, grazers, browsers, and the way these interact in 38 global studies of how ecosystems respond to climate change.

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2.5.3.5 Feedbacks between Ecosystems and Climate

The possibility of feedbacks and interactions between climate drivers and biological systems or ecological 43 processes was identified as a significant emerging issue in AR5, and has since also been highlighted in the 44 CRCCL and the Special Report on 1.5°C. It is virtually certain that land cover changes affect regional and 45 global climate through changes to albedo, evapotranspiration and roughness (very high confidence) (Perugini 46 et al., 2017). There is growing evidence that biosphere-related climate processes are being affected by 47 climate change in combination with disturbance and land use change (high confidence) (Jia et al., 2019). It is 48 49 virtually certain that land surface change caused by disturbances such as forest fire, hurricanes, phenological changes, insect outbreaks and deforestation affect carbon, water, and energy exchanges, thereby influencing 50 weather and climate (very high confidence) (Table 2.4; Figure 2.10; Bright et al., 2013; Brovkin et al., 2013; 51 Naudts et al., 2016; Prăvălie, 2018). 52

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Table 2.4: Terrestrial and freshwater ecosystem feedbacks which affect the Earth's climate system dynamics, following
 (Prăvălie, 2018).

Perturbation	Implications for Warming/Feedback Mechanism the Earth's Climate System Dynamics
Phenological change	Increased primary productivity and plant growth with CO ₂ fertilisation (Mao et al., 2016; Wang et al., 2018a); Increasing growing season length (Peñuelas et al., 2009; Barichivich et al., 2013); reduced diurnal temperature range through evapotranspiration (mid-latitudes) and albedo (high latitudes) caused by vegetation greening (Jeong et al., 2011); increased CO ₂ storage in biomass (cooling) (Keenan et al., 2014); Reduced albedo in snow-covered regions as canopies become taller and darker (warming); increased evapotranspiration, a key component of the global water cycle and energy balance which influences global rainfall, temperature, and atmospheric motion (Zeng et al., 2017).
Insect outbreaks	Reduced carbon uptake and storage (warming); Increased surface albedo (cooling) (Landry et al., 2016); increased CO ₂ emissions (warming); decreased leaf area index and gross primary productivity (Ghimire et al., 2015), leading to reduced evapotranspiration and increased land surface temperature (Bright et al., 2013).
Range shifts	Reduced albedo in snow-covered regions as trees expand poleward (warming) (Chae et al., 2015); enhanced permafrost thawing; expansion of insect outbreak range, increasing forest impact (Pureswaran et al., 2018); biome dependent changes in albedo and evapotranspiration regimes (Naudts et al., 2016). Reduction in snow and ice albedo in freshwater due to loss of ice (warming) (Lang et al., 2018).
Die-off and large-scale mortality events	Decreased Gross primary productivity (GPP); decline in carbon storage (warming); increased CO ₂ emissions; increased solar radiation, reduced soil moisture, higher surface runoff; albedo effects (Lewis et al., 2011; Prăvălie, 2018)
Deforestation	Reduced carbon storage (warming) (Pugh et al., 2019a); increase in (regional) surface air temperature due to reduced evaporation (less cooling); increased albedo in high-latitude systems (regional radiative cooling) (Loranty et al., 2014); increased air temperature and diurnal temperature variation (Alkama and Cescatti, 2016), locally and globally (Winckler et al., 2019); reduced precipitation (Perugini et al., 2017); decreased biogenic volatile organic compounds (BVOC) and aerosol emissions (warming through direct and indirect aerosol effects; cooling associated with reduction in atmospheric methane (Jia et al., 2019)
Forest degradation	Reduced carbon storage (warming) (de Paula et al., 2015; Bustamante et al., 2016; de Andrade et al., 2017; Mitchard, 2018)
Fragmentation	Carbon losses because biomass is less developed in forest edges (Pütz et al., 2014; Chaplin-Kramer et al., 2015; Haddad et al., 2015)
Air pollution	Decreased plant productivity, transpiration and carbon sequestration in forest with lower biomass due to ozone toxicity (Sitch et al., 2007; Ainsworth et al., 2012); increased (regional) productivity due to increase in diffuse solar radiation caused by terrestrial aerosols (Xie et al., 2021)
Declining populations of megafauna	Changes to physical and chemical properties of organic matter, soils and sediments influence carbon uptake and storage (Schmitz et al., 2018); increased or decreased carbon storage biomass and carbon storage, with differences across biomes determined by floristic structure and animal size (Bello et al., 2015; Osuri et al., 2016; Peres et al., 2016; He et al., 2017; Berzaghi et al., 2018; Schmitz et al., 2018; He et al., 2019)
Fire	Increased carbon and aerosol emissions(van der Werf et al., 2017); surface warming (Liu et al., 2019b); albedo effect dependent on ecosystem and species-level traits (Rogers et al., 2015; Chen et al., 2018a) (initial albedo decrease post-fire; increased albedo where snow exposure is increased by canopy removal and species composition changes during recovery); black carbon deposition on snow and sea ice (short-term) (Randerson et al., 2006); indirect increases in carbon emissions due to soil erosion (Caon et al., 2014)
Change in forest composition Woody encroachment in non-forested ecosystems	Reduced carbon storage due to decline in biomass (warming) (McIntyre et al., 2015) Reduced production, increased water use, reduced albedo and altered land-atmosphere feedbacks; increased carbon storage in woody savannas (Zhou et al., 2017; Mureva et al., 2018); Uncertain
Net Primary Productivity (NPP) shifts	feedbacks to C cycle (some suggest an increase, others a decrease) Reduced albedo following high-latitude expansion of trees caused by photosynthetic enhancement of growth (cooling); increased photosynthesis and net ecosystem production (NEP) (Fernandez- Martinez et al., 2019); increased NPP in N-limited ecosystems due to increased nitrogen deposition from agriculture and combustion (Du and de Vries, 2018; Schulte-Uebbing and de Vries, 2018); Nutrient limited lakes are likely to become less productive, while nutrient rich lakes are likely to become more productive due to warming induced prolongation of stable stratification (Adrian et al., 2016; Kraemer et al., 2017).
Biogeochemical shifts	Decline in carbon storage due to nitrogen limitation in N limited systems (warming)(Reich et al., 2014; Wieder et al., 2015); increased carbon storage on land(Peñuelas et al., 2013) and in lakes (Heathcote et al., 2015; Mendonça et al., 2017); Increase in CO_2 and CH4 emissions from freshwater ecosystems due to increased eutrophication (DelSontro et al., 2018), the imbalance

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between losses and gains of CO₂ by photosynthesis and respiration (metabolic theory of ecology), enhanced emissions from exposed river and lake sediments during droughts and re-wetting (Marcé et al., 2019; Keller et al., 2020), enhanced CH4 ebullition of seasonally hypoxic lakes (Aben et al., 2017; DelSontro et al., 2018; Bartosiewicz et al., 2019; Beaulieu et al., 2019; Sanches et al., 2019), increased transfer of organic carbon from land to water (particularly in permafrost areas) (Wauthy et al., 2018)

Terrestrial ecosystem feedbacks which affect the Earth's climate system dynamics

Perturbations & implications for climate system dynamics for the three global forest biomes.

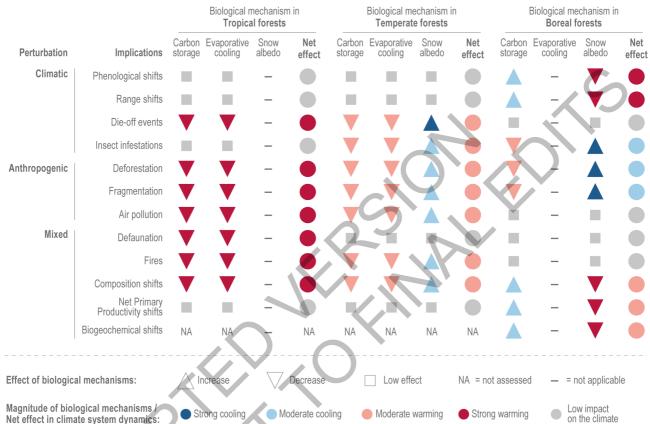


Figure 2.10: Terrestrial ecosystem feedbacks, which affect the Earth's climate system dynamics. Perturbations and implications for climate system dynamics (warming/cooling) are shown for the three global forest biomes (adapted from Figure 5, Prăvălie, 2018). The strength of the mechanism is estimated in general terms based on the magnitude of carbon storage and evaporative cooling processes that characterise each forest biome (Bonan, 2008). Carbon storage includes forest biomass, without accounting for carbon dynamics in soil, peat and underlying permafrost deposits. Implications of biogeochemical shifts were only estimated in relation to the intensification of the carbon cycle and increase in biomass at high latitudes, assuming N availability for the stoichiometric demands of forest vegetation).

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Feedbacks can be positive or negative (i.e., amplify or dampen the original forcing), vary spatially and 13 seasonally, and act over large geographic areas and long time periods (>decades), making them difficult to 14 observe and quantify directly (AR6 WGI Chapter 5; Schimel et al., 2015). Due to the positive impacts of 15 CO₂ on vegetation growth and ecosystem carbon storage (high confidence) (Sections 2.4.4.4; 2.5.5.4; AR6 16 WGI chapter 5), the associated climate feedback is negative (increased removal of atmospheric CO_2 and 17 dampened warming, compared to absence of the feedback). By contrast, projected global losses of carbon in 18 warmer climates (AR6 WGI Chapter 5) imply a positive climate feedback. Chapter 5 of WGI assesses an 19 overall increase in land carbon uptake through the 21st century. However, the overall strength of the carbon 20 cycle-climate feedback remains very uncertain. One of the underlying reasons may be complex interactions 21 with ecosystem water balance and nitrogen and phosphorous availability, which are poorly constrained by 22 observational evidence and incompletely captured in Earth System Models (AR6 WGI Chapter 5, Section 23 2.5.2.10; Huntzinger et al., 2017). 24

Land ecosystems contribute substantially to global emissions of nitrous-oxides and methane As with CO₂, 1 these emissions respond both directly and indirectly to atmospheric CO₂ concentration and climate change, 2 which gives rise to potential additional biogeochemical feedbacks in the climate system. A large part of these 3 emissions stem from land and water management, such as fertiliser application, rice production, aquaculture 4 or animal husbandry (Jia et al., 2019). However, nearly 60% of total nitrous-oxide emissions (2007-2016) is 5 estimated to stem from natural ecosystems, especially in the tropics (AR6 WGI Chapter 5; Tian et al., 2019), 6 while freshwater wetlands and peatlands are estimated to contribute between 83% (top-down estimates) and 7 40% (bottom up estimates) of total natural CH₄ (and 31% or 20% of total methane emissions, respectively) 8 for the period 2008-2017 (AR6 WGI Chapter 5). Median CH4 emissions from northern latitude wetlands in 9 2100 were estimated to be 12.1 and 13.5 Pg C in emission scenarios leading to 1.5°C and 2°C warming, 10 respectively (Comyn-Platt et al., 2018). Likewise, global warming was attributed to soil N₂O emission 11 increases since the pre-industrial period of 0.8 (0.3–1.3) Tg N a⁻¹ (Tian et al., 2020). Overall, climate 12 feedbacks from future altered land ecosystem emissions of CH₄ or N₂O are uncertain, but expected to be 13 small (AR6 WGI Chapter 5). 14 15 Changes in regional biodiversity are integral parts of ecosystem-climate feedback loops, including and 16 beyond carbon-cycle processes (Figure 2.10; Table 2.4). For instance, the impacts of climate-induced altered 17 animal composition and trophic cascades on ecosystem carbon turnover (see Sections 2.4.4.4; 2.5.3.4) could 18 be a substantive contribution to carbon-climate feedbacks (low confidence). Additional surface-atmosphere 19 feedbacks that arise from changes in vegetation cover and subsequently altered albedo, evapotranspiration or 20 roughness (often summarised as biophysical feedbacks) can be regionally relevant and could amplify or 21 dampen vegetation cover changes (Jia et al., 2019). 22 23

Climate-induced shifts towards forests in what is currently tundra would be expected to reduce regional 24 albedo especially in spring but also during parts of winter when trees are snow-free (whereas tundra 25 vegetation would be covered in snow), which amplifies warming regionally (high confidence) (Perugini et 26 al., 2017; Jia et al., 2019). Trees would also enhance momentum absorption compared to low tundra 27 vegetation thus impacting surface-atmosphere mixing of latent and sensible heat fluxes (Jia et al., 2019). 28 Boreal forests insulate and stabilize permafrost and reduce fluctuations of ground temperature: the amplitude 29 of variation of ground surface temperatures was 28°C in a forested site, compared to 60°C in nearby 30 grassland (Section 2.5.2.7; Bonan, 1989; Stuenzi et al., 2021a; Stuenzi et al., 2021b). Likewise, a shift in 31 moist tropical forests towards vegetation with drought-tolerant traits could possibly reduce 32 evapotranspiration, increase albedo, alter heat transfer at the surface and lead to a negative feedback to 33 precipitation (Section 2.5.2.6; Jia et al., 2019). In savannas, restoration of woody vegetation has been shown 34 to enhance cloud formation and precipitation in response to enhanced transpiration and turbulent mixing, 35 leading to a positive feedback on woody cover (Syktus and McAlpine, 2016). While this has not yet been 36 systematically explored, similar feedbacks might also emerge from a CO₂-induced woody cover increase in 37 savannas (low confidence) (Section 2.5.2.5). 38 39

Since biophysical feedbacks can contribute to both surface temperature warming or cooling, analyses so far 40 suggest that, at global scale, the net impact on climate change is small (Perugini et al., 2017; Jia et al., 2019), 41 unless these feedbacks also accelerate vegetation mortality and lead to substantive carbon losses (Zemp et 42 al., 2017a; Lemordant and Gentine, 2019). More than one third of the Earth's land surface has at least 50% 43 of its evapotranspiration regulated by vegetation, and in some regions between 40 and >80% of the land's 44 evaporated water is returned to land as precipitation. Locally, both, direct human-mediated as well as climate 45 change-mediated vegetation cover change can therefore notably affect annual average freshwater availability 46 to human societies, especially if negative feedbacks amplify vegetation cover reduction, reduced 47 evapotranspiration and reduced precipitation (medium confidence) (Keys et al., 2016; Keys and Wang-48 49 Erlandsson, 2018).

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Since AR5, freshwater ecosystems (lakes, reservoirs, rivers, ponds) have been increasingly recognised as important sources of greenhouse gas emissions (CO₂, CH₄, N₂O) into the atmosphere. Key mechanisms which contribute to rising GHG emissions from freshwater ecosystems are the temperature imbalance between photosynthesis and respiration (respiration increases more than photosynthesis with rising temperature), CO₂ and CH₄ emissions from exposed sediments during droughts, increased matter transport from land to water, changes in water retention time in rivers and lakes, and temperature effects on lake

57 stratification and anoxia, favouring CH₄ emissions.

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1 DelSontro et al. (2018) assembled the largest global dataset to date on emission rates from lakes of CO₂, CH₄ 2 and N₂O and found that they co-vary with lake size and trophic state. They estimated that moderate global 3 increases in eutrophication of lakes could translate to 5-40% increases in the GHG effects in the atmosphere. 4 Moreover, they estimated that greenhouse gas emissions from lakes and impoundments in past decades 5 accounted for 1.25-2.30 Pg C-CO₂ yr⁻¹ (DelSontro et al., 2018), thus around 20% of global burning fossil-6 fuel CO₂ emission (9.4 PgC-CO₂ yr⁻¹ (Friedlingstein et al., 2020). 7 8 Global warming will strongly enhance freshwater CH₄ emissions through a disproportionate increase in 9 ebullition (gas flux) by 6–20% per 1°C increase in water temperature (Aben et al., 2017). It can be expected 10that ongoing eutrophication enhanced by climate change-related increases in sediment nutrient release and 11 organic carbon and nutrient loading from catchments will enhance CH4 ebullition at a global scale (Aben et 12 al., 2017; DelSontro et al., 2018; Bartosiewicz et al., 2019; Beaulieu et al., 2019; Sanches et al., 2019). The 13 strongest increase in ebullition is expected in shallow waters where sediment temperatures are strongly 14 related to atmospheric temperature (Aben et al., 2017). Given that small ponds and shallow lakes are the 15 most abundant freshwater ecosystems globally they may become hot spots of CH₄ ebullition in the future 16 (Aben et al., 2017). On average CH₄, CO₂, N₂O account for 75%, 23, and 2% of the total CO₂ equivalent 17 emissions, respectively in lakes (DelSontro et al., 2018). 1819 Further, the exposure of lake and river sediments during droughts activates decomposition of buried organic 20 carbon. In dry river beds, mineralisation of buried organic matter is likely to increase with climate change as 21 anoxic sediments are oxygenated downwards during drying along with pulses of microbial activity following 22 rewetting of desiccated sediment. Conservative estimates indicate that adding emissions from exposed 23 sediments of dry inland waters across diverse ecosystem types and climate zones to current global estimates 24 of CO₂ emissions could result in a 6% (~0.12 Pg C y⁻¹) increase of total inland water CO₂ emission rates 25 covering streams and rivers (334 mmol m⁻² day⁻¹), lakes and reservoirs (320 mmol m⁻² day⁻¹) and small 26 ponds (148 mmol m⁻² day⁻¹) (Marcé et al., 2019; Keller et al., 2020). 27 28

Overall, uncertainty in the quantity of carbon fluxes within freshwater ecosystems and between terrestrial 29 and freshwater systems and subsequent emissions to the atmosphere remain very high (Raymond et al., 30 2013; Catalán et al., 2016; Stanley et al., 2016; Evans et al., 2017; Drake et al., 2018; Seekell et al., 2018; 31 Sanches et al., 2019; Bodmer et al., 2020; Keller et al., 2020) (see Table 2.SM., see also Chapter 5 of WGI). 32 Projections of carbon fluxes are e.g. challenged by the complex interaction between rising water 33 temperature, loss of ice, changes in hydrology, ecosystem productivity, increase in extreme events, and 34 variation in terrestrial matter transport. While we are still short in empirical data, particularly in the tropics 35 (DelSontro et al., 2018), improvements in sensor technology (Eugster et al., 2011; Gonzalez-Valencia et al., 36 2014; Maeck et al., 2014; Delwiche et al., 2015) and the use of statistically robust survey designs (Beaulieu 37 et al., 2016; Wik et al., 2016) have improved the accuracy of GHG emission rate measurements in freshwater 38 ecosystems. Global networks such as GLEON (Global Lakes Ecological Observatory Network) increasingly 39 allow a global view of carbon fluxes improving estimates of the contribution of freshwater ecosystems to 40 global GHG emissions to the atmosphere. 41

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In summary (Drake et al., 2018) aggregated contemporary estimates of CO₂ and CH₄ emissions from 43 freshwater ecosystems with global estimates made by (Raymond et al., 2013) and arrived at an estimate of 44 3.9 Pg C yr⁻¹. Rivers and streams accounted for 85% of the emissions and lakes and reservoirs for 15% 45 (Raymond et al., 2013). This trend will continue under scenarios of nutrient loading to inland waters over the 46 next century where inland water increased CH₄ emission has an atmospheric impact of 1.7–2.6 Pg C-CO₂-eq 47 y^{-1} , which is equivalent to 18–33% of annual CO₂ emissions from burning fossil fuels (*medium evidence*, 48 49 *medium agreement, medium confidence*) (Beaulieu et al., 2019). For comparison, annual uptake of CO_2 in land ecosystems is estimated as $3.4 (\pm 0.9)$ PgC/yr (Friedlingstein et al., 2020). The freshwater numbers 50 combine CO₂ and CH₄ and are thus not directly comparable. However, they are indicative for the importance 51 to account better for freshwater systems in global carbon budgets. 52

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54 2.5.3.6 Risks to Freshwater Ecosystem Services: Drinking Water, Fisheries and Hydropower

AR5 named water supply and biodiversity as freshwater ecosystem services vulnerable to climate change.
 We discuss risks to these and to additional services identified by model projections based both on climate-

change scenarios (Schröter et al., 2005; Boithias et al., 2014; Huang et al., 2019; Jorda-Capdevila et al.,
 2019) and on the Common International Classification of Ecosystem Services (*high agreement, high confidence*) (CICES, 2018). Effects of floods, droughts, permafrost and glacier melting on global changes in
 water quality, particularly with respect to contamination with pollutants, are described in Section 4.2.6.

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6 2.5.3.6.1 Risks to quantity and quality of drinking water

7 Forests and other vegetated ecosystems assist production of drinkable water by facilitating infiltration of

8 rainfall and snowfall into the ground, where water either moves through the soil saturated zone to supply

9 streams and other surface waters or infiltrates further to recharge groundwater aquifers (Ellison et al., 2012;

- Bonnesoeur et al., 2019). Globally, 4 billion people depend on forested watersheds for drinking water
- (Mekonnen and Hoekstra, 2016). Chapter 4 assesses the physical science of water supply, including precipitation, runoff, and hydrology, and social aspects of human water use. This section assesses ecological
- precipitation, runoff, and hydrology, and social aspects of human water use. This section aspects of risks to freshwater supplies for people.
- 14

Reduction of vegetation cover following wildfires (Section 2.5.5.2) and tree mortality (Section 2.5.5.3) can 15 reduce long-term water infiltration, increase soil erosion and flash floods, and release sediment that degrades 16 drinking water quality. Widlfires increase impacts of extreme precipitation events due to climate change, 17 which contribute to increased surface runoff and hence to increased risks of land erosion, landslides and 18 flooding (Ebel et al., 2012; Robinne et al., 2020). Under current conditions, nearly half of global land area is 19 at moderate to high risk of water scarcity due to wildfires (Robinne et al., 2018; Robinne et al., 2020). From 20 1984 to 2014 wildfires in the western USA affected 6-11% of stream and river length (Ball et al., 2021). 21 Under a high emissions scenario of a 3.5°C temperature increase post-fire erosion across the western USA 22 could double sedimentation and degrade drinking water quality in one-third of watersheds by 2050 (Sankey 23 et al., 2017). In Brazil, post-fire vegetation loss tends to increase runoff, reduce infiltration, and reduce 24 groundwater recharge and flow of springs (Rodrigues et al., 2019). Runoff from wildfires can contain 25 dissolved organic carbon precursors for the formation of carcinogenic trihalomethanes during water 26 chlorination for drinking (Uzun et al., 2020), plus chromium, mercury, selenium, and other toxic trace metals 27 (Burton et al., 2016; Burton et al., 2019). 28

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Net effects of deforestation and afforestation on runoff and water supply depend on local factors, leading to 30 conflicting evidence for effects of land cover change (Ellison et al., 2012; Chen et al., 2021b), but 31 combinations of climate change and deforestation are projected to reduce water flows (Olivares et al., 2019). 32 In southern Thailand, the combination of conversion of forest to rubber plantations and a one-third increase 33 in rainfall could increase erosion and sediment load 15% (Trisurat et al., 2016). In the watershed that 34 supplies São Paulo, Brazil, afforestation could increase water quantity and quality (Ferreira et al., 2019). In 35 most regions with dry or Mediterranean subtropical climates, climate change reduces renewable surface 36 water and groundwater resources significantly (Doell et al., 2015). In northeast Spain, reduced precipitation 37 and vegetation cover under a high emissions scenario of a 3.5°C temperature increase could reduce drinking 38 water supplies by half by 2100 (Bangash et al., 2013). 39 40

Changes in algal biomass development and spread of cyanobacteria blooms, related to global warming,
resemble those triggered by eutrophication with well-known negative effects on the services lakes provide,
particularly for drinking water provision and recreation (*robust empirical evidence, high agreement, high confidence*) (Carvalho et al., 2013; Adrian et al., 2016; Gozlan et al., 2019).

45 Based on a 10% increase in precipitation, (de Wit et al., 2016) estimated increased mobilisation of organic 46 carbon from soils to freshwaters by at least 30%, demonstrating the importance of climate wetting for the 47 carbon cycle. Browning negatively affects the taste of drinking water and may be difficult to address 48 49 (Kothawala et al., 2015; Mineau et al., 2016; Kritzberg et al., 2020). It also often reduces attractives for recreational purposes, especially swimming (Arthington and Hadwen, 2003; Keeler et al., 2015). Based on a 50 worst-case climate scenario until 2030 (Weyhenmeyer et al., 2016) projected an increase in browning of 51 lakes and rivers in boreal Sweden by a factor of 1.3. The chemical character of dissolved organic matter, as 52 modified by climate change (Kellerman et al., 2014), determines its amenability to removal by water 53 treatment (Ritson et al., 2014). Therefore, in order to provide safe and acceptable drinking water, more 54 advanced, more expensive and more energy/resource intensive technical solutions may be required 55 (Matilainen et al., 2010). 56

In summary, climate change increases risks to the integrity of watersheds and provision of safe, acceptable freshwater to people (*medium evidence, medium agreement*).

4 2.5.3.6.2 Risks to freshwater fisheries and biodiversity

Climate change will increase water temperatures and decrease dissolved oxygen levels (Section 2.3.1) impacting freshwater fisheries which form an important ecosystem service (Vári et al., 2021). People living in the vicinity of cold lakes will be affected by projected losses of ice. In a worst-case scenario (air temperatures increase of 8°C), 230,400 lakes and 656 million people in 50 countries will be impacted (Reid et al., 2019; Sharma et al., 2019). Winter ice fishing (Orru et al., 2014), transportation via ice roads (Prowse et al., 2011) and cultural activities (Magnuson and Lathrop, 2014) are ecosystem services at stake with ongoing loss of lake ice.

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Eutrophication of central European lakes has wiped out a significant proportion of the endemic fish fauna

14 (Vonlanthen et al., 2012), so climate-induced further eutrophication is expected to represent an additional 15 threat to fish fauna and commercial fisheries (Ficke et al., 2007). Given that the ecological consequences of

16 lake warming may be especially strong in the tropics (Section 2.3.1.1), ecosystem services may be most

affected there. Tropical lakes support important fisheries (Lynch et al., 2016a economic; McIntyre et al.,

18 2016) providing critical sources of nutrition to adjacent human populations. These lakes are especially prone 19 to loss of deep-water oxygen due to warming, with adverse consequences for fisheries productivity and

to loss of deep-water oxygen due to warming, with adverse consequences for fisheries productivit biodiversity (*medium evidence, medium confidence*) (Lewis Jr, 2000; Van Bocxlaer et al., 2012).

21

22 Tropical lakes tend to be hotspots of freshwater biodiversity (Vadeboncoeur et al., 2011; Brawand et al.,

23 2014; Sterner et al., 2020); ancient tropical lakes such as Malawi, Tanganyika, Victoria, Titicaca, Towuti

and Matano hold thousands of animal species found nowhere else (Vadeboncoeur et al., 2011). While

biodiversity and several ecosystem services can be considered synergistic (food webs, tourism, aesthetical
 and spiritual value (Langhans et al., 2019), others can be considered antagonistic in case of a strong

and spiritual value (Langhans et al., 2019), others can be considered antagonistic in case of a strong
 ecosystem service demand (such as water abstraction, water use, food security in terms of over-exploitation).

ecosystem service demand (such as water abstraction, water use, food security in terms of over-exploitation
 Here the balance between biodiversity and ecosystem services is key (Langhans et al., 2019), where

Here the balance between biodiversity and ecosystem services is key (Langhans et al., 2019), where
 biodiversity can be integrated into water policy through Integrated Water Resource Management (IWRM)

towards nature-based solutions (Ligtvoet et al., 2017)

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32 2.5.3.6.3 Risks to hydro power and erosion control

River banks, riparian vegetation and macrophyte beds play important roles in erosion control through 33 reducing current velocities, increasing sedimentation and reducing turbidity (Madsen et al., 2001). Rates of 34 flow in rivers affect and inland navigation (Vári et al., 2021). Changing seasonality in snow-dominated 35 basins is expected to enhance hydropower production in winter, but decrease it during summer (Doell et al., 36 2015). Glacier melt changes hydrological regimes, sediment transport, and biogeochemical and contaminant 37 fluxes from rivers to oceans, profoundly influencing ecosystem services that glacier-fed rivers provide, 38 particularly provision of water for agriculture, hydropower, and consumption (Milner et al., 2017). Loss of 39 glacial mass and snowpack has already impacted flow rates, quantities and seasonality (Hock et al., 2019); 40 see AR6 WGII, Chapter 4 Water). Meltwater yields from glacier ice are likely to increase in many regions 41 during the next decades, but decrease thereafter as glaciers become smaller and smaller and finally disappear 42 (Hock et al., 2019). 43

45 2.5.4 Key Risks to Terrestrial and Freshwater Ecosystems from Climate Change

46 Among numerous risks to terrestrial and freshwater ecosystems from climate change, this IPCC chapter 47 identified five phenomena as the most fundamental risks of climate change to ecosystem integrity and the 48 ecosystem services that support human well-being that are also quantified: species losses to ecosystems, 49 increased wildfire, increased tree mortality, ecosystem carbon losses, and ecosystem structure change (Table 50 2.5, Table 2.S.4, Figure 2.11). These key risks form part of the overall assessment of key risks in Chapter 16. 51 The IPCC Fifth Assessment Report chapter on terrestrial ecosystems (Settele et al., 2014) had also identified 52 three of these key risks – species extinctions, tree mortality, ecosystem carbon losses – and a fourth – 53 invasion by non-native species. This IPCC chater assesses impacts of climate change on invasive species in 54 multiple sections with respect to different processes or systems (e.g. in Section 2.4.2.3.3), and here includes 55 this aspect in a new broader key risk of ecosystem structure change. The IPCC Fifth Assessment Report had 56 included wildfire as a mechanism of the ecosystem carbon loss key risk. Based on additional research and 57

field experience with major wildfires since then, this IPCC chapter sets wildfire apart as a specific key risk to 1 ecosystem integrity and human well-being. These different measures of risk are interconnected but approach 2 assessment of risk to terrestrial and freshwater ecosytems from different angles, using complementary 3 metrics. 4

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Species are the fundamental unit of ecosystems. As species become rare, their roles in the functioning of the 6 ecosystem diminishes, and disappears altogether if they go locally extinct (high confidence) (Isbell et al., 7 2015; Chen et al., 2018b; van der Plas, 2019; Wang et al., 2021b). Loss of species and functional groups 8 reduces the ability of an ecosystem to provide services, and lowers its resilience to climate change (high 9 confidence) (Section 2.6.7; Elmqvist et al., 2003; Cadotte et al., 2011; Harrison et al., 2014; Carlucci et al., 10 2020). For example, among crop systems, a key factor to succesful pollination is the phylogenetic diversity 11 of bee species available, more than total abundances (Drossart and Gérard, 2020). Because many species 12 have obligate interactions with, or are resources for, other species (e.g. predators and their prey, insects and 13 their host plants, plants and their mycorrhizae symbionts), loss of one species affects risk to another species 14 and, ultimately, ecosystem functioning (Mahoney and Bishop, 2017) 15

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Global rates of species extinction are accelerating dramatically (Barnosky et al., 2011), with approximately 17

10% of species having been driven extinct by humans since the late Pleistocene, principally by over-18

exploitation and habitat destruction, a rate estimated to be 1000 times higher than pre-Anthropocene 19

(natural) background extinction rates (De Vos et al., 2015). Therefore, this level—10%—of species 20

becoming endangered (sensu IUCN) due to loss of suitable climate space (Figure 2.8b), is used here as a 21

threshold moving risk to biodiversity from moderate to high, and twice that (20%) as the threshold from high 22 to very high.

23 24

Key risks assessed here are interconnected. Extinction of species is an irreversible impact of climate change, 25 has negative consequences on ecosystem integrity and functioning, and risk increases steeply with even 26 small rises in global temperature (Section 2.5.1.3; Figure 2.6; Figure 2.7; Figure 2.8). Continued climate 27 change substantially increases risks of carbon losses due to wildfires, tree mortality from drought and insect 28 pest outbreaks, peatland drying, permafrost thaw, and ecosystem structure change which could exacerbate 29 self-reinforcing feedbacks between emissions from high-carbon ecosystems and increasing global 30 temperatures (medium confidence). Thaving of Arctic permafrost alone could release 11-200 Gt carbon 31 (medium confidence). Complex interactions of climate change, land use change, carbon dioxide fluxes, and 32 vegetation changes will regulate the future carbon balance of the biosphere, processes incompletely 33 represented in earth system models. The exact timing and magnitude of climate-biosphere feedbacks and 34 potential tipping points of carbon loss are characterized by high ranges of the estimates, yet studies indicate 35 increased ecosystem carbon losses could cause extreme future temperature increases (medium confidence) 36 (Sections 2.5.2.7; 2.5.2.8; 2.5.2.9; 2.5.3.2; 2.5.3.3; 2.5.3.4; 2.5.3.5; Figure 2.10; Figure 2.11; Table 2.4; 37 Table 2.5; Table 2.S.2; Table 2.S.4) 38

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40 Table 2.5: Key risks to terrestrial and freshwater ecosystems from climate change. This IPCC chapter assesses these as 41 the most fundamental risks of climate change to ecosystem integrity and the ecosystem services that support human 42 well-being. Climate factors include the primary variables governing the risk. Non-climate factors include other 43 phenomena that can dominate or contribute to the risk. Detection and attribution comprise cases of observed changes 44 attributed predominantly or in part to anthropogenic climate change (Section 2.4.2, 2.4.3, 2.4.4, 2.4.5, Table 2.2, Table 45 2.3, Table 2.S.1). Adaptation includes options to address the risk (Section 2.6). Risk transitions (defined in Figure 2.11) 46 indicate an approximate global mean surface temperature increase, relative to the pre-industrial period (1850-1900), to 47 move from one level of risk to the other and confidence in the assessment. Table 2.S.4 provides details of the 48 temperature levels for risk transitions. Both tables provides details for the key risk burning embers diagram (Figure 49

2.11). 50

> **Biodiversity Risk:** Increasing high extinction risk (species projected loss of >50% of range) among increasing number of plant and animal species. The transition from non-detectable to moderate was based on the number of local population extinctions, major declines of sub-species and two global species extinctions and that are attributable to climate change. The transition from moderate to high is centred around 1.5°C based on a few taxa that are known from their basic biology and habitat requirements to be at risk of extinction (endangered) at 1.5°C and on the increasing number of taxa that are projected to have high extinction risk (losing >50% of their suitable climate space) affecting >10% of the species in that taxa (1000x natural background rates of extinction). The

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transition to very high comes from the increasing number of taxa projected to have >20% of species at high risk of extinction. In the worst-case scenario (10^{th} percentile of the models), some of the taxa show >50% of the species at high risk of extinction. These assessments are also weighted by role the species in the taxa play in performing ecosystem services (both to the ecosystems and to humans, e.g. pollinators, detritivores). Confidence for the moderate threshold is *high* because it is based on observed trends attributed to climate change. Confidence for future projections are is *medium* as these are based on one large study (covering more than 135,000 species) and primarily based on loss of suitable climate. Based on Sections 2.4.2, 2.5.1, 2.6.1, 2.6.6, Table 2.3, Figure 2.6, Table SM2.1 and Table SM2.2.

Climate factors	Non-climate factors	Detection and attribution	Adaptation	Risk transitions (<i>confidence</i>)
Shifts in geographic placements of climate space; loss of climate space globally; emergence of non-analogue climates, increases in extreme climate events	land use change, habitat degradation from pollution, fertilisation, invasive species	Observed D&A: many cases of population extinctions; two cases of species extinctions (2.4.2.2); species have tracked their climate niches raising confidence in SDM projections (2.4.2.1, 2.4.2.3, 2.4.2.5)	Habitat restoration, habitat creation, increased connectivity of habitats and protected areas, increase in protected areas, assisted colonisation	0.8°C undetectable- moderate (<i>high</i> <i>confidence</i>) 1.58°C moderate- high (<i>medium</i> <i>confidence</i>) 2.07°C high-very high (<i>medium</i> <i>confidence</i>)

Wildfire: Increasing risk of wildfire that exceeds natural levels, damaging ecosystems, increasing illnesses and death of people, and increasing carbon emissions. Field evidence shows that anthropogenic climate change has increased the area burned by wildfire above natural levels across western North America in the period 1984–2017, increasing burned area up to 11 times in one extreme year and doubling burned area over natural levels in a 32-year period. Burned area has increased in the Amazon, the Arctic, Australia, and parts of Africa and Asia, consistent with, but not formally attributed to anthropogenic climate change. These changes have occurred at global mean surface temperature increases of 0.6-0.9°C. Empirical and dynamic global vegetation models project increases in burned area and fire frequency above natural levels on all continents under continued climate change, emergence of an anthropogenic signal from natural variation in fire weather for a third of global area, and increases of burned area in regions where fire had been rare or absent, particularly Arctic tundra and Amazon rainforest, at global temperature increases of 1.5-2.5°C. Models project up to a doubling of burned area globally and wildfire-induced conversion of up to half the area of Amazon rainforest to grassland at temperature increases of 3-4.5°C. (Sections 2.4.4.2, 2.5.3.2)

Climate factors	Non-climate factors	Detection and attribution	Adaptation	Risk transitions (confidence)
Increase in magnitude and duration of high temperatures, decrease in precipitation, decrease in relative humidity	Deforestation, agricultural burning, peatland burning	Increased burned area in western North America above natural levels	Reduce deforestation, reduce use of fire in tropical forests, use prescribed burning and allow naturally ignited fires to burn in targeted areas to reduce fuel loads, encourage settlement in non-fire-prone areas	0.75°C undetectable- moderate (<i>high</i>) 2.0 °C moderate- high (<i>medium</i>) 4.0°C high-very high (<i>medium</i>)

Tree mortality: Tree mortality that exceeds natural levels degrades habitat for plant and animal species, increases carbon emissions, and reduces water supplies for people. Anthropogenic climate change caused three cases of drought-induced tree mortality in the period 1945–2007 in western North America, the African Sahel, and North Africa in temperate and tropical ecosystems. Pest infestations and wildfire due to climate change also caused much of the tree mortality in North America. These changes occurred at global mean surface temperature increases of 0.3–0.9°C above the pre-industrial period. Models project increasingly extensive drought-induced tree mortality at continued temperature increases of 1–2°C. Models project risks of mortality of up to half of forest area in different biomes at temperature increases of 2.5–4.5°C. In Amazon rainforests, insufficient plant moisture reserves during drought increase the risk of tree mortality and, combined with increased fire from climate change and deforestation, the risk of a tipping point of massive forest dieback and a biome shift to grassland. (Sections 2.4.4.3, 2.5.2.6, 2.5.3.3, 2.5.3.5)

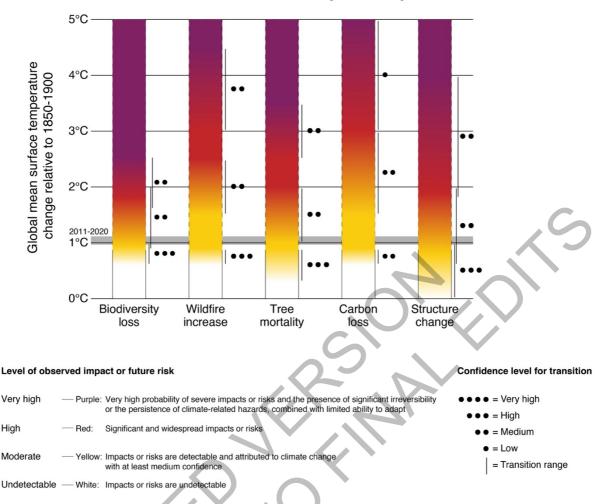
Climate factors	Non-climate factors	Detection and attribution	Adaptation	Risk transitions (confidence)
Increase in temperature, decrease in precipitation, increase in aridity, increase in frequency and severity of drought	Deforestation, land-use change	Tree mortality up to 20% in three regions in Africa and North America	Reduce deforestation, reduce habitat fragmentation, encourage natural regeneration, restore fragmented habitats	0.6°C undetectable- moderate (<i>high</i>) 1.5°C: moderate- high (<i>medium</i>) 3.0°C high-very high (<i>medium</i>)

Ecosystem carbon loss: Increasing risk of ecosystem carbon losses that could substantially raise the atmospheric carbon dioxide level. Measurements have detected emissions of carbon from boreal, temperate, and tropical ecosystems in places where increases in wildfire and tree mortality have been attributed to anthropogenic climate change, at global mean surface temperature increases of 0.6–0.9°C above the pre-industrial period. Many factors govern the carbon balance of ecosystems, so changes have not been attributed to climate change. Tropical forests and Arctic permafrost contain the highest ecosystem stocks of aboveground and belowground carbon, respectively. Primary tropical forests currently emit more carbon to the atmosphere than they remove due to deforestation and forest degradation. Wildfires in the Arctic are contributing to permafrost thaw and soil carbon release. An emissions scenario of 2°C increase could thaw ~15% of permafrost area and emit 20-100 Gt carbon by 2100, Under emissions scenarios of a 4°C global temperature increase, models project possible tipping points of conversion of half of Amazon rainforest to grasslands and thawing of Arctic permafrost that could release 11–200 Gt carbon that could substantially exacerbate climate change. (Sections 2.4.3, 2.4.4.3, 2.4.4.4, 2.5.2.7–10; 2.5.3.2–5; Figure 2.9; Figure 2.10; Figure 2.11; Table 2.4; Table 2.5; Table 2.S.2; Table 2.S.3; Table 2.S.4)

Climate factors	Non-climate factors	Detection and attribution	Adaptation	Risk transitions (confidence)
Increase in temperature, increase in aridity, increase in frequency and severity of drought	Deforestation, road and infrastructure expansion, agricultural expansion	Losses of carbon detected in boreal, temperate, and tropical ecosystems, due to wildfire and tree mortality, not formally attributed to climate change	Reduce deforestation, especially in tropical forests, reduce road and infrastructure expansion, especially in the Arctic, reduce use of fire to clear agricultural land, increase protected areas	0.75°C undetectable- moderate (<i>medium</i>) 2°C: moderate-high (<i>medium</i>) 4°C high-very high (<i>low</i>)

Ecosystem Structure Change: Increasing risk of large-scale changes in ecosystem structure. Ecosystem structural change with most information derived for tropical forest, boreal forest, savannas, and tundra for both observations and future projections. The transition from non-detectable to moderate is based on detected changes attributable to climate change, or interactions between changing disturbance regime, climate and rising CO₂, already observed at 0.5°C above pre-industrial, with shifts initially detected in boreal forests, tundra, and tropical grassy ecosystems. Transition from moderate to high is centred around 1.5°C based on widespread global observations (at current GSAT of 1.09°C above pre-industrial) that agree with projected future impacts with at least 10% area of key ecosystems being affected (Box 2.1). Overall confidence in projections is *medium*, based on existing observations and projections giving *high* confidence of risk for several ecosystems but because data and projections are not available for all biomes, overall confidence lowers to *medium*. Transition from high to very high occurs when more than 50% of multiple ecosystems are projected to experience shifts in structure. (Sections 2.4.2.3, 2.4.3, 2.4.5, 2.5.2, Box 2.1, Figure Box 2.1.1, Table Box 2.1.1, Table 2.S.2, Table 2.S.2, Table 2.S.3, Table 2.S.5)

Climate factors	Non-climate factors	Detection and attribution	Adaptation	Risk transitions (confidence)
Increases in average and extreme temperatures, changes in precipitation amount and timing, increased atmospheric CO ₂	Land use change, livestock grazing, deforestation, fire suppression, loss of native herbivores, food, fiber, or wood production	Individual species ranges shifts, biome shifts	Conservation of potential refugia, habitat restoration, increasing connectivity of habitats and protected areas, increase in protected areas, changes in grazing and fire management	0.5°C undetectable- moderate (<i>high</i>) 1.5°C moderate- high (<i>medium</i>) 2.5°C high-very high (<i>medium</i>)



Terrestrial and Freshwater Ecosystems Key Risks

Figure 2.11: Key risks to terrestrial and freshwater ecosystems from climate change. This IPCC chapter assesses these as the most fundamental risks of climate change to ecosystem integrity and the ecosystem services that support human well-being, based on observed impacts and future risks of: (far left) Losses of animal and plant species from different ecosystems globally with resulting declines in ecosystem integrity, functioning and resilience (Section 2.4.2.1, 2.4.2.2, 2.5.1.3.3); (middle left) wildfire exceeding natural levels (Section 2.4.4.2, 2.5.3.2); (middle) tree mortality exceeding natural levels (Chapter 2.4.4.3, 2.5.3.3), (middle right) ecosystem carbon losses that could occur abruptly and substantially raise atmospheric carbon dioxide (Sections 2.4.3.6—2.4.3.9, 2.4.4.4, 2.5.2.6—2.5.2.10, 2.5.3.4, 2.5.3.5); (far right) major changes occurring in ecosystem structure (Sections 2.4.3, Box 2.1, 2.5.2, Figure 2.9, Figure Box 2.1.1, Table Box 2.1.1, Table 2.S.5). This burning embers diagram shows impacts and risks in relation to changes in global mean surface temperature, relative to the pre-industrial period (1850–1900). Risk levels reflect current levels of adaptation and do not include more interventions that could lower risk. The compound effects of climate change combined with deforestation, agricultural expansion, urbanisation, air, water, and soil pollution, and other non-climate hazards could increase risks. Tables 2.5 and 2.S.4 provide details of the key risks and temperature levels for the risk transitions.

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18 [START FAQ2.4 HERE]

FAQ2.4: How does nature benefit human health and well-being and how does climate change affect this?

Human health and well-being are highly dependent on the "health" of nature. Nature provides material and economic services that are essential for human health and productive livelihoods, but studies also show that being in "direct contact with natural environments" has direct positive effects on well-being, health and socio-cognitive abilities. Therefore, the loss of species and biodiversity under climate change will reduce natural space, decrease biodiversity and in turn, decrease human-well-being and health worldwide.

Human health and well-being are highly dependent on the "health" of nature. Biodiversity – the variety of 1 genes, species, communities and ecosystems – provides services that are essential for human health and 2 productive livelihoods, such as breathable air, drinkable water, productive oceans and fertile soils for 3 growing food and fuels. Natural ecosystems also help store carbon and regulate climate, floods, disease, 4 pollution and water quality. The loss of species, leading to reduced biodiversity, has direct and measurable 5 negative effects on all of these essential services, and therefore, on humankind. A recent demonstration of 6 this is the decline of pollinator species, with potential negative effects on crop pollination, a fundamental 7 ecosystem function crucial for agriculture. The loss of wild relatives of the domesticated varieties humans 8 rely on for agriculture reduces the genetic variability that may be needed to support the adaptation of crops 9 to future environmental and social challenges. 10

The number of species that can be lost before negative impacts occur is not known and is likely to differ in different systems. However, in general, more diverse systems are more resilient to disturbances and able to recover from extreme events more quickly. Biodiversity loss means there are fewer connections within an ecosystem. A simpler food web with fewer interactions means less redundancy in the system, reducing the stability and ability of plants and animal communities to recover from disturbances and extreme weather events such as floods and drought.

In addition to "material" and economic services such as eco-tourism, nature also provides cultural services such as recreation, spirituality and well-being. Specifically, being in "direct contact with natural environments" (versus urban environment) has a high positive and causal impact on human well-being (e.g. mood, happiness), psychological and physical health (energy, vitality, heart rate, depression) and sociocognitive abilities (attention, memory, hyperactivity, altruism, cooperation). Therefore, the loss of species and biodiversity under climate change and urbanisation will reduce natural space, decrease biodiversity and in turn, decrease human-well-being and health worldwide.

Finally, the extent to which humans consider themselves as part of the natural world – known as humannature connectedness – has been demonstrated to be closely associated with human health and well-being. Individuals who are more connected to nature are not only happier and healthier but also tend to engage more in pro-nature behaviours, making the enhancement of human-nature connectedness worldwide a valuable win-win solution for humans and nature to face environmental challenges.

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Figure FAQ2.4.1: Positive relationship between human health and well-being and nature conservation. Nature provides essential services to humans including material and economic services (i.e. ecosystem services) as well as cultural, experiential and recreational services, which, in turn enhance human psychological and physical health, and well-being. People who are more connected to nature are not only happier and healthier but are also more likely to engage in pro-nature behaviours, making the enhancement of human-nature connectedness worldwide a valuable win-win solution for humans and nature to face environmental challenges.

[END FAQ2.4 HERE]

2.6 Climate Change Adaptation for Terrestrial and Freshwater Ecosystems

Adaptation to reduce vulnerability of ecosystems and their services to climate change has been addressed in 13 previous IPCC Reports, with AR4 and AR5 recognising both autonomous adaptation and human assisted 14 adaptation to protect natural species and ecosystems. In AR5, Ecosystem-based Adaptation (EbA), 15 adaptation for people, based on better protection, restoration and management of the natural environment, 16 was identified as an area of emerging opportunity, with a dedicated Cross Chapter Box on the topic. In the 17 SRCCL report, conservation, EbA and related concepts were integrated throughout the report; SR1.5 also 18 noted the role of EbA. Since the last assessment report the scientific literature has expanded considerably, 19 with growing interest in the concept of Nature-based Solutions (NbS). This section assesses this new 20 literature and its implications for the implementation of climate change adaptation. 21

22 Previous sections of this chapter have set out the vulnerability of natural and semi-natural ecosystems to 23 climate change and the risks this poses both to biodiversity and ecosystem services (also sometimes 24 described as 'Nature's Contributions to People'). Natural systems respond to climatic and other 25 environmental changes in variety of ways. Individual organisms can respond through growth, movement 26 and developmental processes. Species and populations genetically adapt to changing conditions and evolve 27 over successive generations. Geomorphological features, such as the path of watercourses, can also change 28 naturally in response to climate change. However, there is a limit to which these natural processes can 29 maintain biodiversity and the benefits people derive from nature, partly because of intrinsic limits, but also 30 because of the pressures that people exert on the natural environment. Most of this section therefore focuses 31 on human interventions to build the resilience of ecosystems, enable species to survive or to adjust 32 management to climate change. Vulnerability is in many cases exacerbated by the degraded state of many 33 ecosystems as a result of human exploitation and land use change, leading to fragmentation of habitats, loss 34 of species and impaired ecosystem function. This interaction between climate change and environmental 35 degradation means that protecting ecosystems in a natural or near-natural state will be an important pre-36 requisite for maintaining resilience and give many species the best chance of persisting in a changed climate 37 (Belote et al., 2017; Arneth et al., 2020; Ferrier et al., 2020; França et al., 2020). Protection from 38 degradation, deforestation and exploitation is also essential to maintain critical ecosystem services, including 39 carbon storage and sequestration and water supply (Dinerstein et al., 2020; Pörtner, 2021). 40

41 It is worth briefly considering some key concepts that are relevant to adaptation in ecosystems. Adaptation 42 for biodiversity and ecosystems can encompass both managing change and building resilience. We use the 43 definition of 'resilience' set out Chapter 1: 'the capacity of social, economic and environmental systems to 44 cope with a hazardous event or trend or disturbance, responding or reorganising in ways that maintain their 45 essential function, identity and structure while also maintaining the capacity for adaptation, learning and 46 transformation', It includes the concept of 'resistance', which is used in some ecological literature to 47 distinguish systems which a resistant to change from those that recover quickly from change. We consider 48 both interventions designed primarily to protect biodiversity and those intended to reduce the risks of climate 49 change to people. 50

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A variety of terms are used to describe using environmental management reduce the impacts of climate change on people in ways that also benefit biodiversity in the scientific literature, particularly Ecosystem based Adaptation (EbA) and Nature-based Solutions (NbS) (see also Section 1.4). EbA is the use of biodiversity and ecosystem services as part of an overall adaptation strategy to help people to adapt to climate change (CBD, 2009). EbA aims to maintain and increase the resilience and reduce the vulnerability of ecosystems and people in the face of the adverse effects of climate change (Vignola et al., 2009). NbS is a broader term which is not restricted to climate change and is also often used to refer to climate change

mitigation; it has been defined by the International Union for the Conservation of Nature (IUCN) as 'Actions 1 to protect, sustainably manage and restore natural or modified ecosystems that address societal challenges 2 effectively and adaptively, simultaneously providing human well-being and biodiversity benefits' (Cohen-3 Shacham et al., 2016). This widely accepted definition excludes actions, which use the natural environment 4 to solve human problems but do not provide benefits for biodiversity and is closely linked to the concept of 5 the Ecosystem Approach. NbS is not a universally accepted term but it is increasingly used in the scientific 6 literature. It is a concept which recognises the importance of biodiversity in ecosystem service provision and 7 offers the opportunity to address climate change and loss of biodiversity together in an efficient integrated 8 way (Chong, 2014; Seddon et al., 2020a; Ortiz et al., 2021). Given the focus of this chapter is on adaptation 9 we primary use the term EbA as it is more specific, but we do so understanding that it can be regarded as a 10 subset of NbS. The wider concept of NbS for climate change adaptation and mitigation is covered in a Cross-11 Chapter Box on the topic (see Cross-Chapter Box NATURAL this Chapter). 12 13

Whilst we distinguish between adaptation for biodiversity and EbA, it is important to recognise that the two are linked in that if ecosystems themselves are not resilient to climate change, they will not be able to provide adaptation benefits for people. The case for resourcing biodiversity conservation and building the resilience of ecosystems is also strengthened when there are direct benefits for people in addition to the more general benefits of biodiversity.

Ecosystems are specifically included in the adaptation goals set out in the Paris Agreement and are addressed in most national adaptation plans (Seddon et al., 2020b). There is also now a large number of adaptation programmes and plans for local governments and governmental and non-governmental conservation organisations. Adaptation for and by ecosystems needs to be understood and developed in the wider contexts of conservation, Climate Resilient Development and Sustainable Development: there is significant potential synergies, but also conflicts between different objectives, which require an integrated approach (covered further in 2.6.7).

28 2.6.1 Limits to Autonomous (Natural) Adaptation

Natural ecosystems often have a high degree of resilience and can to some extent adjust to change. Species
can adjust through evolutionary adaptation, distribution change, behavioural change, developmental
plasticity and ecophysiological adjustment. There are, however, limits to autonomous adaptation, because of
intrinsic limitations, the rate at which the climate is changing and the degraded state of many ecosystems.

34 None of the evolutionary changes either documented or theorised would enable a species to survive and 35 reproduce in climate spaces that it does not already inhabit. Evolutionary responses are very unlikely to 36 prevent species extinctions in the case of that species losing its climate space entirely on a regional or global 37 scale (Parmesan and Hanley, 2015). At highest risk are the world's most cold-adapted species (whose 38 habitats are restricted to polar and high mountaintop areas). Examples include the polar bear (Regehr et al., 39 2016), "sky-island" plants in the tropics (Kidane et al., 2019), mountain-top amphibians in Spain (Enriquez-40 Urzelai et al., 2019), mountain-top lichens in the Appalachians (USA) (Allen and Lendemer, 2016), and 41 silverswords in Hawaii (Krushelnycky et al., 2013). However, there is potential for using evolutionary 42 changes to enhance the adaptive capacity of target species, as is being done in the Great Barrier reef by 43 translocating symbionts and corals that have survived recent intense heat-induced bleaching events into areas 44 that have had large die-off (Rinkevich, 2019). Multiple studies assessed when and how evolution might be 45 able to help wild species adapt to climate change (Ratnam et al., 2011; Sgro et al., 2011). 46

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48 Some of the reasons cited in the literature as limits to autonomous adaptation are:

1) Genetic changes in populations require many generations and for many species operate on longer time
 scales than those, on which the climate is currently changing.

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2) Many species are moving to higher latitudes as the climate warms, but not all are keeping pace with
 changes in suitable climate space (Valladares et al., 2014; Mason et al., 2015). Such climate debt indicates
 an inability for non-genetic autonomous adaptation (e.g. evidence-limited ability for plastic responses, such
 as stemming from dispersal limitations, or behavioural restrictions, or physiological constraints).

3) Some species have low capacity for dispersal, which, combined with increased fragmentation of habitats, creates barriers to range shifts to match climate warming. Studies have shown that changes in distribution of species and composition of communities are limited by the presence of intensively managed agricultural land fragmenting natural habitats (Oliver et al., 2017).

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There are a variety of mechanisms which promote the resilience of ecosystems through persistence, 6 recovery, and reorganisation (Falk et al., 2019). Changes in the balance of different plant species within a 7 community can maintain the persistence of the community itself, maintaining its value as a habitat for other 8 species and providing ecosystem services (add reference?). In some cases there are negative feedback 9 mechanisms between biological and physical processes, for example in peatlands, lowered water tables 10 resulting from drier conditions can lead to reduced permeability of peat, increasing rates of water loss (Page 11 and Baird, 2016). There are limits to this resilience and the concept of tipping points beyond which 12 ecosystems change state and returning to the original state has been subject of much recent research (van Nes 13 et al., 2016). There is clear evidence that the degradation of ecosystems has reduced their resilience and 14 restoration can help to reduce risks to biodiversity and ecosystem services, discussed below (see Section 15 2.6.2, 2.6.3). However, as rates of climate change increase, the limits of this approach will start to be 16 reached and losses, including some with potentially catastrophic consequences, cannot be prevented; this is 17 discussed further in Section 2.6.6. 18

20 2.6.2 Adaptation for Biodiversity Conservation

21 A variety of approaches have been identified as potential adaptation measures which people can take to 22 reduce the risks of climate change to biodiversity. (Heller and Zavaleta, 2009) (quoted in AR5) identified 23 113 categories of recommendation for adaptation from a survey of 112 papers and reports. Since this time 24 the literature has greatly expanded, with thousands of relevant publications. Whilst there is increasing 25 interest in adaptation for biodiversity conservation and a wide range of plans and strategies, there is less 26 evidence of these plans being implemented. Since AR5 a number of studies, predominantly from Europe and 27 North America, have investigated the extent to which adaptation has been integrated into conservation 28 planning and is being implemented at site and regional scale (Macgregor and van Dijk, 2014; Delach et al., 29 2019; Prober et al., 2019; Clifford et al., 2020; Barr et al., 2021; Duffield et al., 2021). A common pattern in 30 these studies is that vulnerability has been assessed and potential adaptation actions identified, but 31 implementation has been limited beyond actions to improve ecological condition, which may increase 32 resilience at a local scale. 33

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To date most scientific literature on adaptation to reduce risk to biodiversity from climate change has been based on ecological theory rather than observations or practical experience. A recent review (Prober et al., 2019) concluded that out of 473 papers on adaptation, only 16% presented new empirical evidence and very few assessed the effectiveness of actual adaptation actions. It is also the case that relatively little research is focussed on local-scale management interventions rather than larger scale strategies (Ledee et al., 2021), although there are some exceptions (Duffield et al., 2021).

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Although direct assessments of the effectiveness of adaptation actions are rare, since AR5 there has been an 42 increasing number of empirical analyses of how different land use and management influences the 43 vulnerability of species and habitats. As climate change often interacts with other factors including 44 ecosystem degradation and fragmentation (Oliver et al., 2015a), actions to address these other interacting 45 factors is expected to build resilience to climate change. Table 2.6 summarises evidence that supports the 46 main categories of proposed adaptation measures. We have taken an inclusive approach and included studies 47 that address extreme weather events such as droughts, which may be exacerbated by climate change as well 48 49 as long-term changes in climate variables. We have not distinguished between studies in which climate change adaptation was an explicit focus and those in which lessons for adaptation can be learnt from studies 50 conducted for other reasons but inform the assessment of the impacts of actions identified as potential 51 adaptation measures. 52

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Table 2.6: Evidence to support proposed climate change adaptation measures for biodiversity. Highlights that adaptation for biodiversity is a broad concept, encompassing a wide range of actions. It includes targeted interventions to change the microclimate for particular species (for example by shading); through to changing national conservation objectives to take account of changing distributions of species and communities. It includes targeted actions addressing

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both climate change and protection and restoration of ecosystems, with multiple additional benefits including reducing vulnerability to climate change. Most of the studies are not direct tests of the impacts of adaptation actions, which, as

noted above, is an important evidence gap. There is also a major limitation in that reported studies are predominantly
 from Europe, North America and Australasia, with little research in other regions.

Proposed Adaptation	Confidence	Comment	Selected References
Measures for Biodiversity	Assessment		
Protect large areas of natural and semi- natural habitat	Robust evidence, high agreement	Considerable evidence that intact systems provide better quality and quantity of ecosystem services; that larger intact areas provide better ecosystem services; that species' extinction risk with disturbances, including climate change, are reduced by having large, connected populations; that more biodiverse systems provide higher levels of ecosystem services and are more resilient to climate change than degraded systems that have lost species	Dinerstein et al. (2019); Woodley et al. (2019); Brooks et al. (2020); Zhao et al. (2020); Sala et al. (2021); Pimm et al. (2018); Hannah et al. (2020); Luther et al. (2020)
Increase connectivity in terrestrial habitats – corridors, stepping stones	Medium evidence, Medium agreement	Good evidence that some species move more quickly in more connected landscapes. However, not all species do and some species that benefit are invasive / pest / disease species and there is limited empirical evidence showing connectivity has reduced climate change impacts on species to date.	Keeley et al. (2018); Stralberg et al. (2019); von Holle et al. (2020)
Increase connectivity in river networks	Limited evidence, High agreement	Connectivity is needed to maintain species and population movements, but river reaches and catchments lack integrated protection.	Hermoso et al. (2016); Thieme et al. (2016); Abell et al. (2017); Brooks et al. (2018)
Increase habitat patch size site and expand protected areas	Limited evidence High agreement	Generally increase resilience because of functioning natural processes, large species populations and refugial areas	Eigenbrod et al. (2015); Oliver et al. (2015a)
Increase replication and representation of protected areas	Limited evidence, High agreement	Various benefits inferred, including, wider range of climatic and other conditions, less risk of extreme events affecting many rather than few areas. More sites available for colonisation by range expanding species and better conditions to maintain species in situ under range contraction.	Mawdsley et al. (2009); Thomas et al. (2012a); Virkkala et al. (2014); Gillingham et al. (2015); Pavón-Jordán et al. (2020b)
Protect microclimatic refugia	Medium evidence, High agreement	Locally cool areas can be identified and there is evidence species can survive better in such areas.	Haslem et al. (2015); Suggitt et al. (2015); Isaak et al. (2016); Morelli et al. (2016); Merriam et al. (2017); Bramer et al. (2018); Suggitt et al. (2018); Massimino et al. (2020)
Creating shade to lower temperatures for vulnerable species	Limited evidence, High agreement	Creating shade has been used as an adaptation strategy, for example by watercourses but improvements in species survival under warming conditions have yet to be demonstrated.	Broadmeadow et al. (2011); Lagarde et al. (2012); Patino-Martinez et al. (2012); Thomas et al. (2016)
Restoring hydrological processes of wetlands, rivers and catchments, including by raising water tables and restoring original channels of watercourses,	Medium evidence, High agreement	Wetland restoration is well established as a conservation measure in some countries. Can reduce vulnerability to drought with climate change but evidence to demonstrate effectiveness as an adaptation measure is limited and requires long-term monitoring of a range of sites. Little restoration of degraded tropical peatlands to date	Carroll et al. (2011); Hossack et al. (2013); Dokulil (2016); Timpane-Padgham et al. (2017); Moomaw et al. (2018)
Restoration of natural vegetation dynamics	Medium evidence,	Includes reintroduction of native herbivores and reversing woody encroachment of	Coffman et al. (2014); Valkó et al. (2014);

	Medium agreement	savannas. Benefits for biodiversity are well established in a wide range of different regions	Batáry et al. (2015); Smit et al. (2016); Stevens et al. (2016); Hempson et al. (2017); Bakker and Svenning (2018); Cromsigt et al. (2018); Fulbright et al. (2018); Olofsson and Post (2018)
Reduce non-climatic stressors to increase resilience of ecosystems	Limited evidence, Medium agreement	As a general principal climate change is recognised as a 'threat multiplier' but specific details are often unclear	Oliver et al. (2017); Pearce-Higgins et al. (2019)
Assisted translocation and migration of species	Limited evidence, Medium agreement	Assisted translocation has been commonly suggested as an adaptation measure, but there have been very few examples of this being trialed. Translocations have been carried out for other reasons and lessons for climate change adaptation have been inferred	Willis et al. (2009); Brooker et al. (2018); Skikne et al. (2020)
Intensive management for specific species	Medium Evidence, Medium Agreement	A variety of approaches including manipulating microclimate and competition between species to improve chances of survival under climate change.	Angerbjörn et al. (2013); Greenwood et al. (2016); Pearce-Higgins et al. (2019)
Ex-situ conservation (seedbanks/genetic stores, etc.)	Not possible to assess at present time	Seed banks have been established but long- term effectiveness could only be evaluated at a later point.	Christmas et al. (2016)
Adjusting conservation strategies and site objectives to reflect changing species distributions and habitat characteristics	Robust evidence, High Agreement	Conservation management will need to take account of changes that cannot be prevented, for example in the distribution of species and composition of communities, in order to protect and manage biodiversity as effectively as possible in a changing climate.	Stein et al. (2013); Rannow et al. (2014); Oliver et al. (2016); Stralberg et al. (2019); Duffield et al. (2021)
Softening the matrix of unsuitable habitats between patches to increase permeability for species movement in response to climate change	Limited evidence	Potential for agri-environment schemes to do this in hostile farmed landscapes.	Donald and Evans (2006); Stouffer et al. (2011)

Many climate adaptation actions for biodiversity operate at the landscape scale (von Holle et al., 2020). The 3 total area of habitat, how fragmented it is, the size of habitat patches and the connectivity between them are 4 interlinked properties at this scale. A growing number of studies have investigated how these properties 5 affect species ability to persist in situ and colonise new areas. Overall, larger areas of semi-natural habitat are 6 associated with increased resilience to ongoing climate change and extreme events and the capacity to 7 colonise new areas (Haslem et al., 2015; Oliver et al., 2017; Papanikolaou et al., 2017). Larger habitat 8 patches can support larger populations, which are more likely to maintain themselves and recover from 9 periods of adverse conditions. A large patch size has been found to increase resilience of some populations 10 of species to extreme events such as droughts (Oliver et al., 2015b). They are also more likely to provide a 11 range of different resources and microclimate conditions, which may increase chances of species persistence 12 under climate change. A larger area of habitat may also enables greater connectivity between patches and 13 increases the chances of species colonising new areas as they track climate change. (Oliver et al., 14 2015b)Protecting and restoring natural processes is a general principle for maintain and building resilience to 15 climate change for biodiversity (Timpane-Padgham et al., 2017). One element of this is ensuring naturally 16 functioning hydrology for wetlands and river systems (Table 6), which is particularly important in a context 17 of changing rainfall patterns and increased evapo-transpiration. An important development in approaches to 18

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conservation over recent decades has been the concept of re-wilding (Schulte To Bühne et al., 2021); this encompasses a number of elements of restoring natural processes, including the reintroduction of top predators, larger conservation areas and less prescriptive outcomes than much previous conservation. There are elements of re-wilding which may well contribute to building resilience to elimete change but it will be

are elements of re-wilding which may well contribute to building resilience to climate change but it will be
 increasingly important to factor climate change adaptation into the planning of rewilding schemes (Carroll
 and Noss, 2021).

The most consistently cited climate change adaptation measure for species is increasing connectivity to 8 facilitate colonisation of new areas. This reflects the fact that many species' habitats are highly fragmented 9 in areas with more intensive land management, which prevents them naturally changing their range to track 10 changing climatic conditions. Advances and innovations in modelling techniques can support decision 11 making on connectivity (Littlefield et al., 2019). There is evidence from empirical as well as modelling 12 studies that species can disperse more effectively in better connected areas in terrestrial habitats (Keeley et 13 al., 2018). The issues are different in more natural landscapes—species may still be threatened in 14 intrinsically isolated habitats, such as mountain top, but connectivity cannot be created in the same way. 15 Evidence suggests that increased connectivity will only benefit a subset of species, probably those which are 16 intermediate habitat specialists that are able to disperse (Pearce-Higgins et al., 2014). Generalists do not 17 require corridors or stepping stones whilst many corridors or stepping stones will not be of sufficient quality 18to be used by the most habitat specialists. There should also be a caveat to the general principle that 19 increasing connectivity is a benefit for climate change adaptation. It can increase the spread of invasive, pest 20 and disease-causing species into newly suitable regions. In some places isolated refugia may better allow 21 vulnerable species and biological communities to survive. 22

23 There are many different approaches to increasing connectivity, ranging from increasing overall area of 24 suitable habitat through to 'corridors' and 'stepping stones', with different strategies likely to be more 25 effective for different species and circumstances (Keeley et al., 2018). Connectivity can also be important in 26 increasing resilience of populations to extreme climatic events (Newson et al., 2014; Oliver et al., 2015b). 27 Within freshwater environments, connectivity of watercourses is essential. Fluvial corridors are necessary to 28 ensure migrating fish population survival, even without climate change; with climate change, connectivity 29 becomes crucial for relatively cold-adapted organisms to migrate upstream to colder areas. Connectivity is 30 also important for the larvae of benthic invertebrates to be able to drift downstream and hence to disperse 31 (Brooks et al., 2018); for adult benthic invertebrates, riparian and terrestrial habitat features can potentially 32 affect dispersal. Connectivity within river and wetland systems for some species can also mediated by more 33 mobile animal species such as fish and birds (Martín-Vélez et al., 2020) Which factors are the most 34 important in either promoting their colonisation of new sites or persisting *in situ* will differ between species 35 and locations. Some general principle have been recognised and can guide conservation policy and practice 36 (England and RSPB, 2020; Stralberg et al., 2020) but this will often require additional investigation and 37 planning based on understanding individual the niche of specific species. 38

Managed, translocation by moving species from areas where the climate is becoming unsuitable to places where there persistence under climate change is more likely has been discussed as an adaptation option for many years. So far there have been very few examples of this and it is likely to be a last resort in most cases as in many cases it requires a large investment of resources, the outcome is uncertain and there may be adverse impacts on receiving sites. Nevertheless there are cases where it may be a be a viable option (Stralberg et al., 2019). This is discussed in more detail as a case study in section 2.6.5.1.

46 The evidence that species can persist in microclimatic refugia where suitable conditions for them are 47 maintained locally (for example, because of variations in topography) has increased in recent years. This has 48 49 opened up the potential to include refugia in conservation plans and strategies to facilitate local survival of species (Jones et al., 2016; Morelli et al., 2016; Morelli et al., 2020). For example, in targeting management 50 actions (Sweet et al., 2019) aimed at supporting populations of species and is likely to become an important 51 aspect of climate change adaptation for biodiversity conservation in future. It is also possible to manipulate 52 microclimate for example by creating shelters for birds' nests (see case study of African Penguins 2.6.5.5) 53 (Patino-Martinez et al., 2012). One specific approach of this sort is planting or retention of trees and wooded 54 corridors to shade water courses (see also case study 2.6.4.5 below; Thomas et al., 2016). In the latter case, 55 riparian shading can also possibly help to reduce phytoplankton and benthic diatom growth in smaller 56

streams and rivers (Halliday et al., 2016).

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1 Refugia often refer to locally places in a landscape, such as on shaded slopes or high elevations, but they can 2 also include places where water supply may continue during dry periods (Morelli et al., 2016) Monitoring 3 can reveal which streams, wetlands, springs, and other aquatic resources retain suitable discharges, water 4 quality, wetland area, and ecological integrity, especially during dry years (Cartwright et al., 2020). 5 Measures to conserve drought refugia may include protecting springs and other groundwater-fed systems 6 from groundwater extraction, contamination, salinisation, surface-water diversion, channelisation of 7 streams, livestock trampling, recreation and invasive species, as well as effects from surrounding landscape 8 disturbances (Cartwright et al., 2020; Krawchuk et al., 2020). Restoration of degraded aquatic ecosystems 9 can include removal of flow-diversion infrastructure, exclusion of livestock, reduction of other human 10 impacts, geomorphic restructuring, invasive species removal, and planting of native riparian vegetation. 11 12 In fire prone areas, fire suppression and management is a key element of protecting refugia (Section 2.6.5.8 13 below). In ecosystems in which a natural fire regime has been suppressed restoration practices such as 14 prescribed fires, thinning trees, and allowing some wildfires where it benefits the ecosystem can be 15 introduced to reduce increasing risks from severe wildfires (Meigs et al., 2020). 16 17 Protected areas-areas of land set aside for species and habitat protection with legal protection from 18development or exploitation-have been a cornerstone of nature conservation for many years. Their 19 effectiveness under a changing climate has been the subject of debate and investigation. There is now a large 20 body of evidence demonstrating that colonisations by range shifting species are more likely to occur on 21 protected sites compared to non-protected sites for a wide range of taxa (e.g. Thomas et al., 2012b; 22 Gillingham et al., 2015), including across continents (Pavón-Jordán et al., 2020a). This is probably because 23 by protecting large areas of natural and semi-natural habitats they provide suitable places for colonising 24 species (Hiley et al., 2013) which may not be available in the surrounding landscape. Although the evidence 25 for protected areas being associated with reduced extinctions is weaker, the finding in Gillingham et al. 26 (2015) that protected sites were associated with reduced extinction rates at low latitudes and elevations is 27 strongly suggestive that they can help species' persistence in the face of climate change. 28 29 It is intrinsically difficult to assess the effectiveness of climate change adaptation measures, the benefit of 30 which will be realised in years and decades ahead (Morecroft, 2019, Measuring success). Nevertheless, 31 taking account of the wide range of evidence reported above, including theory, modelling and observations 32 of climate change impacts in contrasting circumstances, it is possible to make an overarching assessment that 33 appropriate adaptation measures can reduce the vulnerability of many aspects of biodiversity to climate 34 change (robust evidence, high agreement). It is also however clear that to be most effective and avoid 35 unintended consequences, measures need to be carefully implemented taking account of specific local 36 circumstances (robust evidence, high agreement) and include the management of inevitable changes (robust 37

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agreement).

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41 42 2.6.3 Nature-based Solutions: Ecosystem-based Adaptation

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Ecosystem-based Adaptation is an increasingly important element of Nature-based Solutions (see 2.6 above).
A study published in 2020 found that out of 162 Intended Nationally Determined Contributions (covering
189 countries) submitted to the United Nations Framework Convention on Climate Change, as commitments
to action under the Paris Agreement, 109 indicated 'ecosystem-orientated visions' for adaptation, although
only 23 use the term 'Ecosystem-based Adaptation' (Seddon et al., 2020b).

evidence, high agreement). It is also clear that whilst there are now many plans and strategies for adapting

biodiversity conservation to climate change, many have yet to be implemented fully (medium evidence, high

EbA includes a range of different approaches. Examples include restoring coastal and river systems to 50 reduce flood risk and improve water quality and the creation of natural areas within urban areas to reduce 51 temperatures through shading and evaporative cooling. EbA is closely linked with a variety of other concepts 52 such as ecosystem services, natural capital and Disaster Risk Reduction (DRR). EbA was becoming a well-53 recognised concept at the time of AR5 but implementation was still at an early stage in many cases. Since 54 then pilot studies have been assessed and EbA projects have been initiated around the world. The evidence 55 base continues to grow (Table 2.7) and this has led to increasing confidence in approaches which have been 56 shown to work leading to further expansion in some countries (Table 2.7). However, this is not uniform and 57

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1	there is relatively little synthesis a	cross disciplines and regions (Sedd	lon et al., 2020a). Chausson et al. (2020)
1	• •	1 0 (
2	used a systematic mapping metho	dology to characterise 386 publishe	ed studies. They found that interventions
3	in natural or semi-natural ecosyste	ems ameliorated adverse climate ch	ange impacts in 66% of cases, with
4	fewer trade-offs than for more art	ificial systems such as plantation fo	prest. However, the evidence base has

substantial gaps. Most of the evidence has been collected in the Global North and there is a lack of robust, 5 site-specific investigations of the effectiveness of interventions compared to alternatives and of more holistic 6

appraisals accounting for broader social and ecological outcomes.

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Table 2.7: Examples of key Ecosystem- based Adaptation measures with assessments of confidence. Note only adaptation related services are shown - many measures also provide a range of other benefits to people. All also provide benefits for biodiversity.

Ecosystem Based Adaptation Measures	Confidence Assessment	Ecosystem Services for climate change adaptation	Climate Change Impact addressed	Social Benefits from adaptation	Relevant Ecosystems and contexts	Selected References
Natural Flood risk management in river systems –restoring natural river courses (removing canalisation), restoring and protecting wetlands and riparian vegetation	Medium evidence Medium agreement	Flood regulation; sediment retention; water storage; water purification	Increased rainfall intensity	Reduction of flood damage Increased water security (quality and supply)	Multiple	Iacob et al. (2014); Meli et al. (2014); Burgess-Gamble et al. (2017); Dadson et al. (2017); Rowiński et al. (2018)
Shade rivers and streams by restoration of riparian vegetation or trees.	Medium evidence High agreement	Provision of fish stocks	Warmer water temperatures	Food security income benefits	Multiple	Broadmeadow et al. (2011); Isaak et al. (2015); Williams et al. (2015b); Thomas et al. (2016)
Managed realignment of coastlines; re- establishing and protecting coastal habitats including mangroves, salt marsh,	Robust evidence High agreement	Coastal storm and flood protection Coastal erosion control Salt water intrusion prevention	Rising sea level Increasing storm energy	Protection of life, property and livelihoods Water security	Coastal	Høye et al. (2013); Spalding et al. (2014); Narayan et al. (2016); Morris et al. (2018); Chowdhury et al. (2019); Powell et al. (2019)
Agroforestry and other agro- ecological/conservation agricultural practices on agricultural land	Medium evidence Medium agreement	Local climate regulation; soil conservation; soil nutrient regulation; water conservation; pest control; food provisioning	High temperature or changing temperature regimes Changing precipitation regimes	Food security income benefits	Multiple	Vignola et al. (2015); Torralba et al. (2016); Paul et al. (2017); Blaser et al. (2018); Nesper et al. (2019); Verburg et al. (2019); Aguilera et al. (2020); Tamburini et al. (2020)
Restore and maintain urban and peri-urban green space – trees, parks, local nature reserves, created wetlands	Robust evidence High agreement	Local climate regulation Flood regulation Water purification Water storage	Higher temperatures and heatwaves Increased rainfall intensity or	Cooler microclimate Reduced flood damage, water security	Urban areas	Norton et al. (2015); Liquete et al. (2016); Liu (2016); Bowler et al. (2017); Aram et al. (2019);

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		Erosion control	reduced rainfall intensity			Stefanakis (2019); Ziter et al. (2019)
Ecological restoration for fire risk reduction through restoration of natural vegetation and herbivory and by re- instating natural fire regimes	Medium evidence High agreement	Regulation of wildfires	Mega-fires from increases in drought and heat	Reduce deaths and infrastructure damage from fires	Fire- adapted ecosystems	Waldram et al. (2008); Stephens et al. (2010); van Mantgem et al. (2016); Boisramé et al. (2017); Johnson et al. (2018); Parisien et al. (2020a); Parisien et al. (2020b); Stephens et al. (2020)
Invasive non-native aquatic plant control to improve water security	Robust evidence High agreement	Water provision	Increasing droughts	water security	Water scarce regions prone to an increase in droughts	van Wilgen and Wannenburgh (2016)
Woody plant control (of encroaching biomass) in open grassy ecosystems to restore and maintain grassy vegetation (see 2.4.3.5)	Medium evidence Medium agreement	Fodder biomass production	Elevated CO ₂ increasing tree growth/ increases in rainfall promoting tree growth	income through bush clearing, fuelwood supplies, restore grazing	Savanna and grasslands	Haussmann et al. (2016)
Rangeland rehabilitation and management such as through livestock enclosures, appropriate grazing management, re-introducing native grassland species	Medium evidence Medium agreement	Fodder biomass production; soil erosion control; soil formation; nutrient cycling; water retention	Changing precipitation and temperature regimes including prolonged dry seasons and increased drought frequency	Food security Water security, income benefits	Rangelands	Descheemaeker et al. (2010); Wairore et al. (2016); Kimiti et al. (2017)
Sustainable forestry of biodiverse managed forests, maintaining forest cover and protecting soils	Medium evidence Medium agreement	Timber production	Increased frequency and severity of storms Higher temperatures Changing precipitation regimes (more intensive wet and dry periods) Increased incidents of wildfire, pest, and disease outbreaks	Livelihood and income benefits	Boreal, temperate, subtropical, tropical forests	Gyenge et al. (2011); Barsoum et al. (2016); Jactel et al. (2017); Cabon et al. (2018)

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Watershed reforestation and conservation for hydrological services	Medium evidence Medium agreement	Flood control; erosion control; water provisioning; water purification	precipitation	Food security; Water security; Flood Protection	Boreal, temperate, subtropical, tropical forests	Filoso et al. (2017); Bonnesoeur et al. (2019)
Multifunctional forest management and conservation to provide climate resilient sources of food and livelihoods and protect water sources	Medium evidence Medium agreement	Timber and non-timber forest production; fuel wood production; water provisioning; water purification	Multiple	Food security; Water security; income benefits	Boreal, temperate, subtropical, tropical forests	Lunga and Musarurwa (2016); Strauch et al. (2016); Adhikari et al. (2018)
Slope revegetation for landslide prevention and erosion control	Robust evidence High agreement	Soil retention; slope stabilisation	Increased rain frequency	Reduced landslide damage; prevention of loss of life	Montane and other steep sloped regions	Fox et al. (2011a); Krautzer et al. (2011); Leal Filho et al. (2013); Bedelian and Ogutu (2017); Getzner et al. (2017); de Jesús Arce-Mojica et al. (2019)

2 Restoring coasts, rivers and wetlands to reduce flood risk have probably seen the largest investment in EbA 3 and it is becoming an increasingly accepted approach in some places (e.g. case studies in Sections 2.6.5.2, 4 2.6.5.7) although significant social, economic and technical barriers remain (Wells et al., 2020; Bark et al., 5 2021; Hagedoorn et al., 2021). Natural flood management encompasses a wide range of techniques in river 6 systems and at the coast and have been used in varied locations around the world. In tropical and sub-7 tropical areas, the restoration of mangroves to reduce the risk of coastal flooding is widely advocated, 8 evidence-based approach (for example (Høye et al., 2013; Sierra-Correa and Kintz, 2015; Powell et al., 9 2019)). In temperate regions salt marsh is a similarly important habitat (Spalding et al., 2014). Both provide 10 buffering against rising sea levels and storm surges. Managed realignment of the coast, by creating new 11habitats can lead to a loss of terrestrial and freshwater ecosystems, but it can protect them and the services 12 they provide by reducing the risks of catastrophic failure from hard-engineered sea defences. In river 13 systems (Iacob et al., 2014) management of both catchments and the channel itself is important: restoring 14 natural meanders in canalised water courses and allowing the build-up of woody debris can slow flows rates; 15 restoring upstream wetlands or creating them in urban and peri-urban situations can store water during flood 16 events if they are in the right place in a catchment (Acreman and Holden, 2013; Ameli and Creed, 2019; Wu 17 et al., 2020). There is less data on the potential for natural flood management in tropical compared to 18 temperate catchments, however (Ogden et al., 2013) showed that flooding was reduced from a secondary 19 forested catchment compared to those which were pasture or a mosaic of forest, pasture and subsistence 20 agriculture. EbA approaches to reduce flooding can be applied within urban areas, as well as in rural 21 catchments, as in Durban (Section 2.6.5.7), although its effectiveness will depend on its being implemented 22 at a sufficient scale and in the right locations (Hobbie and Grimm, 2020; Costa et al., 2021). which may in 23 turn provide protection to downstream urban communities. 24 25

Protecting and restoring natural river systems, natural vegetation cover within catchments and integrating agro-ecological techniques into agricultural systems can also help to maintain and manage water supplies for human use, under climate change, including during drought periods, by storing water in catchments and improving water quality (Taffarello et al., 2018; Agol et al., 2021; Khaniya et al., 2021) (Lara et al., 2021) showed that replacing a non-native *Eucalyptus* plantation in Chile with native forest caused base flow increased by 28% to 87% during the restoration period compared to pre-treatment, and found it remained during periods with low summer precipitation FINAL DRAFT

1 EbA can operate at a range of different scales, from local to catchment to region. At the local scale, there is 2 a variety of circumstances in which microclimates can be managed and local temperatures lowered by the 3 presence of vegetation (Table 2.7), and these EbA techniques are now being used more widely. In both urban 4 and agricultural situations, shade trees are a traditional technique, which can be applied to contemporary 5 climate change adaptation. Shading of water courses can lower temperatures, as reported in Section 2.6.2, 6 above, which can allow species to survive locally, as well as supporting diversity it can help to maintain are 7 important fisheries, including of Salmonid fish (O'Briain et al., 2020). Within cities, green spaces, including 8 parks, local nature reserves and green roofs and walls can also provide cooling as a result of evapo-9 transpiration (Bowler et al., 2010a; Aram et al., 2019; Hobbie and Grimm, 2020), although this may be 10 reduced in drought conditions. 11 12 Wildfire is an increasing risk for people as well as to ecosystems, in many parts of the world. As discussed 13 in Section 2.4.4.2, this is the result not just of climate change but also past management practices, including 14 fire suppression. Better fire management including reinstating more natural fire regimes can reduce risks. 15 16 EbA is usually a place-specific approach and a number of studies have documented how attempts to 17 implement it without an understanding of local circumstances and full engagement of local communities 18have been unsuccessful (Nalau et al., 2018). Since AR5, a number of studies have considered the factors that 19 are important for environment adaptation programmes and projects (UNFCCC, 2015; Nalau et al., 2018; 20 Duncan et al., 2020; Network and ENCA, 2020; Townsend et al., 2020). Considering these sources, others 21 described above and the case studies presented, in 2.6.5, a number of requirements for effective 22 implementation of EbA can be identified, including the following: 23 24 • Targeting of the right EbA measure in the right location 25 • Decision-making at the appropriate level of governance with participation from all affected communities 26 • Integration of Local Knowledge and Indigenous Knowledge & capacity into decision-making and 27 management of project 28 • Involvement of government and non-government stakeholders 29 • Full integration of EbA with other policy areas, including agriculture, water resources and natural 30 resource protection 31 • Protection and if possible improvement of incomes of local people. 32 • Effective institutional support to manage finances and implementation of projects and programmes. 33 • Time -many EbA interventions take time to establish e.g. trees to grow, wetlands recover 34 • Monitoring of intended outcomes and other impacts and communication of results 35 36 Whilst it is essential to develop place-specific EbA measures, with full engagement of local communities, it 37 is worth noting that new opportunities may emerge that would not have been possible in the past. As the 38 climate changes, novel ecosystems may emerge with no present day analogue which have the potential to 39 provide different adaptation benefits and societies may be more willing to adopt transformative approaches 40 (Colloff et al., 2017; Lavorel et al., 2020). 41 42 Increasingly it is essential to integrate adaptation and the protection of biodiversity with land based climate 43 change mitigation initiatives; this is discussed in more detail in Cross-Chapter Box NATURAL, this 44 Chapter. The new IUCN standard (IUCN, 2020, Global standard) offers a basis for assessing whether actions 45 are true Nature-based Solutions and take account of the wider factors necessary for success. 46 47 Whilst policy interest is growing and there is an increasing deployment of EbA there is still a long way to go 48 in delivering it full potential (Huq and Stubbings, 2015) and significant institutional and cultural barriers 49 remain(Huq et al., 2017; Nalau et al., 2018). Nevertheless it is increasingly clear that EbA can offer a 50 portfolio of effective measures to reduce risks from climate change to people at the same time as benefiting 51 biodiversity (robust evidence, high agreement), providing they are deployed with careful planning in a way 52 that is appropriate local ecological and societal contexts (robust evidence, high agreement). 53 54 2.6.4 Adaptation for Increased Risk of Disease 55 56

Low-probability events can be very high impact (for example, the transmission of the SARS-CoV-2 from 1 wild animals to humans). A robust disease risk reduction policy would include utilizing a One Biosecurity 2 approach (Hulme, 2020) with actions to reduce disease risk across multiple sectors and from a variety of 3 anthropogenic drivers, including climate change, even if there is high uncertainty in projected risk (see 4 Cross-Chapter Boxes: ILLNESS, COVID, DEEP). Kraemer et al. (2019) found that vector importation was a 5 key risk factor and focus should be on preventing invasive species introductions. Further, many neglected 6 tropical diseases (NTDs) are also VBDs, and the UN SDG of good health and well-being explicitly calls for 7 increased control and intervention with a focus on emergency preparedness and response (Stensgaard et al., 8 2019a). Online tools are being developed to warn conservation biologists when species of conservation 9 concern are at greater risk of disease outbreaks due to environmental changes (e.g., for Hawaiian 10 Honeycreepers and avian malaria (Berio Fortini et al., 2020) and for coral diseases (Caldwell et al., 2016)). 11 Forecasting models to warn of human disease outbreaks like malaria and dengue are also now available, with 12 findings that multiple model ensemble forecasts outperform individual models (Lowe et al., 2013; Lowe et 13 al., 2014; Lowe et al., 2018; Zhai et al., 2018; Johansson et al., 2019; Tompkins et al., 2019; Muñoz et al., 14 2020; Colón-González et al., 2021; Petrova et al., 2021). Improving vector-borne disease and NTD public 15 health responses will require multi-disciplinary teams capable of interpreting, analyzing, and synthesizing 16 diverse components of complex ecosystem-based studies for effective intervention (Mills James et al., 2010; 17 Rubin et al., 2014; Valenzuela and Aksoy, 2018), broad epidemiological and entomological surveillance 18 (Depaquit et al., 2010; Lindgren et al., 2012; Springer et al., 2016), and community-based disease control 19 programs that build local capacity (Andersson et al., 2015; Jones et al., 2020b). 20

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23 [START CROSS-CHAPTER BOX ILLNESS HERE]

Cross-Chapter Box ILLNESS: Infectious Diseases, Biodiversity and Climate: Serious Risks Posed by Vector- and Water-borne Diseases

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Climate change is altering the life cycles of many pathogenic organisms and changing the risk of vector- and 38 water-borne infectious diseases transmission to humans (high confidence). Re-arrangement and emergence 39 of some diseases are already observed in temperate-zone and high-elevation areas, and coastal areas 40 (medium to high confidence). Shifts in the geographic and seasonal range suitability of pathogens and 41 vectors are related to climatic-impact drivers (warming, extreme events, precipitation, humidity) (high to 42 very high confidence), but there are substantial non-climatic drivers (land use change, wildlife exploitation, 43 habitat degradation, public health and socio-economic conditions) that affects the attribution of the overall 44 impacts on prevalence or severity of some vector- and water-borne infectious diseases over recent decades 45 (high confidence). Adaptation options that involve sustained and rapid surveillance systems, and the 46 preservation and restoration of natural habitats, with their associated higher levels of biodiversity, both 47 marine and terrestrial, will be key to reducing risk of epidemics and large-scale disease transmissions 48 (medium confidence). 49 50

51 Since AR5, further evidence is showing that climate-related changes in the geographic and seasonal range

⁵² suitability of pathogens and vectors and in the prevalence or new emergence of vector- and water-borne

infectious diseases have continued across many regions worldwide and are sustained over decadal time

scales (*medium to high confidence*) (WGII Sections 2.4.2.5, 3.5.5.3, 7.2, 7.3, 9.10.1.2.1; Harvell et al., 2009;

- ⁵⁵ Garrett et al., 2013; Burge et al., 2014; Guzman and Harris, 2015; Baker-Austin et al., 2018; Watts et al.,
- ⁵⁶ 2019; Semenza, 2020; Watts et al., 2021). Ecosystem-mediated infectious diseases at risk of increase from ⁵⁷ climate change include water-borne diseases associated with pathogenic *Vibrio* species (e.g., those causing

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cholera and vibriosis), and harmful algal blooms (e.g., ciguatera-fish poisoning) (SROCC 5.4.2.1.1, Box 5.4, 1 AR6 WGII 3.5, Table 3.S.3, 5.12; Baker-Austin et al., 2013; Levy, 2015; Trtanj et al., 2016; Ebi Kristie et 2 al., 2017; Mantzouki et al., 2018; Nichols et al., 2018), and vector-borne diseases associated with arthropods 3 (e.g., malaria, dengue, chikungunya, Zika virus, West Nile virus, and Lyme disease), helminths (e.g., 4 schistosomiasis) and zoonotic diseases associated with cattle and wildlife (e.g., leptospirosis) (medium to 5 high confidence) (Sections 2.4.2.7, 3.5, 7.2, 7.3, 9.10.1.1.1, 13.7.1.2, 14.4.6, Cross-Chapter Box COVID in 6 Chapter 7; Table Cross-Chapter Box ILLNESS.1; SR1.5 3.4.7.1; Ebi et al., 2021). 7

8

The attribution of observed changes in disease incidence partly or fully to climatic-impact drivers remains 9 challenging because of the difficulty of accurately capturing the contributions of multiple, interacting, and 10 often nonlinear underlying responses of host, pathogen, and vector, which can be influenced further by non-11 climate stressors and the long history of anthropogenic disturbance. Disease emergence in new areas requires 12 independent drivers to coincide (i.e., increasing climate suitability for pathogen or vector survival and 13 competence/capacity, and introduction of the pathogen via mobility of human populations). Further, the 14 extent to which changes in ecosystem-mediated diseases impact human health is highly dependent upon 15 local socio-economic status, sanitation, medical systems, and practices (Section 2.4.2.5; Figure FAQ2.3.1; 16 Gething et al., 2010; Lindgren et al., 2012; Mordecai et al., 2013; Liu-Helmersson et al., 2014; Bhatt et al., 17 2015; Morin et al., 2015; Ryan et al., 2015; Wesolowski et al., 2015; Stanaway et al., 2016; Yamana et al., 182016; Mordecai et al., 2017; Tesla et al., 2018; Ryan et al., 2019; Shah et al., 2019; Iwamura et al., 2020; 19 Mordecai et al., 2020; Colón-González et al., 2021; Ryan et al., 2021). Thus, the links between climate 20 change, ecosystem change, health and adaptation need to be considered concurrently (AR6 WGII 2.4, 3.5.3, 21 7.2, 7.3, 4.3.3, 6.2.2.3, Table 2.S.1). 22

23

24 Table Cross-Chapter Box ILLNESS.1: Observed climate change impacts on cholera, dengue, and malaria. 1) 25 26 Cholera: Endemicity based on Ali et al., 2015. Changes (2003-2018) in suitability for coastal Vibrio cholerae estimated 27 from model observations driven by sea-surface temperature (SST) and chlorophyll-a (CHL) concentration (Escobar et al., 2015; Watts et al., 2019). Vulnerabilities based on Sigudu et al., 2015, Agtini et al., 2005, and Sack et al., 2003. 2) 28 Dengue: Endemicity based on Guzmen et al., 2015. 3) Malaria: Endemicity based on Phillips et al. 2017, and WHO 29 Global Malaria Programme. Impacts of climate change on diseases and their vectors are most evident at the margins of 30 current distributions. However, climate change is difficult to implicate in areas with extensive existing transmission and 31 vector/pathogen abundance, and in particular is difficult to separate from concurrent directional trends in disease 32 control, changes in land use, water access, socioeconomic and public health conditions. As a result, while many studies 33 indicate increasing climate suitability of some areas for cholera, dengue, and malaria, the degree to which these changes 34 can be attributed to climate change remains challenging. For these cases, confidence statements of low, medium, or high 35 36 reflect confidence that variation in the disease and/or vector/pathogen is associated with variation in climate drivers, 37 rather than with directional climate change per se. Acronyms: ONI (Oceanic Niño Index), Tmin (minimum temperature), SPI (Standardised Precipitation Index), LST Land-surface temperature. Full references for this table can be seen in supplemental table 6 (see Table 2.S.6). 39

38

	Cholera	Dengue	Malaria
Africa			
Endemicity	Endemic	Endemic in sub-Saharan Africa but not S South Africa	Endemic
Climate drivers	Disease incidence: NE Africa, Central Africa & Madagascar: Rainfall (<i>medium confidence</i>) SE Africa: Rainfall, LST, SST, Plankton (<i>medium to high</i> <i>confidence</i>) ES Africa: SST, CHL (<i>low</i> confidence due to <i>limited</i> <i>evidence</i>) W Africa: Rainfall (floods), LST, SST (<i>medium to high confidence</i>)		W Africa: Temp (<i>medium to</i> <i>high confidence</i>) E Africa: Temp <i>medium to high</i> <i>confidence</i>)
Change and Confidence	Area of coastline suitable for outbreak: N&W Africa: Increase (low to medium confidence) C & E Africa: No change (low to medium confidence) S Africa: Decrease (low to medium confidence)	Potentially expanding (<i>low</i> <i>confidence</i>) Dengue and <i>Ae. aegypti</i> present but underdetected in climatically suitable areas.	E Africa: Upward shift and increase in malaria & Anopheles spp. in highland areas (medium to high confidence) Widespread decreases due to malaria control (medium confidence) and warming climate (low confidence)

Vulnerabilities	ES Africa: women of all ages more affected than men by outbreaks		
Asia	P 1 1		
Endemicity	Endemic	Endemic in S Asia, SE Asia, and E Asia	Endemic in S Asia, SE Asia,
Climate drivers	Disease incidence: E Asia: SST, CHL, Sea Level (medium to high confidence) S Asia: SST, CHL, LST, Rainfall(floods) (high to very high conficende)	S Asia: Rainfall, Temp, Humidity (medium confidence) SE Asia: Rainfall, Temp medium confidence) E Asia: Rainfall, Temp, Typhoons (low confidence)	Partially endemic in E Asia S Asia: Rainfall, Temp (medium to high confidence) SE Asia: Rainfall, Temp (medium confidence)
Change and Confidence	Area of coastline suitable for outbreak: Increase (<i>low to</i> <i>medium confidence</i>)	SE Asia: Increase (low confidence) S Asia: Increase (medium confidence) E Asia: Increase (low confidence)	S Asia: Increase (medium confidence)
Vulnerabilities	SE Asia: infants (<9 years) with highest incidences of cholera S Asia: older children and young adults (16-20 years old) more frequently reported with cholera than non-cholera diarrhoea	5	
Australasia			
Endemicity	Not endemic	Partially endemic in N Australia	Not endemic
Climate drivers	No evidence for disease incidence	Rainfall, Temp (low confidence)	×
Change and Confidence	Area of coastline suitable for outbreak: No change (<i>low to</i> <i>medium confidence</i>)	Increase in sporadic outbreaks due to climate change (<i>low</i> <i>confidence</i>)	No change
Central America			
Endemicity	Not endemic	Endemic	Partially endemic
Climate drivers	No evidence for disease incidence	ONI, SST, Tmin, Temp, Rainfall, Drought (<i>low</i> <i>confidence</i>)	
Change and Confidence	Areas of coastline suitable for outbreak: Decrease (<i>low to</i> <i>medium confidence</i>)	Increasing due to climate (<i>low</i> confidence) Upward expansion of <i>Ae. aegypti</i> (<i>low confidence</i>)	Overall decrease not linked to climate change. Focal increases due to human activities.
South America			F 1 .
Endemicity	Epidemic	Endemic in all regions except S South America	Endemic
Climate drivers	Abundance of coastal V.	Temp Prec Drought	
	cholerae: NW South America: SST, Plankton (low to medium confidence)	Temp, Prec, Drought	N South America: Temp (<i>low</i> <i>confidence</i>) N SE South America: Tmax, Tmin, humidity (<i>low confidence</i>)
Change and	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for	Increasing due to urbanization	<i>confidence)</i> N SE South America: Tmax,
Change and Confidence	SST, Plankton (low to medium confidence)		<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i>
Confidence	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to</i>	Increasing due to urbanization and decreased vector control programmes, not strongly linked	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions:
Confidence Europe	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to</i>	Increasing due to urbanization and decreased vector control programmes, not strongly linked to climate	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions:
Confidence	SST, Plankton (low to medium confidence) Area of coastline suitable for outbreak: No change (low to medium confidence)	Increasing due to urbanization and decreased vector control programmes, not strongly linked	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions: Increase <i>(low confidence)</i>
Confidence Europe Endemicity Climate drivers	SST, Plankton (<i>low to medium</i> <i>confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to</i> <i>medium confidence</i>) Not endemic No evidence for disease incidence Abundance of coastal V. <i>cholerae</i> : N Europe: SST, Plankton (medium confidence)	Increasing due to urbanization and decreased vector control programmes, not strongly linked to climate S Europe: focal outbreaks	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions: Increase <i>(low confidence)</i> Not endemic
Confidence Europe Endemicity	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to medium confidence</i>) Not endemic No evidence for disease incidence Abundance of coastal <i>V. cholerae</i> : N Europe: SST, Plankton (medium confidence) Area of coastline suitable for outbreak: Increase (<i>low to</i>	Increasing due to urbanization and decreased vector control programmes, not strongly linked to climate S Europe: focal outbreaks Mediterranean regions of S Europe: Outbreaks	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions: Increase <i>(low confidence)</i>
Confidence Europe Endemicity Climate drivers Change and Confidence	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to medium confidence</i>) Not endemic No evidence for disease incidence Abundance of coastal <i>V. cholerae</i> : N Europe: SST, Plankton (medium confidence) Area of coastline suitable for	Increasing due to urbanization and decreased vector control programmes, not strongly linked to climate S Europe: focal outbreaks Mediterranean	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions: Increase <i>(low confidence)</i> Not endemic
Confidence Europe Endemicity Climate drivers Change and	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to medium confidence</i>) Not endemic No evidence for disease incidence Abundance of coastal <i>V. cholerae</i> : N Europe: SST, Plankton (medium confidence) Area of coastline suitable for outbreak: Increase (<i>low to</i>	Increasing due to urbanization and decreased vector control programmes, not strongly linked to climate S Europe: focal outbreaks Mediterranean regions of S Europe: Outbreaks <i>(low confidence)</i> Partially endemic in S North	<i>confidence)</i> N SE South America: Tmax, Tmin, humidity <i>(low confidence)</i> Higher elevation regions: Increase <i>(low confidence)</i> Not endemic
Confidence <i>Europe</i> Endemicity Climate drivers Change and Confidence <i>North America</i>	SST, Plankton (<i>low to medium confidence</i>) Area of coastline suitable for outbreak: No change (<i>low to medium confidence</i>) Not endemic No evidence for disease incidence Abundance of coastal V. <i>cholerae</i> : N Europe: SST, Plankton (medium confidence) Area of coastline suitable for outbreak: Increase (<i>low to medium confidence</i>)	Increasing due to urbanization and decreased vector control programmes, not strongly linked to climate S Europe: focal outbreaks Mediterranean regions of S Europe: Outbreaks <i>(low confidence)</i>	<pre>confidence) N SE South America: Tmax, Tmin, humidity (low confidence) Higher elevation regions: Increase (low confidence) Not endemic Not endemic</pre>

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Change and Confidence	No evidence for disease incidence Abundance of coastal V. cholerae: EN America: SST (Low due to limited evidence)	Declining	No change
Small Islands			
Endemicity	Epidemic	Endemic in many small islands the Tropics	in Endemic in many small islands in the Tropics
Climate drivers	Disease incidence: Caribbean: SST, LST, Rainfall (low to medium confidence)	Caribbean: SPI, Tmin (low confidence)	
Change and Confidence	Area of coastline suitable for outbreak: Caribbean & Pacific Small Island: Decrease (low to medium confidence)	Increasing (low confidence)	Decrease in Caribbean not linked to climate

3

Observed and projected changes

4 In aquatic systems, at least 30 human pathogens with water infection-routes (freshwater and marine) are 5 affected by climate change (Section 3.5.3 Table SM3.G; Nichols et al., 2018). Warming, acidification, 6 hypoxia, sea-level rise, and increases in extreme weather and climate events (e.g. marine heatwaves, storm 7 surges, flooding, and drought), which are projected to intensify in the 21st century (high confidence) (AR6 8 WGI SPM B2.2, B.3.2), are driving species' geographic range shifts and global rearrangements in the 9 location and extent of areas with suitable conditions for many harmful pathogens, including viruses, bacteria, 10 algae, protozoa, and, helminths (high confidence) (Sections 2.3, 2.4.2.7, 3.5.5.3, SROCC 5.4.2.1.1, Box 5.4; 11 Trtanj et al., 2016; Ebi Kristie et al., 2017; Manning and Nobles, 2017; Pecl et al., 2017; Mantzouki et al., 12 2018; Nichols et al., 2018; Bindoff et al., 2019; Kubickova et al., 2019; Watts et al., 2019; Watts et al., 2020; 13 Watts et al., 2021). 14 15

Incidence of cholera and *Vibrio*-related disease outbreaks has been shown to originate primarily in coastal regions, and then spread inland via human transportation. Our understanding of impacts of climate drivers on the dynamics of *Vibrio*-related infections have been strengthened through improved observations from longterm monitoring programs (e.g., (Vezzulli et al., 2016)), and statistical modelling supported by large-scale and high-resolution satellite observations of climate drivers (*high confidence*) (e.g., Baker-Austin et al., 2013; Escobar et al., 2015; Jutla et al., 2015; Martinez et al., 2017; Semenza et al., 2017; Racault et al., 2019; Campbell et al., 2020).

23 The coastal area suitable for V. cholerae (the causative agent for cholera) has increased by 9.9% globally 24 compared to a 2000s baseline (Escobar et al., 2015; Watts et al., 2019). The poleward expansion of the 25 distribution of Vibrio spp. has increased the risk of vibriosis outbreaks in northern latitudes. Specifically, the 26 coastal area suitable for Vibrio infections in the past 5 years has increased by 50.6% compared with a 1980s 27 baseline at latitudes of 40-70°N; in the Baltic region, the highest-risk season has been extended by 6.5 28 weeks over the same periods (Watts et al., 2021). Already, studies have noted greater numbers of Vibrio-29 related human infections, and most notably disease outbreaks linked to extreme weather events such as heat 30 waves in temperate regions such as Northern Europe (Baker-Austin et al., 2013; Baker-Austin et al., 2017; 31 Baker-Austin et al., 2018) (high confidence). By the end of the 21st century, under RCP6.0, the number of 32 months of risk of Vibrio illness is projected to increase in Chesapeake Bay by 10.4±2.4%, with largest 33 increases during May and September, which are the months of strong recreational and occupational use, 34 compared to a 1985-2000 baseline (Jacobs et al., 2015; Davis et al., 2019a). In the Gulf of Alaska, the 35 coastal area suitable for Vibrio spp. is projected to increase on average by 58%±17.2% in summer under 36 RCP6.0 by the 2090s, compared to a 1971-2000 baseline (low to medium confidence) (Jacobs et al., 2015). 37 38

39 On land, increased global connectivity and mobility, unsustainable exploitation of wild areas and species,

⁴⁰ land conversion (agricultural expansion, intensification of farming, deforestation, infrastructure

41 development), together with climate-change-driven range shifts of species and human migration (Cross-

Chapter Box MOVING PLATE in Chapter 5), have modified interfaces between people and natural systems
 (IPBES, 2018a). Climate-driven increase in temperature, frequency and intensity of extreme events, and

(IPBES, 2018a). Climate-driven increase in temperature, frequency and intensity of extreme events, and
 changes in precipitation and relative humidity, have provided opportunities for re-arrangements of disease

45 geography and seasonality, and emergence into new areas (*high confidence*) (Section 2.4.2.7). In particular,

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malaria has expanded into higher elevations in recent decades and although climate change attribute remains 1 challenging (Hay et al., 2002; Pascual et al., 2006; Alonso et al., 2011; Campbell et al., 2019c), evidence that 2 the elevational distribution of malaria has tracked warmer temperatures is compelling for some regions (Siraj 3 et al., 2014). Models based on both empirical relationships between temperature and the Anopheles mosquito 4 and Plasmodium parasite traits that drive transmission (Mordecai et al., 2013; Yamana Teresa and Eltahir 5 Elfatih, 2013; Johnson et al., 2015) and existing mosquito distributions (Peterson, 2009) predict that 6 warming will increase the risk of malaria in highland East Africa and Southern Africa, while decreasing the 7 risk in some lowland areas of Africa, as temperatures exceed the thermal optimum and upper thermal limit 8 for transmission (Peterson, 2009; Yamana Teresa and Eltahir Elfatih, 2013; Ryan et al., 2015; Watts et al., 9 2021). 10

11 In contrast to malaria, dengue has expanded globally since 1990, particularly in Latin America and the 12 Caribbean, South Asia, and sub-Saharan Africa (Stanaway et al., 2016). While urbanization, changes in 13 vector control, and human mobility play roles in this expansion (Gubler, 2002; Åström et al., 2012; 14 Wesolowski et al., 2015), the physiological suitability of temperatures for dengue transmission is also 15 expected to have increased as climates have warmed (Colón-González et al., 2013; Liu-Helmersson et al., 16 2014; Mordecai et al., 2017; Rocklöv and Tozan, 2019). Models predict that dengue transmission risk will 17 expand across many tropical, subtropical, and seasonal temperate environments with future warming 18 (Åström et al., 2012; Colón-González et al., 2013; Ryan et al., 2019; Iwamura et al., 2020; Watts et al., 19 2021)). 20

22 Adaptation options

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23 During the 21st century, public health adaptation measures (Figure Cross-Chapter Box ILLNESS.2) have 24 been put in place in attempts to control or eradicate a variety of infectious diseases by improving 25 surveillance and early detection systems; constraining pathogen, vector, and/or reservoir host distributions 26 and abundances; reducing likelihood of transmission to humans; and improving treatment and vaccination 27 programs and strategies (medium to robust evidence, medium to high agreement) (Chinain et al., 2014; 28 Adrian et al., 2016; Friedman et al., 2017; Konrad et al., 2017; Semenza et al., 2017; Borbor-Córdova et al., 29 2018; Rocklöv and Dubrow, 2020). In addition, effective management and treatment of domestic and 30 wastewater effluent, through better infrastructure and preservation of aquatic systems acting as natural water 31 purifiers, have been key to securing the integrity of the surrounding water bodies, such as groundwater, 32 reservoirs and lakes, and agricultural watersheds, as well as protecting public health (*high confidence*) 33 (Okeyo et al., 2018; Guerrero-Latorre et al., 2020; Kitajima et al., 2020; Sunkari et al., 2021). The 34 preservation and restoration of natural ecosystems, with their associated higher levels of biodiversity, have 35 been reported as significant buffers against epidemics and large-scale pathogen transmission (medium 36 confidence) (Johnson and Thieltges, 2010; Ostfeld and Keesing, 2017; Keesing and Ostfeld, 2021). Further, 37 the timely allocation of financial resources and sufficient political will in support of a "One Health" 38 scientific research approach, recognising the health of humans, animals and ecosystems as interconnected 39 (Rubin et al., 2014; Whitmee et al., 2015; Zinsstag et al., 2018), holds potential for improving surveillance 40 and prevention strategies that may help to reduce the risks of further spread, and new emergence of 41 pathogens and vectors (medium confidence) (Destoumieux-Garzón et al., 2018; Hockings et al., 2020; 42 Volpato et al., 2020; Hopkins et al., 2021; Services and Ecosystem, 2021). 43 44

45

Туре		Description of adaptation optio	ns	Climate impact	Confidence
Warning systems	Early surveillance in wildlife & humans	Seasonal & dynamic forecasts of disease outbreaks; detailed risk mapping	Early Warning systems targeted locally	(t-2)	+++
Diagnostic abilities	Technology & trained personnel to permit rapid diagnosis and awareness of cases	Reporting in near-real time, for efficient response & resources mobilization for mitigation	Rapid response to disease emergence events, with public health and medical resources		+++
Capacity building	Training health & environmental officials to respond to new disease emergent risks	Awareness of local populations of the health risks from pathogens & vectors	Robust healthcare systems with good facilities, access & epidemic protocols	(t-)	++
Public policy	Policy-making and international cooperation within a One Health framework	Large-scale public health programs for diseases/vectors eradication	Herd-immunity level vaccination for pathogens with few host species		++
Financing	Green recovery funds to tackling biodiversity loss & climate change	National funds for nature-based projects for forest conservation, water services	Funds to provide jobs for tribal groups in plantation work, forest & water management		8+
Technology	Non-insecticide-based controls of vectors	Other control alternatives (avoiding use of antibiotics & chemical drugs)	Genetic surveillance & control of disease vectors & pathogens		<u> </u> ++
Management	Planning aligned with climate targets	Long-term observing & monitoring systems	Environmental regulations & sustainable agriculture, livestock & fisheries farming practice		8++
Infrastructure	Urban forests & green spaces, standing water removal	Drinking water access, sewage & drainage maintenance	Better homes keeping mosquitoes out of habited-indoor areas	(t-e)	= +++
Nature-based solutions	Natural habitats restoration, reforestation	Reducing habitats fragmentation & limiting human proximity to risky environments	Ecosystem-based management to regulate pathogens & vectors population		8++
Practice change	Diets diversification, more resilient food systems	Reduction of wildlife trade	Alternatives to reduce reliance on bushmeat and usage of wild animals		++
Co-benefits from mitigation	Reducing local emission from energy systems	Clean transport systems	Better access to food, water & energy		8++

Adaptation measures to reduce risks of ecosystem-mediated diseases under climate change

Figure Cross-Capter Box ILLNESS.1: Adaptation measures to reduce risks of climate change impact on water- and vector-borne diseases. Impacts are identified at three levels: 1) impact on pathogen, host/vector distributions and abundance; 2) impact on pathogen-host transmission cycle occurrence and efficiency; and 3) impact on likelihood of transmission to humans. Adaptation typology is based on (Biagini et al., 2014; Pecl et al., 2019). For each type of adaptation, examples are provided with their level of evidence and agreement.

[END CROSS-CHAPTER BOX ILLNESS HERE]

2.6.5 Adaptation in Practice: Case Studies and Lessons Learned

Adaptation plans for biodiversity and EbA have been adopted in many places and different scales but it is difficult to get a systematic overview of adaptation in practice. We have therefore reviewed a series of contrasting case studies to illustrate the some key issues. There is a pressing need for more thorough monitoring and evaluation of adaptation to assess effectiveness. Climate change adaptation is conceptually difficult to measure but it is possible to test which techniques work in reducing vulnerability and monitor their deployment (Morecroft et al., 2019).

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Adaptation can take place at a range of scales with specific projects nested within overarching national strategies. Small scale projects can be adaptation focused, but in larger scale adaptation is often integrated with wider objectives. Within an urban or peri-urban context, the benefits of natural and semi-natural areas for health and well- help to justify support for EbA. Economic wellbeing is also an important factor in many cases, whether as, in Durban (Section 2.6.5.7), by providing new job opportunities or, in the Andes (Section 2.6.5.4) by supporting long-established agricultural practice. Action on the ground often depends on factors

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1			and international funding. Within Durban, academic community were essential,
2 3			amples of communities using traditional or
4			r no external input, (Section 2.6.5.4). They
5		cope to adapt by factors beyond th	
6		1 1 2 2	
7		t species from climate change, suc	
8			Heritage Area (Section 2.6.5.8), are rare.
9 10	· · · · · · · · · · · · · · · · · · ·		ed or where ecosystem restoration projects dapt to weather conditions and their
11			od practice, but it may not be sufficiently
12			. Training and resources to support
13	conservation practitioners are b	ecoming available to help address	s this. Examples include the Climate Change
14		(Section 2.6.5.2), and The Alliand	
15			he sustainable management of freshwater
16 17	biodiversity (Darwall et al., 20)	18).	100
18	Adaptation is widely recognise	d as important for national conser	vation policy and is being considered in a
19			strategic context includes decisions about
20			tifying places which can act as refugia. It
21			ant but will support a changing range of
22		important for directing resources	r uncertainties and the extent to which there
23 24			in reducing greenhouse gas emissions
24 25			nental adjustments to conservation
26			require radical, transformational changes to
27		system services (Morecroft et al.,	
28			
29		conservation are relatively comm	
30			ck of resources for conservation in many
31			value protected sites in their present form.
32 33		ential in gaining support for such c	Initiatives to engage wider communities in
33 34	discussions are likely to be esse	sitial in gaining support for such c	changing approaches.
35	EbA and adaptation for biodive	ersity are intrinsically linked and t	he largest scale interventions for adaptation
36	in ecosystem have tended to bri	ing together both elements. For ex	ample adaptation to reduce flood risk by
37			oss-Chapter Box SLR in Chapter 3), such as
38			s for water storage, offers the potential to
39	meet multiple objectives and ha	s increased overall funding available	ble for ecosystem restoration.
40	2651 Case study: Assisted Co	olonisation / Managed Relocation	in Practice
41 42	2.0.3.1 Case study. Assisted Co	ionisation / Managea Relocation	in i ruciice
43	Scale: Global		
44	Issue: Helping species move to	track shifting climate space	
45			
46			the movement of species, populations, or
47			ons (Hoegh-Guldberg et al., 2008) and may
48	· ·	A	urally. It requires careful consideration of
49 50		0	of relocation and the receiving ecosystem
50 51	(110egn-Guidderg et al., 2008; J	Richardson et al., 2009; Schwartz	ct al., 2012).
51 52	Individual cases show that assis	sted migration can be successful	Anich & Ward (2017) extended the
53			phaga kirtlandii, by 225km by using song
54			cessfully transplanted an annual legume,
55	Chamaecrista fasciculata, to si	tes beyond its current poleward ra	nge limit, while Liu (2012) found that all
56	but one of 20 orchid species sur	rvived when transplanted to highe	r elevations than their current range limits.

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1 2	margins, Willis (2009) observed that for at least 8 years.	both introduced populations	grew, expanded their ranges and survived
3 4 5 6 7 8	they are easy to move and their range butterfly, <i>Carterocephalus palaemon</i> either the species' poleward or equate parameter for re-introduction, using S	e dynamics have been extension, became locally extinct in E borial range limits. Nonetheles	esponse to regional climate warming, since ively studied. The Chequered Skipper ngland in the 1970's, in an area not close to ss, Maes (2019) consider climate a crucial source population in Belgium and
9 10 11	introduction site. Success of assisted migration for con		
12 13 14 15 16 17 18	fasciculata were more successful wh	led failure was most strongly al source of failure is local ac source populations can be in en sourced from the most po	associated with low numbers of
19 20 21	2.6.5.2 Case study: Adaptation for co Kingdom	onservation and Natural Floo	od Management in England, United
22 23 24	Scale: National Issue: National approach to adaptatic	on in the natural environment	
25 26 27 28 29 30 31 32 33 34 35 36	of more frequent extreme weather even management: with habitats fragmente people, which are exacerbated by enven A National Adaptation Programme, p on the natural environment. There are England, the conservation agency an including flood defence. The principle conservation community and resource Adaptation Manual jointly with the F (England and RSPB, 2020) and spati	rents such as droughts. These ed, land often drained and riv vironmental factors, including provides a broad policy frame e also adaptation plans produ d the Environment Agency, v les of climate change adaptat res are available. Natural Eng Royal Society for the Protecti al mapping tool for climate c	with a wide range of responsibilities ion are well established in the UK gland has published a Climate Change on of Birds–a major conservation NGO hange vulnerability (Taylor et al., 2014).
 37 38 39 40 41 42 43 44 45 	managers and that they were implem restoring ecosystem processes and re to change management objectives of implementation of such changes (Du coast where rising sea level is causin marine ones.	enting actions that might build ducing fragmentation. Ther protected sites to adjust to ch ffield et al., 2021). The main g transitions from terrestrial	tion was common amongst nature reserve lding resilience to climate change, such as e is a recognition that it will be necessary nanging circumstances but there was little examples of managing change, was at the and freshwater systems to coastal and
45 46 47	Natural Flood Management (NFM):	restoring natural processes an	n in England but the best developed is nd natural habitats to reduce flood risk ojects have been established in local areas.

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NFM includes a broad range of techniques, some of which, deliver real benefits for biodiversity and allow natural ecological process to re-establish. Others, such as creating 'woody debris dams' – barriers artificially constructed from tree trunks and branches in water courses to slow flow of water –will have fewer benefits, although they may benefit some species. Dadson et al. (2017) concluded that 'the hazard associated with small floods in small catchments may be significantly reduced' by natural flood management techniques.

The Environment Agency collated the evidence base on NFM (Burgess-Gamble et al., 2017) and was able to

draw on 65 case studies (Ngai et al., 2017), covering river and floodplain management, woodland

management, run-off management and coast and estuary management.

FINAL DRAFT Chapter 2 IPCC WGII Sixth Assessment Report However, they noted that the most extreme flood events may overwhelm any risk management measures and 1 failed to find clear evidence of NFM in reducing flood risk downstream in large catchments. 2 3 There remain challenges in deploying NFM at larger scales, partly reflecting the time necessary to 4 demonstrate the effectiveness of pilot studies and build confidence and building stakeholder support is 5 important (Huq et al., 2017). There are now a number of examples of where collaborative initiatives between 6 local communities, land owners and government agencies have been successful in establishing effective 7 NFM schemes (Short et al., 2019). 8 9 2.6.5.3 Case Study: Protected Areas Planning in Response to Climate Change in Thailand 10 11 Scale: National 12 Issue: Protected area network planning 13 14 Many countries in the Association of South East Asian Nations (ASEAN) are expanding protected area 15 networks to meet the Aichi target 11 of at least 17% of terrestrial area protected and it is important to take 16 the effects of climate change into account. Existing protected areas in Thailand cover approximately 21% of 17 the land area, and it is one of the few tropical countries that passes the Aichi Target 11. Most protected areas 18 in Thailand were established on an ad hoc basis to protect remaining forest cover, and as a result do not 19 represent diverse habitats and their associated species (Chutipong et al., 2014; Tantipisanuh, 2016) and may 20 not be resilient to the interacting impacts of future land use and climate change (Klorvuttimontara et al., 21 2011; Trisurat, 2018). 22 23 Recent research conducted in northern Thailand indicated that the existing protected areas (31% of the 24 region area) cannot secure viability of many medium- size and large mammals. Most species climate space 25 would substantially shift, bringing a risk of extinction. The model results based on the spatial distribution 26 model and network flow determined there was a need for expansion areas of 5,200 km² or 3% of the region 27 to substantially minimise the high-risk level and increase the average coping capacity of the protection of 28 suitable habitats from 82% as the current plan to 90%. These results were adopted by the Thailand's 29 Department of National Parks, Wildlife and Plant Conservation and included in the National Wildlife 30 Administration and Conservation Plan (2021-2031). 31 32 Case Study: Effects of Climate Change on Tropical High Andean Social Ecological Systems 2.6.5.4 33 34 Scale: Regional 35 Issue: Complex ramifications of glacial retreat on vegetation, animals, herders and urban populations 36 37 Accelerated warming is shrinking tropical glaciers at rates unseen since the middle of the Little Ice Age 38 (Rabatel et al., 2013; Zemp et al., 2015). Climate-driven upward migration of species associated with 39 warming and glacier retreat has modified species distribution and richness, and community composition 40 along the Andes altitudinal gradient (Seimon et al., 2017; Carilla et al., 2018; Zimmer et al., 2018; Moret et 41 al., 2019). Climate-driven glacier retreat alters hydrological regimes, impacting Andean pastoralists directly 42 (López-i-Gelats et al., 2016; Postigo, 2020; Thompson et al., 2021), and water provisioning to lowland 43 regions (Vuille et al., 2018; Hock et al., 2019; Orlove et al., 2019; Rasul and Molden, 2019). Drying 44 wetlands has modified alpine plant communities, which are relevant to storing carbon, regulating water, and 45 providing food for local livestock, leading to negative impacts on herders' livelihoods (Dangles et al., 2017; 46 Polk et al., 2017; Postigo, 2020) and differently affecting the wild vicuña and the domesticated alpaca and 47 llama. Vicuña (Vicugna vicugna) and alpaca (Vicugna pacos) wool are important income sources for 48 indigenous communities and Llama (Lama glama) is the main source of meat. Vicuña is adjusting its feeding 49 behaviour and spatial distribution as vegetation migrates upwards (Reider and Schmidt, 2020), causing them 50 to roam outside protected areas and become vulnerable to illegal poaching. 51 52

Andean herders have responded to drying of grasslands by increasing livestock mobility, accessing new grazing areas through kinship and leases, creating and expanding wetlands through building long irrigation canals (of several km), limiting allocation of wetlands to new households, and sometimes cultivating grasses (Postigo, 2013; López-i-Gelats et al., 2015; Postigo, 2020). These adaptive responses to regional climate change are enabled by deeply-embedded indigenous institutions that have traditionally governed Andean

2	2010; Postigo, 2019; Postigo, 2020). For instance, water quality, access and control by local pastoralists has
3	declined due to new mining concessions in headwaters of Andean watersheds (Bebbington and Bury, 2009)
4	and diversion of water to lowland coastal desert for agricultural irrigation (Mark et al., 2017).
5	
6	Glacier mass and runoff in the tropics are projected to reduce >70% and >10%, respectively, by 2100 under
7	RCP 2.6, RCP 4.5 and RCP 8.5 (Huss and Hock, 2018; Hock et al., 2019). In Peru, montane ice-field melt-
8	water provides 80% of the water resources for the arid coast where half the population lives (Thompson et
9	al., 2021). Increasing variability of precipitation has compromised rain-fed agriculture and power generation,
10	particularly in the dry season, exacerbating pressures for new water sources (Bradley et al., 2006; Bury et al.,
11	2013; Buytaert et al., 2017). Thus, there is risk of increasing conflicts between climate change adaptation to
12	benefit high Andean human and natural communities and adaptation to maintain water provisioning for
13	lowland agricultural and urban areas.
14	
15	2.6.5.5 Case Study: Helping African Penguins Adapt to Climate Change
	2.0.5.5 Cuse blady. Helping fijnean i enganis flaapi to chinate change
16	
17	Scale: Regional / local
18	Issue: Adaptation for a threatened species
19	
20	The African penguin, Spheniscus demersus, is the only resident penguin species on mainland Africa and
21	breeds in a handful of colonies in South Africa and Namibia. In 2017, penguins in Cape Town's Boulders
22	Beach colony attracted almost one million visitors, providing 885 jobs and \$18.9m revenue (Van Zyl and
23	Kinghorn, 2018). Ninety-six percent of the population has been lost since 1900, with a 77% decline in the last
24	two decades (Sherley et al., 2018) and by 2019 only 17,700 pairs remained (Sherley et al., 2020). The
25	species is listed as Endangered on the IUCN Red List (IUCN, 2018) and if this trajectory persists the African
26	penguin will become functionally extinct in the near future (Sherley et al., 2018).
27	
28	Historically, hunting, egg and guano collection were the species' main threats, but three aspects of climate
	change now predominate. Firstly, a several-hundred-kilometre eastward shift in distributions of their main
29	
30	prey species, anchovies and sardines, has reduced food availability (Roy et al., 2007; Crawford et al., 2011).
31	While adult penguins typically forage up to 400 km from their colonies, they are restricted to a ~20 km
32	radius from their colonies during breeding months (Ludynia et al., 2012; Pichegru et al., 2012). The resulting
33	food shortage at this critical time is compounded by competition with commercial fisheries and
34	environmental fluctuations (Crawford et al., 2011; Pichegru et al., 2012; Sherley et al., 2018). This has
35	impacted adults' survival and their ability to raise high-quality offspring (Crawford et al., 2006; Crawford et
36	al., 2011; Sherley et al., 2013; Sherley et al., 2014).
	al., 2011, Sheney et al., 2015, Sheney et al., 2014).
37	
38	Increasing heat wave frequency and intensity recorded in recent decades presents a second threat (van
39	Wilgen and Wannenburgh, 2016; Van Wilgen et al., 2016; Mbokodo et al., 2020). Nests were historically
40	built in insulated guano burrows, but are now frequently sited on open ground (Kemper et al., 2007;
41	Pichegru et al., 2012; Sherley et al., 2012). High temperatures frequently expose the birds to severe heat
42	stress, causing adults to abandon nests and resulting in mortality of eggs and chicks (Frost et al., 1976;
	Shannon and Crawford, 1999; Pichegru et al., 2012).Intensifying storm surges and greater wave heights can
43	
44	cause nest flooding (Randall et al., 1986; de Villiers, 2002).
45	
46	The African penguin's survival in the wild is dependent on the success of adaptation action. Increasing
47	access to food resources is a management priority (IUCN, 2018). One approach is to reduce fishing pressure
48	immediately around breeding colonies. An experiment excluding fishing around colonies since 2008 has
49	demonstrated positive effects (Pichegru et al., 2010; Pichegru et al., 2012; Sherley et al., 2015; Sherley et al.,
50	2018; Campbell et al., 2019b). A second approach is to establish breeding colonies closer to their prey. An
51	ongoing translocation initiative aims to entice birds eastwards to recolonise an extinct breeding colony and
52	potentially to establish a new one (Schwitzer et al., 2013; Sherley et al., 2014; International, 2018). Penguin
53	"look-alikes" or decoys, constructed from rubber and concrete, have been placed at the extinct colony site
54	and, along with call play-backs, give the illusion of an established penguin colony (Morris and Hagen,
55	2018). This approach has not yet proven successful.
	2010). This approach has not yet proven successful.

2-125

Chapter 2

pastoralists, but have become severely compromised by national socio-economic pressures (Valdivia et al.,

2010; Postigo, 2019; Postigo, 2020). For instance, water quality, access and control by local pastoralists has

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1 2 3 4 5 6	nesting sites that also provide s 2018; International, 2018). Art have been explored in combina	torm protection and are sufficiently ficial nest boxes of various designs tion with use of natural vegetation. emper et al., 2007; Sherley et al., 2	atives are underway to provide cooler y above the high water level (Extinction, s and constructed from a range of materials . Some designs have proven successful, 012), but the same designs have had less
7 8 9 10 11 12 13 14 15 16 17	moulting parents, being shore- orphaned chicks have been rele of 77% (Schwitzer et al., 2013; at Boulders Beach aims to use of penguin heat stress as trigge established chick-rearing facili	bound, are unable to feed late-hatch cased into the wild as part of the Ch Sherley et al., 2014; Klusener et a real-time weather station data, with rs for implementing a Heat Wave F ties and a large body of expertise, t	m eggs, has long proven valuable because hing chicks. Since 2006, over 7,000- hick Bolstering Project with a success rate 1., 2018; SANCCOB, 2018). A new project hin-nest temperatures and known thresholds Response Plan. Drawing on well- his includes removing heat-stressed eggs ed that such birds may be released at the
17 18 19 20 21 22	<i>California, United St</i> Scale: Local		e Joshua tree in Joshua Tree National Park,
23 24 25 26 27 28 29 30	the national park is arid, with a annual precipitation of 170 ± 8 temperature increased at a sign precipitation decreased at a sig Anthropogenic climate change	1971-2000 average summer temper 0 mm y ⁻¹ (Gonzalez et al., 2018). F ificant ($p < 0.0001$) rate of 1.5 ± 0 . nificant ($p = 0.0174$) rate of -32 ± 1	
31 32 33 34 35 36 37 38 39 40 41	particularly the Joshua tree (Yu small animals and holds cultura Joshua tree. Paleobiological da ground sloth (<i>Nothrotheriops s</i> present a across a wider range, 2010; Cole et al., 2011). A maj coinciding with a warming in t	Il significance. The national park p ta from packrat (<i>Neotoma spp.</i>) mic <i>hastensis</i>) show that Joshua trees g extending as far as 300 km south in or retraction of the range began ~1 he region of 4°C, caused by Milank	ral features unique to the region, ca that provides habitat for birds and other rotects the southernmost populations of the ddens and fossil dung of the extinct Shasta rew 13 000–22 000 thousand years before nto what is now México (Holmgren et al., 1 700 thousand years before present, covitch cycles, that marked the end of the ggesting a sensitivity of Joshua trees of 300
42 43 44 45 46 47	Joshua tree could shift north an mortality would increase from	d the species become extirpated fro drought stress and wildfires, which e and may continue to fuel (Brooks	s over 4°C by 2100, suitable climate for the om the park (Sweet et al., 2019). Plant have been rare or absent in the Mojave, and Matchett, 2006; DeFalco et al., 2010;
48 49 50 51 52 53 54 55 56 57	of extirpation prompted them to conditions may persist for the s analyses of suitable climate to (Barrows and Murphy-Marisca invasive grasses and fire contro- using native plants, including r	b adapt conservation to focus on pr pecies into the future (Barrows et a identify potential refugia under all 1, 2012; Sweet et al., 2019). The pa ol (Barrows et al., 2020)and works sursery-grown Joshua tree seedlings	er in the park it was found. The future risk otecting potential refugia, where suitable al., 2020). The national park used spatial emissions scenarios except for the highest ark prioritises the refugia for removal of to restore refugia that have burned in fires, s. The park and its partners are monitoring warning of any changes (Barrows et al.,

2.6.5.7 Case Study: Ecosystem based Adaptation in Durban, South Africa

4 Scale: Local

5 Issue: Ecosystem based adaptation (EbA) in a city and surrounding area

6 Durban was an early pioneer of EbA in a city context, establishing a Municipal Climate Protection 7 Programme (MCPP) in 2004 (Roberts et al., 2012). The City, situated in a global biodiversity hotspot 8 (Bank, 2016), has a rapidly growing population (approximately 3.5 million) and is highly fragmented 9 (Roberts et al., 2013). High levels of development, particularly in peri-urban areas, have encroached into 10 natural habitats (Bank, 2016). Degradation of the natural resource base in this way has direct economic and 11 financial costs, is threatening the City's long-term sustainability, and is exacerbated by climate change 12 (Bank, 2016; Municipality, 2020). The impacts of climate change are anticipated to increase unless 13 appropriate mitigation and adaptation interventions are prioritised (Municipality, 2020). High rates of 14 poverty, unemployment and health problems have pushed Durban to explore a climate change adaptation 15 work stream within its MCPP (Roberts et al., 2013; Roberts et al., 2020b). 16

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A single approach to adaptation is likely to be insufficient (Archer et al., 2014), and community-based adaptation should be integrated as part of a package of tools applied at the city level. Durban's climate change adaptation work stream is composed of three separate components: municipal adaptation (adaptation activities linked to the key functions of local government); community-based adaptation (focused on improving the adaptive capacity of local communities); and a series of urban management interventions (addressing specific challenges such as the urban heat island, increased storm water runoff, water conservation and sea level rise) (Roberts et al., 2013).

25

39

Lessons learnt from Duban's experience include the importance of meaningful partnerships, long-term 26 financial commitments (Douwes et al., 2015) and significant political and administrative will (Roberts et al., 27 2012; Roberts et al., 2020b). Securing these requires strong leadership (Douwes et al., 2015), including from 28 local champions (Archer et al., 2014), even if EbA is considered cost-effective (Roberts et al., 2012). Natural 29 habitat restoration projects are seen as an ideal tool, as they combine mitigation outcomes with increased 30 adaptation capacity that not only reduces vulnerability of ecosystems and communities (Douwes et al., 31 2016), but creates economic opportunities. These include direct job creation (Diederichs and Roberts, 2016; 32 Douwes and Buthelezi, 2016) with various spinoffs such as better education for schoolchildren (Douwes et 33 al., 2015). Indirect benefits, include better water quality and reduced flooding, are generated as a result of 34 improved ecosystem service delivery (Douwes and Buthelezi, 2016). In areas that are already developed, 35 opportunities for green roof infrastructure can yield reductions in roof storm water run-off (by approximately 36 60 ml/m²/minute during a rainfall event), slow release of water over time, and reduced temperatures on roof 37 surfaces (Roberts et al., 2012). 38

- 40 2.6.5.8 Case Study: Protecting Gondwanan refugia against fire in Tasmania, Australia
- 4142 Scale: Local
- 43 Issue: Protection of rare endemic species

The Tasmanian Wilderness World Heritage Area (TWWHA) has a high concentration of 'paleo-endemic' plant species restricted to cool, wet climates and fire free environments, but recent wildfires have burnt substantial stands, which are unlikely to recover (Harris et al., 2018b, Bowman et al., 2021, The 2016 Tasmanian). The fires led to government inquiries and a fire-fighting review, which have suggested changes to management as that climate change will make such fires more likely in the future (Council, 2016; Press, 2016; Council, 2019).

The majority of the TWWHA is managed as a Wilderness Zone, where management is currently carried out in a manner that allows natural processes to predominate. The exclusion of fire from stands of fire-sensitive trees such as the Pencil pine, *Athrotaxis cupressoides*, is part of this management strategy, possible in the past due to the moisture differential and lower flammability of these areas. However, in recent years, the threat posed by extensive and repeated wildfires, and an increasing awareness that fire risk is likely to increase (Fox-Hughes et al., 2014; Love et al., 2017; Love et al., 2019) have meant that more direct

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1 2 3 4	managing the threat will not be su the remote wilderness area, but lin	fficient to protect the paleo-endem	ealisation that a "hands off" approach to ics. Not only is fire-fighting difficult in agers must prioritise where fires will be ate simultaneously.
5 6 7 8 9 10 11 12 13	After wildfires in 2016 caused ext were spent trying to protect the re- including the strategic application lines (Council, 2019). These appro- the high value vegetation in some extent of fire-sensitive vegetation interventions may have adverse eff	ensive damage (Bowman et al., 20) naining stands of Pencil pine durin of long-term fire retardant and the paches are thought to have been eff situations. Impact reports are curre communities that have been affects fects on the values of the TWWHA	20a), significant effort and resources ing the 2019 fires, using new approaches installation of kilometres of sprinkler fective at halting the fire and protecting ently being finalised to quantify the ed. However, there is concern that these A if applied widely, so while research is ant is not being applied to some areas).
14 15 16 17 18 19 20 21	the area, along with their knowled and their connection with the area protect important sites from fire ar (DPIPWE, 2016). Increasingly, th	ge of plants, animals, marine resou as a living and dynamic landscape and ensure that management does no ere is an acknowledgment that the	management as an important value of urces, minerals (ochre and rock sources), e. Fire management planning aims to ot impact Aboriginal cultural values cessation of traditional fire uses has led knowledge into fire management of the
22 23 24	2.6.5.9 Case Study: Bhojtal Lak	e, Bhopal, India	
24 25 26 27	Scale: Local Issue: Protection of water resource	es and biodiversity	R
28 29 30 31 32 33	man-made lake bordering the city, conservation site with wetlands pr 2006). Bhojtal also provides a wid	for its water supply (Everard et al otected under the Ramsar conventi e range of other benefits to people	on and diverse flora and fauna (WWF,
34 35 36 37	weather patterns (Anonymous, 20 illegal digging of bore wells and u	19), exacerbated by a series of prob nsustainable water extraction/explo	
38 39 40 41	Services (RAWES) approach, incl	uding water quality analysis from	l Assessment of Wetland Ecosystem the lake. Information on the geology, seline Biodiversity Assessment was
42 43 44		d Management Project (JICA, 200	7) was developed with the following
45 46 47 48			; prevention of pollution (sewerage t and management of water quality
49 50 51		tion of small check dams or percol	neering structures particularly at upper lation tanks for recharge purposes in
52 53 54	3. Afforestation initiatives.		
55 56 57	based solutions are more resilient		s improved the lake's health. Nature- ate change. Restoration not only reduced area (Kabisch et al., 2016).

2.6.5.10 Case Study: Addressing Vulnerability of Peat Swamp Forests in South East Asia (SEA)

Scale: Regional 4

Issue: Protecting peatland biodiversity, carbon and ecosystem services from climate change and land 5 degradation 6

7

1

2 3

SEA peatlands have undergone extensive logging, drainage and land-use conversion that have caused habitat 8 loss for endemic species, i.e., orangutan (Pongo spp) (Gregory et al., 2012; Struebig et al., 2015). Prolonged 9 droughts associated with El Niño (Section 4.4.3.2) compound the effects of drainage, leading to large 10 recurrent fires (Langner and Siegert, 2009; Gaveau et al., 2014; Putra et al., 2019). Under RCP 8.5, it is 11 projected that by the end of the century, the annual rainfall will significantly decrease (30%) over SEA, and 12 the number of consecutive dry days will significantly increase (60%) over Indonesia and Malaysia (Supari et 13 al., 2020). Peat degradation and losses to fire result in large GHG emissions (Miettinen et al., 2016) as well 14 as haze pollution that is a trans-boundary problem in the region (Heil et al., 2007). 15 16

Improving resilience to fire and climate change in SEA peatlands through restoration is extremely difficult 17

and presents many challenges. The Indonesian government has tasked the Badan Restorasi Gambut 18

- (Peatland Restoration Agency) to restore peatlands (Darusman et al., 2021; Giesen, 2021). Other local 19
- initiatives exist, such as fire management programmes and restoration projects (Puspitaloka et al., 2020). 20 Since 2016, the Government of Indonesia has rewetted ~380,000 hectares of degraded peatlands mainly 21
- through canal blocking and flooding, but less than 2000 hectares were successfully restored to native plant 22
- species common to peat swamp forests (Giesen, 2021).Replanting native trees has had relatively low success 23
- (Lampela et al., 2017) because they have a low tolerance to prolonged inundation and a lack of fire 24
- adaptation strategies (Page et al., 2009; Roucoux et al., 2013; Dohong et al., 2018; Cole et al., 2019; Luom, 25 2020; Giesen, 2021). Barriers to successful management are complex, and include the disparity in
- 26 timeframes between ecological restoration and political/socioeconomic needs (Harrison et al., 2020) and an 27
- over-focus on fire-fighting rather than fire prevention (Mishra et al., 2021a). Early protection of peat forests 28
- has been highlighted as a more effective management strategy than restoration, not only in insular SEA but 29 also in areas like Papua New Guinea, which may be targeted for expansion of estate crop plantations (Neuzil 30
- et al., 1997; Dennis, 1999; Anshari et al., 2001; Anshari et al., 2004; Hooijer et al., 2006 Assessment of; Heil 31
- et al., 2007; Page et al., 2009; Page et al., 2011; Posa et al., 2011; International, 2012; Miettinen et al., 2012; 32
- Biagioni et al., 2015; Miettinen et al., 2016; Rieley and Page, 2016; Adila et al., 2017; Cole et al., 2019; 33
- Vetrita and Cochrane, 2019; Harrison et al., 2020; Hoyt et al., 2020; Ruwaimana et al., 2020; Ward et al., 34 2020; Cole et al., 2021).
- 35 36

37

Limits to Adaptation Actions by People 2.6.6

38 The evidence summarised above (Sections 2.6.2 - 2.6.4) shows that by restoring ecosystems it is possible to 39 increase their resilience to climate change, including the resilience of the populations of species they support 40 and of human communities. However, changes to healthy ecosystems and biodiversity are already happening 41 as described in this chapter (robust evidence, high agreement) and further changes are inevitable even under 42 low greenhouse gas emissions scenarios (robust evidence, high agreement). Planning to manage the 43 consequences of inevitable changes and prioritise investments in conservation actions where they have best 44 chance of succeeding (e.g. Section 2.6.4.6) will be an increasingly necessary component of adaptation 45 (robust evidence, high agreement) (Table 2.6). 46

47

It is possible to help species survive by active interventions such as translocation but as described above 48 49 (Section 2.6.4.1) it is not a straightforward process, is not suitable for all species and is resource intensive. Modifying local microclimate or hydrological conditions can work for some species (Sections 2.6.2, 2.6.4.5), 50 but is likely to be less effective at higher levels of climate change. It will also be less successful for larger 51 species and more mobile ones. The microclimate of a tree is much more closely coupled with wider 52 atmospheric conditions than that of a small plant or animal in the boundary layer and mobile species such 53 birds and large mammals range over large areas rather than being confined to discrete locations where 54 conditions can be manipulated. There is potential for using evolutionary changes to enhance the adaptive 55 capacity for target species, such as is being done in the Great Barrier reef by translocating symbionts and

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1	However, known limitations to geneti	c adaptation preclude speci	es-level adaptation to climates beyond their
2	ecological and evolutionary history (S	Sections 2.2.4.6; 2.6.1). All	of these interventionist approaches are
3	constrained by requiring significant fi	nancial resources and expen	tise, they also require a high level of
4	understanding of individual species at	utecology, which can take y	ears to acquire, even if resources were

understanding of individual species autecology, which can take years to acquire, even if resources were
 available. Ex situ conservation (for example in seed banks may be the only option to conserve some species,
 especially as levels of warming increase and this will not be feasible for all.

7

8 While the science of restoration has generated many successes, some habitats are very difficult to restore, 9 making certain decisions effectively irreversible. For example, *Acacia nilotica* was introduced into Indonesia 10 in the 1850s for gum arabic with planting expanded for fire breaks in the 1960s. This tree became invasive 11 and has already replaced >50% of savanna habitat in Baluran National Park, with complete replacement 12 expected in the near future. This shift from savanna to Acacia forest is causing large declines in native

expected in the near future. This shift from savanna to Acacia forest is causing large declines in nati species, including the charismatic wild banteng, *Bos javanicus* and wild dog (dhole, *Cuon alpinus*),

(Caesariantika et al., 2011; Padmanaba et al., 2017; Zahra et al., 2020). Multiple approaches to controlling
the spread of this Acacia have been ineffective, highlighting the difficulty of reversing the decision to plant
this tree (Zahra et al., 2020). Another example is the difficulties in restoring tropical peat forests of South
East Asia (Section 6.5.10).

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EbA when implemented well can reduce risks to people but there are limits, for example, an extreme flood event may exceed the capacity of natural catchments to hold water or slow its flow (Dadson et al., 2017), urban shade trees and green space can make a few degrees difference to temperatures experienced by people but that may not be enough in the hottest conditions.

In general adaptation measures can substantially reduce the adverse impacts of 1-2°C of global temperature rise, but beyond this losses will increase (IPCC, 2018b), including species extinctions and changes, such as major biome shifts which are cannot be reversed on human timescales. Some adaptation measures will also become less effective at higher temperatures. Whilst adaptation is essential to reduce risks, it cannot be regarded as a substitute for effective climate change mitigation (*robust evidence, high agreement*).

30 2.6.7 Climate Resilient Development

Climate Resilient Development (CRD) is the subject of Chapter 18. This section briefly assesses some of the 32 fundamental issues for CRD relating to ecosystems; an overview of the importance of specific ecosystem 33 services for CRD is presented in Box 18.7 in Chapter 18. A large body of evidence has demonstrated the 34 extent to which human life, well-being and economies are dependent on healthy ecosystems and the range of 35 threats they are under (high confidence) (IPBES, 2019; Dasgupta, 2021; Pörtner et al., 2021). An analysis of 36 64 studies found a strong positive synergy among eight critical regulating services of healthy ecosystems, 37 including climate regulation, water provisioning, pest and disease control and adjacent crop pollination (Lee 38 and Lautenbach, 2016). Health of ecosystems is, in turn, reliant upon maintenance of natural levels of 39 species' richness and functional diversity (high confidence) (Lavorel et al., 2020) (see Section 2.5.4). A 40 meta-analysis of 74 studies documented the mechanism for increased ecosystem stability was increased 41 asynchrony among species, that itself was a product of higher species diversity (Xu, 2021, consistently 42 positive effect). Responding to these threats requires the protection and restoration of natural and semi-43 natural ecosystems, together with sustainable management of other areas. 44

45

The Convention on Biological Diversity set the Aichi 2020 target of 17% of each country to be protected for 46 biodiversity. Analyses suggest that 30% or even 50% of land and sea needs to be protected or restored to 47 confer adequate protection for species and ecosystem services (high confidence) (Pimm et al., 2018; 48 Dinerstein et al., 2019) (Woodley et al., 2019; Brooks et al., 2020; Hannah et al., 2020; Luther et al., 2020; 49 Zhao et al., 2020; Sala et al., 2021). Hannah (2020) estimated that limiting warming to 2°C and protecting 50 30% of high biodiversity regions (Africa, Asia and Latin America) reduced risk of species' extinctions in 51 half (medium confidence). Placement of protected areas is as important as total area (Pimm et al., 2018), and 52 quality of protection (strictness and enforcement) is as important as the official land designation (Shah et al., 53 2021). Pimm et al. (Pimm et al., 2018) found that many small protected areas are successful because they are 54 in areas of very high biodiversity containing species of small ranges size, while many large regions identified 55

as wild often are of low biodiversity value, though they may have high mitigation value (e.g. high Arctic
 tundra). Finally, a global meta-analysis of 89 restoration projects, biodiversity increased by 44% and

ecosystem services by 25% after restoration, but values remained lower than in intact reference systems 1 (Benayas José et al., 2009). 2

3 There is also increasing evidence, reported in this chapter, that the loss and degradation of natural and semi-4 natural habitats exacerbates the impacts of climate change and climatic extreme events on biodiversity and 5 ecosystem services (high confidence); example references include: (Ogden et al., 2013; Eigenbrod et al., 6 2015; Struebig et al., 2015; Stevens et al., 2016; Oliver et al., 2017; McAlpine et al., 2018; Taffarello et al., 7 2018; Lehikoinen et al., 2019; Birk et al., 2020; Chapman et al., 2020; Agol et al., 2021; Khaniya et al., 8 2021; Lara et al., 2021; Lehikoinen et al., 2021). Considering these two sets of evidence together, it is clear 9 that climate change adaptation and ecosystem degradation both need to be addressed if either is to be tackled 10successfully (robust evidence, high agreement) as a number of recent publications have concluded (Haddad 11 et al., 2015; Hannah et al., 2020; Arneth et al., 2021; Pörtner, 2021). Taking this combined body of evidence 12 together, this assessment is that the protection and restoration of natural and semi-natural ecosystems are key 13 adaptation measures (robust evidence, high agreement) (Section 2.5.4).

14 15

Large scale protection and restoration of ecosystems can also make a significant contribution to climate 16

- change mitigation (Dinerstein et al., 2020; Roberts et al., 2020a; Soto-Navarro et al., 2020). Globally, there 17 is a 38% overlap between areas of high carbon storage and high intact biodiversity (mainly in the peatland
- 18tropical forests of Asia, Western Amazon and the high Arctic), but only 12% of that is protected (high 19
- confidence) (Soto-Navarro et al., 2020). Peatlands are particularly important carbon stores but are threatened
- 20 human disturbance, land use change (Leifeld et al., 2019) and fire (Turetsky et al., 2015). Restoration of 21
- peatlands is not only an efficient nature climate solution in terms of GHG (Nugent et al., 2019), but it may
- 22 also increase ecosystem resilience (Glenk et al., 2021). Global restoration efforts are ongoing to target 23
- degraded temperate peatlands in the American and Europe (Chimner et al., 2017), as a result of their 24
- importance for climate change mitigation being recognised (Paustian et al., 2016; Bossio et al., 2020; 25
- Humpenöder et al., 2020; Drever et al., 2021; Tanneberger et al., 2021). It has been estimated that the global 26 GHG saving potential of peatland restoration is similar to the most optimistic sequestration potential from all 27 agricultural soils (Leifeld and Menichetti, 2018). However the pressure on peatlands from human activity 28
- remains high in many parts of the world (Humpenöder et al., 2020; Tanneberger et al., 2021). Currently, the 29
- rapid destruction of tropical peatlands overshadows any current restoration efforts in temperate peatlands or 30 any potential carbon gain from natural high-latitude peatlands (Roucoux et al., 2017; Wijedasa et al., 2017; 31
- Leifeld et al., 2019). (Sections 2.4.3.8, 2.4.4.4.2, 2.4.4.4, 2.5.2.8, 2.5.3.4) 32 33

Recent studies have highlighted the importance of ensuring that ecosystem protection is not implemented in 34 a way which disadvantages those who live in or depend on the most intact ecosystems (Mehrabi et al., 2018; 35 Schleicher et al., 2019) or risk food security. The actual area of land to be protected and the balance between 36 sustainable use and protection will need careful planning and targeting to where it can have most benefit 37 (Pimm et al., 2018); it will also be important to ensure that protection measures are effective in preventing 38 damage (Shah et al., 2021). 39

40

At a local level, EbA can often provide a wide range of additional benefits for sustainable development in 41 both rural and urban areas (Wilbanks, 2003; Nelson et al., 2007; Cohen-Shacham et al., 2016; Hobbie and 42 Grimm, 2020; Martín et al., 2020). A number of the case studies above, such as those in Durban and at 43 Bhojtal Lake illustrate this (2.6.5). A key element of Climate Resilient Development is ensuring that actions 44 taken to mitigate climate change do not compromise adaptation, biodiversity and human needs. This depends 45 on choosing appropriate actions for different locations (Box 2.2, Cross-Chapter Box NATURAL this 46 Chapter). A particularly notable case of this is woodland creation as described in Box 2.2: re-afforestation of 47 previously forested areas can provide multiple benefits (Lee et al., 2018; Lee et al., 2020) including for 48 49 climate change mitigation, adaptation and biodiversity. However planting trees where they would not naturally grow can create multiple problems include the loss of native biodiversity and disruption of 50 hydrology (Box 2.2). It is also the case that protection of existing natural forest ecosystems is the highest 51 priority for reducing greenhouse gas emissions (Moomaw et al., 2019) and restoration may not always be 52 practical (see Section 2.6.4.10). (Sections 2.4.3.6, 2.4.3.7, 2.4.4.3, 2.4.4.4, 2.5.2.6, 2.5.2.7, 2.5.3.3, Box 2.2, 53 Cross-Chapter Box NATURAL this Chapter) 54

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In some cases actions supported by international donors and presented as addressing climate change 56 adaptation and mitigation in the natural environment can have damaging consequences for people and nature 57

as well as failing to deliver adaptation and mitigation. One example of this was presented by (Work et al., 1 2019) who reviewed three climate change mitigation and adaptation projects in Cambodia: an irrigation 2 project, a protected-area forest management project, and a reforestation project. In each case they found 3 evidence of local communities rights being violated, maladaptation and destruction of biodiverse habitats. 4 They concluded that the potential for maladaptation and adverse social and environmental impacts had been 5 ignored by international donors as well as national authorities and that there was a need for much strict 6 accountability mechanisms. Moyo et al. (Moyo et al., 2021), using case studies from South Africa, 7 documented higher success of ecosystem restoration projects when they embraced broader SDGs, 8 particularly enhancement of people's livelihoods. Better assessment of the impacts of adaptation and 9 mitigation measures on people and ecosystems, before they are implemented, will be increasingly necessary 10 to avoid unintended and damaging consequences as their deployment is scaled up (Larsen, 2014; Enríquez-11 de-Salamanca et al., 2017; Pour et al., 2017). This applies to ostensibly nature-based approaches as well as 12 more engineering-based ones. 13 14 Another aspect of the benefits to people from ecosystems that needs to be taken into account in Climate 15 Resilient Development is increasingly strong evidence of the benefits of natural environments for human 16 health and well-being beyond the provision of basic necessities, such as food and water (Bratman et al., 17 2019; Marselle et al., 2021). Meta-analyses of 162 studies across 51,738 people documented that individuals 18 with high levels of contact with nature through their lives felt significantly happier, healthier and more 19 satisfied with their lives, and engage in more pro-nature behaviours, than those with little or no contact with 20 nature (high confidence) (Capaldi et al., 2014; Mackay and Schmitt, 2019; Pritchard et al., 2020; Whitburn et 21 al., 2020). Meta-analyses of manipulative human trials across 65 studies document a significant increase in 22 positive feelings and attitudes, and declines in negative feelings, after experimental treatments involving 23 nature (medium confidence) (Bowler et al., 2010b; McMahan and Estes, 2015; Soga et al., 2017). Within the 24 context of CRD improving the extent to which humans see themselves as part of the natural world – known 25 as human-nature connectedness (HNC) – by increasing access to natural areas, particularly within urban 26 areas, can provide additional health, cultural and recreation benefits of NbS as well as increasing public 27 engagement and support (robust evidence, high agreement) (Wilbanks, 2003; Nelson et al., 2007; Bowler et 28 al., 2010b; Capaldi et al., 2014; McMahan and Estes, 2015; Cohen-Shacham et al., 2016; Soga et al., 2017; 29 Mackay and Schmitt, 2019; Work et al., 2019; Hobbie and Grimm, 2020; Pritchard et al., 2020; Whitburn et 30 al., 2020).

31 32 33

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34 [START CCB NATURAL HERE]

36 Cross-Chapter Box NATURAL: Nature-Based Solutions for Climate Change Mitigation and 37 Adaptation

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52 Nature-based Solutions provide adaptation and mitigation benefits for climate change as well as

contributing to other sustainable development goals (high confidence). Effective nature-based climate

54 *mitigation stems from inclusive decision-making and adaptive management pathways that deliver climate-*

resilient systems serving multiple sustainable development goals. Robust decision-making adjusts management pathways as systems are impacted by on-going climate change. Poorly conceived and

designed nature-based mitigation efforts have the potential for multiple negative impacts, including

competing for land and water with other sectors, reducing human well-being and failing to provide 1 mitigation that is sustainable in the long-term (high confidence). 2 3 The concept of Nature-based Solutions (NbS) is broad and debated but has become prominent in both the 4 scientific literature and policy since AR5, including earlier concepts, including Ecosystem based Adaptation 5 (EbA). The key point is that these are actions benefiting both people and biodiversity (IUCN 2020; WGII 6 Glossary). In the context of climate change, NbS provide adaptation and mitigation benefits for climate 7 change in ways that support wild species and habitats, often contributing to other sustainable development 8 goals (robust evidence, high agreement) (Keesstra et al., 2018; Lavorel et al., 2020; Malhi et al., 2020) 9 (Griscom et al., 2017; Hoegh-Guldberg et al., 2019; IPCC, 2019a; Lewis et al., 2019; Seddon et al., 10 2020b)(AR6 WGIII Chapter 12, see Sections 2.2, 2.5.4, 2.6.3, 2.6.5, 2.6.7). Well-designed and implemented 11 NbS mitigation schemes can increase carbon uptake or reduce greenhouse gases emissions at the same time 12 as protecting or restoring biodiversity and incorporating elements of food provisioning (Mehrabi et al., 13 2018). A variety of measures can be part of NbS, ranging from the protection of natural terrestrial, 14 freshwater and marine ecosystems, the restoration of degraded ones (this Cross-Chapter Box; Section 13.3), 15 to more sustainable management of naturally regenerating ecosystems used for food, fibre and energy 16 production (Figure Cross-Chapter Box NATURAL 1, Chapter 5, Cross-Working Group Box 17 BIOECONOMY in Chapter 5). Agroecological practices mitigate and adapt to climate change and can 18 promote native biodiversity (high confidence) (Sinclair et al., 2019; Snapp et al., 2021). 19 20 The role of restoration in NbS 21 22 Where natural ecosystems have been degraded or destroyed, re-establishing them and restoring natural 23 processes can be a key action for adaptation and mitigation, and the science of restoration is well-established 24 (de los Santos et al., 2019; Duarte et al., 2020) (Section 13.4.1). Such restoration activities need to adapt to 25

on-going climate change risks for the landscape and ocean scape and the species composition of biological
 communities. Indeed, climate-change impacts may overwhelm attempts at restoration/conservation of
 previous or existing ecosystems, particularly when the ecosystem is already near its tipping point, as are
 tropical coral reefs (Bates et al., 2019; Bruno et al., 2019).

30

Lands (e.g. forests) and oceans (e.g. fisheries) managed for products using sustainable practices (whether applied by private, state, or indigenous peoples) can also be carbon- and biodiversity-rich, and so effective NbS (Paneque-Gálvez et al., 2018; Soto-Navarro et al., 2020). Indigenous people and private forest owners manage, use or occupy at least one-quarter of global land area, over one-third of which overlaps with protected areas, thus combining both protection and production (Jepsen et al., 2015; Garnett et al., 2018; IPBES, 2019; Santopuoli et al., 2019).

37 Protection/restoration of natural systems, including reducing non-climate stressors, and sustainable 38 management of semi-natural areas emerge as necessary actions for adaptation to minimise extinctions of 39 species, reaching tipping points that cause regime shifts in natural systems, loss of whole ecosystems, and 40 their associated benefits for humans (Scheffer et al., 2001; Folke et al., 2005; Luther et al., 2020) (Chapters 2 41 and 3, AR6 WGIII Chapter 7). Such measures are critical for conservation of biodiversity and the provision 42 of ecosystem goods and services in the face of projected climate change (Duarte et al., 2020). Supporting 43 local livelihoods and providing benefits to indigenous, local communities and millions of private 44 landowners, together with their active engagement in decision-making, is critical to ensuring support for 45 NbS and its successfully delivery (high confidence) (Chapter 05; Figure Cross-Chapter Box NATURAL 1; 46 Ceddia et al., 2015; Blackman et al., 2017; Nabuurs et al., 2017; Smith et al., 2019a; Smith et al., 2019b; 47 Jones et al., 2020a; McElwee et al., 2020; Cao et al., 2021). 48

50 Forests

51

49

Intact natural forest ecosystems are major stores of carbon and support large numbers of species that cannot survive in degraded habitats *(very high confidence)*. Extensive areas of natural forest ecosystems remain in

- 54 tropical, boreal and (to a lesser extent) temperate biomes regions, but in many regions, they are managed
- (sustainably and unsustainably) or have been degraded or cleared. Deforestation and degradation continue to
- be a source of global greenhouse gas emissions (very high confidence) (Friedlingstein et al., 2019).
- 57 Protecting existing natural forests and sustainable management of semi natural forests providing goods and

services is a highly effective NbS (Bauhus et al., 2009) (high confidence). Natural forests and sustainably 1 managed diverse forests play important roles for climate change mitigation and adaptation while providing 2 many other ecosystem goods and services (very high confidence) (Bradshaw and Warkentin, 2015; Favero et 3 al., 2020; Mackey et al., 2020). Contributions to climate change mitigation are estimated at medians of 5-7 4 Gt CO₂/y (Roe et al., 2019). Forests influence the water cycle at local, regional and global scales (Creed and 5 van Noordwijk, 2018) reducing surface runoff, increasing infiltration to groundwater and improving water 6 quality (Bruijnzeel, 2004; Zhou et al., 2015a; Ellison et al., 2017; Alvarez-Garreton et al., 2019). Recent 7 evidence shows that downwind precipitation is also influenced by evapo-transpiration from forests (Keys et 8 al., 2016; Ellison et al., 2017). Protecting existing natural forest and sustainably managing production 9 forests, in a holistic manner, can optimise the provision of the many functions forests fulfil for owners, 10 conservation, mitigation and for society as a whole (Bauhus et al., 2009; Nabuurs et al., 2013). 11 12 Reforestation of formerly forested land can help to protect and recover biodiversity and can be one of the

Reforestation of formerly forested land can help to protect and recover biodiversity and can be one of the most practical and cost effective ways of sequestering and storing carbon *(high confidence)* (Nabuurs et al.,

15 2017; Hoegh-Guldberg et al., 2018b; Paneque-Gálvez et al., 2018; Smith et al., 2018; Cook-Patton et al.,

¹⁶ 2020; Cowie et al., 2021; Drever et al., 2021). This can be achieved through planting or by allowing natural

colonisation by tree and shrub species. The most effective method to employ depends upon local
 circumstances (such as the presence of remnant forest cover) or socio-cultural and management objectives.

circumstances (such as the presence of remnant forest cover) or socio-cultural and management objectives.
 Reforestation with climate-resilient native or geographically near species restores biodiversity at the same

time as sequestering large amounts of carbon (Lewis et al., 2019; Rozendaal et al., 2019). It can also restore

hydrological processes, improving water supply and quality (Ellison et al., 2017) and reducing risks of soil

erosion and floods (high confidence) (Locatelli et al., 2015).

23

Climate change may mean that in any given location, different species will be able to survive and become 24 dominant, and restoring the former composition of forests may not be possible (Sections 2.4; 2.5). Severe 25 disturbances such as insect/pathogen outbreaks, wildfires, and droughts, which are an increasing risk, can 26 cause widespread tree mortality resulting in sequestered forest carbon being returned to the atmosphere 27 (Anderegg et al., 2020; Senf and Seidl, 2021) thus suggesting we need to adapt (Sections 2.4, 2.5, 13.3) 28 14.4.1, Box 14.1). Adaptation measures, such as increasing the diversity of forest stands through ecological 29 restoration rather than monoculture plantations can help to reduce these risks (confidence). When plantations 30 are established without effective landscape planning and meaningful engagement including free prior and 31 informed consent, they can present risks to biodiversity and the rights, well-being and livelihoods of 32 Indigenous and local communities, as well as being less climate-resilient than natural forests (very high 33 confidence) (Section 5.6; Corbera et al., 2017; Mori et al., 2021). 34

35

Afforesting areas such as savannahs and many temperate peatlands, which would not naturally be forested, damages biodiversity and increase vulnerability to climate change *(high confidence)* so is not a Nature-based Solution and can exacerbate greenhouse gas emissions (Sections 2.4.3.5, 2.5.2.5, Box 2.2 this Chapter). Remote sensing based assessments of suitability for tree planting can over-estimate potential, due to failure to adequately distinguish between degraded forest and naturally open areas (Bastin et al., 2019; Veldman et al., 2019; Bastin et al., 2020; Sullivan et al., 2020).

43 **Peatlands**

44 Peatlands are naturally high-carbon ecosystems, which have built up over millennia. Draining, cutting and 45 burning peat lead to oxidation and the release of CO₂ (very high confidence). Rewetting by blocking 46 drainage, preventing cutting and burning can reverse this process on temperate peatlands (medium 47 confidence) although can take many years (Bonn et al., 2016). Trees are naturally found on many tropical 48 49 peatlands and restoration can involve removing non-native species such as oil palm and re-establishing natural forest. However, peatland tropical forest is difficult to fully restore, and native pond fish, that are 50 vital as a local food, often do not return. Protection of intact peat forests, rather than attempting to restore 51 cleared forest, is by far the more effective pathway both in terms of cost, CO₂ mitigation, and protection of 52 food sources (Kreft and Jetz, 2007). Naturally treeless temperate and boreal peatlands have in some cases 53 been drained to enable trees to be planted, which leads to CO₂ emissions, and restoration requires removal of 54 trees as well as re-blocking drainage. (high confidence) (Sections 2.4.3.8; 2.5.4.8; 2.6.5.10). 55 56

57 Blue Carbon

1 Blue Carbon ecosystems (mangroves, saltmarshes and seagrass meadows; see glossary) often have high rates 2 of carbon accumulation and sequestration (Section 3.5.5.5; Macreadie et al., 2019). However, quantification 3 of their overall mitigation value is difficult due to variable production of CH_4 and N_2O (Adams et al., 2012; 4 Rosentreter et al., 2018; MacLean et al., 2019b), uncertainties regarding the provenance of carbon 5 accumulated (Macreadie et al., 2019), and the release of CO_2 by biogenic carbonate formation in seagrass 6 ecosystems (Saderne et al., 2019). Therefore, blue carbon strategies, referring to climate change mitigation 7 and adaptation actions based on conservation and restoration of blue carbon ecosystems, can be effective 8 NbS, with evidence of recovery in carbon stocks following restoration, although their global or regional 9 carbon sequestration potential and net mitigation potential may be limited (medium confidence) (Sections 10 3.6.3.1.6; 13.4.3, AR6 WGI 5.6.2.2.2; Duarte et al., 2020). They can also significantly attenuate wave 11 energy, raise the seafloor thus counteracting sea level rise effects, and buffer storm surges and flooding 12 erosion (high confidence) (Sections 13.2.2; 13.10.2). Additionally, they provide a suite of cultural (for 13 example, tourism, livelihood and well-being for native and local communities), provision (e.g. mangrove 14 woods, edible fish and shellfish) and regulation (e.g. nutrient cycling) services (high confidence) (Section 15 3.5.5.5). These services have motivated the implementation of management and conservation strategies of 16 these ecosystems (Sections 3.6.3.1.6; 13.4.2). Blue carbon strategies are relatively new, with many of them 17 experimental and small scale; therefore there is *limited evidence* of their long-term effectiveness. There is 18 also limited information on the potential emission of other GHGs from restored blue carbon ecosystems, 19 although reconnecting hydrological flow in mangroves and saltmarsh restoration are effective interventions 20 to reduce CH₄ and CO₂ (*limited evidence, medium agreement*) (Kroeger et al., 2017; Al-Haj and Fulweiler, 21 2020). 22

24 Urban NbS

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Nature-based Solutions can be a key part of urban climate adaptation efforts. Direct human adaptation 26 benefit may stem from the cooling effects of urban forests and green spaces (parks and green roofs), from 27 coastal wetlands and mangroves reducing storm surge and flooding, and from sustainable drainage systems 28 designed to reduce surface flooding from extreme rainfall, as well as general benefits to human health and 29 well-being (high confidence) (Sections 2.2; 2.6; Chapter 6; Frantzeskaki et al., 2019; Keeler et al., 2019) 30 (Kowarik, 2011). Not all green schemes are considered "Nature Based Solutions" if they do not benefit 31 biodiversity, but carefully designed urban greening can be effective NbS. Careful planning also helps limit 32 negative equity consequences, benefiting wealthy neighbourhoods more than poor (Geneletti et al., 2016; 33 Pasimeni et al., 2019; Grafakos et al., 2020). Effective planning should also consider what is appropriate for 34 the climate and conditions of each city. For example, some trees emit volatiles (e.g. isoprene) that, in the 35 presence of certain atmospheric pollutants, can increase surface ozone that in turn can cause human 36 respiratory problems (Kreft and Jetz, 2007). Wetlands restoration close to human settlements needs to be 37 paired with mosquito control to prevent negative impacts on human health and well-being (Stewart-Sinclair 38 et al., 2020), but has been shown to provide better filtration and toxicity reduction with lower environmental 39 impact than other forms of waste-water treatment (Vymazal et al., 2021), including "green roofs" and "green 40 walls" (Chapter 06; Addo-Bankas et al., 2021). 41 42

43 Agroecological Farming

Agroecological farming (AF) is a holistic approach that incorporates ecologic and socio-economic
principles. It strives to enhance biodiversity, soil health and synergies between agroecosystem components,
reduce reliance on synthetic inputs (e.g., pesticides), builds on Indigenous knowledge and local knowledge,
and fosters social equity (e.g., supporting fair, local markets (HLPE, 2019; Wezel et al., 2020). AF practices
include intercropping, mobility of livestock grazing across landscapes, organic agriculture, integration of
livestock, fish and cropping, cover crops and agroforestry. (Sections 5.14; FAQ 12.5, 13.5.)

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Agroforestry, cover crops and other practices that increase vegetation cover and enhance soil organic matter, carefully management and varying by agroecosystem, mitigate climate change (*high confidence*) (Zomer et al., 2016; Aryal et al., 2019; Nadège et al., 2019). Global meta-analyses demonstrate agroforestry storing 20 -33% more soil carbon than conventional agriculture (De Stefano and Jacobson, 2018; Shi et al., 2018) and reducing the spread of fire (Sections 5.6, 13.5.2, 7.4.3, Box 7.7). Minimising synthetic inputs such as Nbased fertilisers reduces emissions (Gerber et al., 2016). Cover crops can reduce N₂O emissions and increase

soil organic carbon (Abdalla et al., 2019). Conservation farming (no-till with residue retention and crop
rotation) increases soil organic carbon particularly in arid regions (Sun et al., 2020). Silvo-pastoral systems
(pastures with trees), and other practices that increase vegetation cover and enhance soil organic matter
increase sequestered carbon in vegetation and soils (Zomer et al., 2016; Aryal et al., 2019; Nadège et al.,
2019; Ryan et al. 2019). Agroecologically improved cropland and grazing land management have significant
mitigation potential, estimated at 2.8- 4.1 GtCO₂e per year (Smith et al., 2020). (Sectiond 5.10, 5.14, Box
5.10, Cross Working-Group Box BIOECONOMY in Chapter 5; WGIII 7.4.3; Box 7.7).

8

AF enhances adaptation to climate change, including resilience to extreme events. Building organic matter 9 improves soil water-holding capacity and buffers against drought; increased perenniality and high levels of 10 ground cover reduce soil erosion during storms; agroforestry shelters stock and crops in heat waves; 11 landscape complexity and agrobiodiversity increase resilience to disease and pests and stabilise livestock 12 production and restoration of oyster reefs provides thermal refugia and storm surge protection (Allred et al., 13 2014; Henry et al., 2018; Kremen and Merenlender, 2018; Beillouin et al. 2019; ; Kuyah et al. 2019; Gilby 14 et al., 2020; Niether et al. 2020; Richard et al. 2020; Howie and Bishop, 2021; Snapp et al. 2021) (;). 15 Livestock mobility enables adjusting to increased climatic variability while maintaining pastoral systems' 16 productivity (Turner and Schlecht, 2019; Scoones, 2020). Thus, adoption of agroecology principles and 17 practices will be highly beneficial to maintaining healthy, productive food systems under climate change 18 (high confidence) (Sections 5.4.4; 13.5.2; FAQ 12.4). 19

20 AF practices such as hedgerows and polycultures maintain habitat and connectivity for biodiversity and 21 support ecosystem functioning under climate stress compared to conventional agriculture (high confidence), 22 Section 5.4.4.4; Buechley et al., 2015; Kremen and Merenlender, 2018; Albrecht et al., 2020). Increasing 23 farm biodiversity benefits pollination, pest control, nutrient cycling, water regulation and soil fertility (Snapp 24 et al. 2021; Beillouin et al. 2019; Tamburini et al. 2020). Biodiverse agroforestry systems increase ecosystem 25 services and biodiversity benefits compared to simple agroforestry and conventional agriculture (high 26 confidence); up to 45% more biodiversity and 65% more ecosystem services compared to conventional 27 timber, crop or livestock production in the Brazilian Atlantic Forest (Santos et al., 2019), including for birds 28 (M. Greenler and Ebersole, 2015; Karp et al., 2019), local tree species (Braga et al., 2019) and fewer 29 invasive exotic plants species (Cordeiro et al., 2018). AF includes conservation of semi-natural woodlands, 30 which can conserve bird predators of insect pests (Gonthier et al., 2019). Organic production increases insect 31 species richness and abundance in and around the farm, including essential pollinators (Sections 5.10, 12.6; 32 Kennedy et al., 2013; Haggar et al., 2015; Lichtenberg et al., 2017). 33

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AF significantly improves food security and nutrition by increasing access to healthy, diverse diets and rising incomes for food producers, through increased biodiversity of crops, animals, and landscapes (*high confidence*) (Garibaldi et al., 2016; D'Annolfo et al., 2017; Isbell et al., 2017; Dainese et al., 2019; Bezner Kerr et al. 2021). Livestock mobility improves the site-specific matching of animals' needs with food availability (Damonte et al., 2019; Mijiddorj et al., 2020; Postigo, 2021), and can generate a form of rewilding that restores lost ecosystem functioning (Gordon et al., 2021). Conservation of crop wild relatives in situ supports genetic diversity in crops for the range of future climate scenarios (Redden et al., 2015).

System-level agroecological transitions require policy support for farmer experimentation and knowledge 43 exchange, community-based participatory methodologies and market and policy measures e.g. public 44 procurement, local and regional market support, regulation or payments for environmental services (HLPE 45 2019; Snapp et al. 2021; Mier y Teran et al. 2018). Scientific consensus about the food security and 46 environmental implications of agroecological transitions at a global scale is lacking. Yields in agroforestry 47 and organic can be lower than high-input agricultural systems, but conversely, AF can boost productivity and 48 49 profit, varying by timeframe, socio-economic, political and ecosystem context (medium confidence) (Section 5.14; Muller et al., 2017; LaCanne, 2018; Barbieri et al., 2019; Rosa-Schleich et al. 2019; Smith et al., 50 2019b; Smith et al., 2020). :. Such contrasting results and the limited investment in agroecological research 51 to date make paramount assessing the global and regional impacts of agroecological transitions on food 52 production, ecosystems and economy (Section 5.14; DeLonge et al., 2016; Muller et al., 2017; Barbieri et al., 53 2019). 54

- 55
- 56 Conclusions
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Nature-based Solutions provide adaptation and mitigation benefits for climate change as well as contributing 1 to other sustainable development goals (high confidence). NbS avoid further emissions and promote CO_2 2 removal using approaches that yield long-lasting mitigation benefits and avoid negative outcomes for other 3 sustainable development goals. Poorly conceived and designed mitigation efforts have the potential for 4 multiple negative impacts: (1) They can have cascading negative effects on long-term mitigation by 5 promoting short term sequestration over existing long-term accumulated carbon stocks, (2) They can be 6 detrimental to biodiversity, undermining conservation adaptation, and (3) They can erode other ecosystem 7 services important for human health and well-being (high confidence). Conversely, well-designed and 8 implemented mitigation efforts have the potential to provide co-benefits in terms of climate-change 9 adaptation, as well as multiple goods and services, including conservation of biodiversity, clean and 10 abundant water resources, flood mitigation, sustainable livelihoods, food and fibre security, and human 11 health and well-being (high confidence). A key aspect to such 'smart' climate mitigation is implementation of 12 inclusive and adaptive management pathways (Section 1.4.2). These entail acceptance of the inherent 13 uncertainty in projections of future climate change, especially at a regional or local level, and using decision 14 making processes that keep open as many options as possible, for as long as possible, with periodic re-15 evaluation to aid in choosing pathways forward even as systems are being impacted by on-going climate 16 change (Figure Cross-Chapter Box NATURAL 1; Cross-Chapter Box DEEP in Chapter 17; Section 1.4.2). 17 18

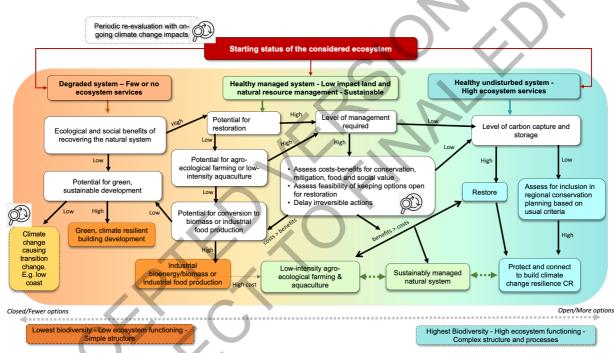


Figure Cross-Chapter Box NATURAL.1: Decision-making framework to co-maximise adaptation and mitigation benefits from natural systems. Decision-making pathways are designed to add robustness in the face of uncertainties in future climate change and its impacts. Emphasis is on keeping open as many options as possible, for as long as possible, with periodic re-evaluation to aid in choosing pathways forward even as systems are being impacted by on-going climate change

Table Cross-Chapter Box NATURAL.1: Assessment of benefits and tradeoffs between mitigation and strategies for both biodiversity and human adaptation to future climate change. Best practices highlight approaches that lead to maximal positive synergies between mitigation and adaptation; worst practices are those most likely to lead to negative tradeoffs for adaptation. Many best practices have additional societal benefits beyond adaptation, such as food provisioning, recreation and improved water quality. Mitigation Potential (Mit. Pot.) and Restoration Potential (Rest. Pot.) are considered.

System	Mit. Pot. Res	t. Pot. Best practices and adaptation benefits	Worst practices and negative adaptation tradeoffs	Additional societal benefits	References
Forests					

FINAL DRAI	FT		Chapter	2 IPCC	C WGII Sixth Asses	sment Report
Boreal Forests	Medium	Medium	Maintain or restore species and structural diversity, reduce fire risk, spatially separate wood production, and sustainably intensify management in some regions	Very large scale clear cuts, aiming for one or few tree species, although boreal is characterised by few tree species and a natural fire risk	Providing goods and services, improved air quality, improved hydrology, jobs	Drever et al. (2021)
Temperate forests	Very high	High	Maintain or restore natural species and structural diversity, leading to more biodiverse and resilient system	native monocultures which	Providing goods and services, jobs, improved hydrology and biodiversity	2.5; Box 2.2;
Tropical wet forests	High	Moderate	Maintain or restore natural species and structural diversity, high biodiversity, more resilient to climate change	monocultures, loss of biodiversity, poor climate change resilience, soil	Indiginous foods, medicines and other forest products, including sustainable selective logging	Section 2.4.3 Edwards et al. (2014)
Tropical dry forests	High	Moderate	Integrated landscape management	Planting non-native monocultures, Loss of biodiversity, poor climate change resilience, soil erosion		Foli et al. (2018)
Tropical peatland forests	Very high	Low	Integrated landscape management	Cutting native rainforest and planting palm oil for biodiesel results in very high carbon emissions from exposed peat soils	major food for local communities	
Blue Carbon						AR6 WGI 5.6.2.2.2
Mangroves	Moderate	High	Conservation, restoration of hydrological flows, revegetation with native plants, livelihoods diversification, landscape planning for landward and upstream migration	C	Improved fisheries and biodiversity, coastal protection against SLR and storm surges, recreation and cultural benefits	Sections 3.4.2.5; 3.5.5.5; 3.6.3.1; Macreadie et al. (2019); Duarte et al. (2020); Sasmito et al. (2020)
Saltmarshes	Moderate	High	Conservation, reduce nutrient loads, restoration of hydrological flows and sediment delivery, revegetation with native plants, landscape planning for landward and upstream migration	Potential NH4 emissions	Improved fisheries and biodiversity, protection against SLR and storm surges, recreational and cultural benefits	Sections 3.4.2.5; 3.5.5.5; 3.6.3.1; Macreadie et al. (2019); Duarte et al. (2020)
Seagrasses	Moderate	High	Conservation; restoration; improve water quality and reduce local stressors (reduction of industrial sewage, anchoring and trawling regulation)	Potential NH4 emissions	Improved fisheries and biodiversity; protection from shoreline erosion; recreational benefits	Section 3.4.2.5; 3.5.5.5; 3.6.3.1; de los Santos et al. (2019); Macreadie et al. (2019); Duarte et al. (2020)
Urban Ecosyst	ems					
Urban forests	Moderate to High*	Moderate	Integrated landscape management. Species richness (including exotics) can be high	monoculture of an exotic tree lowers resilience and reduces biodiversity	Recreation & aesthetics; stormwater absorption benefits;	WGII Chapter 06

			*			
					heat mitigation, air quality improvements	
Urban wetlands	Moderate*	Moderate	Integrated landscape management.		Recreation & aesthetics; stormwater absorption; heat mitigation; coastal flood protection	WGII Chapter 06
Urban grasslands	Moderate*	Moderate	Integrated landscape management	fertilized commercial grass monocultures often require irrigation and are less resilient to droughts than native, mixed grasses and forbs	Recreation & aesthetics; stormwater absorption; heat mitigation	WGII Chapter 06
Open grasslan savanna	ds &					6
	High	Moderate	Blocking drainage channels; Raise water level to natural condition; remove planted trees; revegetation of bare peat; No burns; Increases biodiversity resilience; Reduce flood risk	Inappropriate hydrological restoration, e.g., flood surface depth greater than natural depth leading to methane emissions	Improved water quality in some conditions.	Sections 2.4.3; 2.5; Bonn et al. (2016); Nugent (2019); Taillardat et al. (2020)
Tropical savannas and grasslands (including rangelands)	Moderate	High	Control of feral herbivores; Reintroduce indigenous burns; reintroduce native Herbivores, controlled grazing; strategic design of water-holes; Community-based natural resource management, grass reseeding, clearing of invasive and encroaching woody plants	Afforestation, over- grazing/stocking; No burns; inappropriate placement and design of watering points. All leads to loss of biodiversity, and resilience; soil erosion; water insecurity	production, sustainable wildlife	Sections 2.4.3; 2.5; Box 2.1; Stafford et al. (2017); Moura et al. (2019); Shackelford et al. (2021); Stringer et al. (2021); Wilsey (2021)
Temperate Grasslands and rangelands	Moderate to High	Moderate to High	Integrated landscape management; sustainable grazing; Community- based natural resource management; Native grassland species more resistant to drought than introduced species	Monoculture of introduced species; over-fertilising with chemical or organic amendments; Failure to manage human-wildlife clashes; Failure to distribute income equitably; inadequate enabling policy to facilitate integrated landscape management	Sustainable harvest of wildlife, livestock and dairy production, wild fruits, medicinal plants, construction material, fuelwood; income from ecotourism	Sections 2.4.3; 2.5, Box 2.1; Farai, (2017); Baker et al. (2018); Homewood et al. (2020) Wilsey (2021)
Agroecologial farming and aquaculture	High	High (context specific)	Biodiverse systems at landscape scale; participatory adaptation to context; Short value chains; Farmer incentives; Biodiversity synergies; reduced climate risk	Poorly chosen species, practices and amendments can lead to low yields; Simplified agroforestry systems or industrial scale organic agriculture lacks holistic system-wide approach. Over-fertilising with organic amendments	Food security, human health, livelihoods, socio- cultural benefits e.g. culturally- appropriate foods.	Sections 5.4, 5.10, 5.12, 5.14; Coulibaly et al. (2017); HLPE (2019); Quandt et al. (2019); Sinclair et al. (2019); Smith et al. (2019b); Muchane et al. (2020); Reppin et al. (2020)

[END CROSS-CHAPTER BOX NbS-NATURAL HERE]

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[INSERT FAQ2.5 HERE]

FAQ2.5: How can we reduce the risk of climate change to people by protecting and managing nature better?

5 Damage to our natural environment can increase the risk climate change poses to people. Protecting and 6 restoring nature can be a way to adapt to climate change, with benefits for both humans and biodiversity. 7 Examples include reducing flood risk by restoring catchments and coastal habitats, the cooling effects of 8 natural vegetation and shade from trees and reducing the risk of extreme wildfires by better managing of 9 natural fires. 10

Protecting and restoring natural environments, such as forests and wetlands can reduce the risks climate 12 change poses to people, as well as supporting biodiversity, storing carbon and providing many other benefits 13 for human health and wellbeing. Climate change is bringing an increasing number of threats to people, 14 including flooding, droughts, wildfire, heat waves and rising sea levels. These threats can however be 15 reduced or aggravated, depending on how land, sea and freshwater are managed or protected. There is now 16 clear evidence that 'Nature-based Solutions' (NbS) can reduce the risks that climate change presents to 17 people. This is sometimes called 'Ecosystem-based Adaptation (EbA) and includes:: 18

- Natural flood management: As warm air holds more water, and in some places, because of changing 20 seasonal rainfall patterns, we are seeing more heavy downpours in many parts of the world. This can create serious flooding problems, with loss of life, homes and livelihoods. The risk of flooding is higher where natural vegetation has been removed, wetlands drained or channels straightened. In these circumstances, water flows quicker and the risk of flood defenses being breached is increased. 24 Restoring the natural hydrology of upstream catchments, including by restoring vegetation, creating wetlands and re-naturalising watercourse channels and reinstating connections with the flood plain can reduce this risk. In a natural catchment with trees or other vegetation, water flows slowly overland and much of it soaks into the soil. When the water reaches a watercourse, it moves slowly down the channel, both because of the longer distance it travels when the channel bends and because vegetation and fallen trees slow the flow. Wetlands, ponds and lakes can also hold water back and 30 slowly release it into river systems.
- Restoring natural coastal defences: Rising sea levels as a result of climate change, mean that coasts • 33 are eroding at a fast rate and storm surges are more likely to cause damaging coastal flooding. 34 Natural coastal vegetation, such as saltmarsh and mangrove swamps can, in the right places, stabilise 35 the shoreline and act as a buffer, absorbing the force of waves. On a natural coast, the shoreline will 36 move inland and as sea level rises, the coastal vegetation will gradually move inland with it. This 37 contrasts with hard coastal defences such as sea walls and banks, which can be overwhelmed and 38 fail. In many places however, coastal habitats have been cleared and where there are hard sea 39 defences behind the coastal zone, the vegetation disappears as the coast erodes rather than moving 40 inland. This is often referred to as 'coastal squeeze' as the vegetation is squeezed between the sea 41 and the sea wall. Restoring coastal habitats and removing hard sea defences, can help reduce the 42 risks of catastrophic flooding. 43
- *Providing local cooling:* Climate change is bringing higher temperatures globally, which can result • 45 in heatwaves affecting people's health, comfort and agriculture. In cities, this can be a particular 46 problem for health as temperatures are typically higher than in the countryside. Trees give shade, 47 which people, in both rural and urban areas, have long used to provide cool places for themselves 48 and for crops such as coffee and livestock. Planting trees in the right place can be a valuable, low-49 cost Natural-based Solution to reduce the effects of increasing heat, including in reducing water 50 temperatures in streams and rivers, which can help to maintain fisheries. Trees and other vegetation 51 also have a cooling effect as a result of water being lost from their leaves through evaporation and 52 transpiration (loss of water through pores in the leaves, known as stomata). Natural areas, parks, 53 gardens in urban areas can help reduce air temperatures by up to a few degrees. 54
- Restoring natural fire regimes: Some natural ecosystems are adapted to burning, such as savannas 56 • and boreal forests. Where fire has been suppressed or non-native species of trees planted in more 57

	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
1 2 3 4	fires. Solutions can includ decrease people and ecosy	e restoring natural fire regimes	tes, which can result in larger and hotter and removing non-native species to cerbated fire risk climate change is bringing g rainfall patterns.
5 6 7 8 9 10	important part in reducing greenho	buse gas emissions and taking c ge of other ways, including thro ion. There is increasing eviden	roves, forests and peatlands, also play an carbon dioxide out of the atmosphere. They ough providing food, materials and ace that spending time in natural
11 12 13 14 15		sions about their local environm	the right place and that local communities nent if Nature-based Solutions are to be lerge and benefits can be lost.
16 17 18 19	gases in the atmosphere, it is impo	rtant to note that there are limit ature, it will be essential to rad	e and reduce the amount of greenhouse ts to what they can do. To provide a safe lically reduce greenhouse gas emissions,
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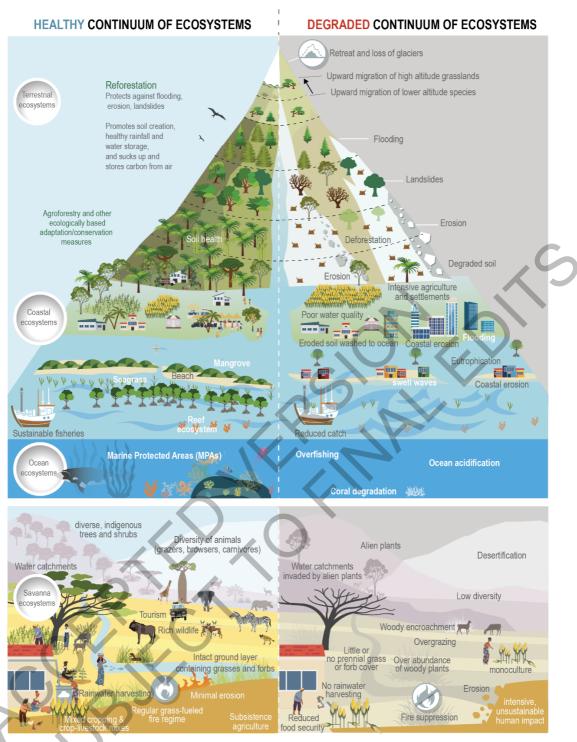


Figure FAQ2.5.1: Different Nature-based Solutions strategies.

[END FAQ2.5 HERE]

2.6.8 Feasibility of Adaptation Options

IPCC (2018a) defined feasibility as "the degree to which climate goals and response options are considered
 possible and/or desirable" (IPCC 2018) and set out an approach to assessing feasibility of pathways to limit
 global temperature rise to 1.5 °C. (Singh et al., 2020) have developed this approach for adaptation,
 recognising 6 different dimensions of feasibility: economic, technological institutional social cultural

recognising 6 different dimensions of feasibility: economic, technological, institutional, socio-cultural,

environmental / ecological and geophysical (Table 2.9). Feasibility is considered more fully in other chapters
 of this report, including Cross-Chapter Box FEASIB in Chapter 18. Adaptation for biodiversity conservation

and EbA encompasses a large range of approaches and techniques (Sections 2.6.2, 2.6.3) and will vary in 1 different contexts globally, as illustrated by the range of case studies (Section 2.6.5). It is important to take 2 account of specific regional and local circumstances as well as the type of adaptation action that envisaged 3 before making a feasibility assessment. It is also important to note that what is a feasible adaptation response 4 may change with the level of warming experienced – some techniques will be become less effective at higher 5 levels of warming. With global temperature rises of less than 2°C, in many cases it will be realistic to build 6 resilience and maintain species and ecosystems in situ, but at higher levels of warming, this will become 7 increasingly difficult and managing inevitable change, including the consequences of loss and damage will 8 be important (Prober et al., 2019). Similarly to be effective at higher levels of warming may require the 9 adaptation of EbA approaches themselves (Calliari et al., 2019; Martín et al., 2021; Ossola and Lin, 2021). 10 We have therefore not attempted a global scale assessment of the feasibility of adaptation options, but rather 11 present some key cross cutting considerations in assessing feasibility for adaptation for and through 12 ecosystems. 13

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Many of the necessary techniques for climate change adaptation for biodiversity and EbA have been demonstrated and shown to provide a wide range of additional benefits. This does however depend on deploying the right techniques in the right place (Box 2.2) and the engagement of local communities (see Section 2.6.6). There is also a challenge where there is high demand for land for other purposes, especially for agriculture and urban developments. Table 2.8 summarise the main feasibility considerations, drawing on previous sections. An assessment of constraints on EbA by Nalau et al. (2018) addressed similar issues.

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23 **Table 2.8**: Considerations in assessing feasibility of ecosystem restoration for climate change adaptation, following

24 Singh et al. (2020)

Singh et al. (2020)				
Feasibility characteristics	Feasibility indicators	Factors relevant to ecosystem restoration		
Economic	Micro-economic viability Macro-economic viability Socio-economic vulnerability reduction potential Employment & productivity enhancement potential	Costs are highly variable, depending on techniques and whether land purchase is required. Costs will depend on local rates for labour and materials. Economic benefits to local communities where employment is created and where loss from extreme events are avoided. (Section 2.6.4; De Groot et al., 2013)		
Technological	Technical resource availability Risks mitigation potential (stranded Assets, unforeseen Impacts)	Techniques are available for restoration of most ecosystems (Sections 2.6.2; 2.6.3) although it can be very difficult to achieve in some circumstances and take a long time, e.g. the restoration of peat swamp forests (Section 2.6.5.10). Successful implementation may also require skills which are in short supply and training may be required.		
Institutional	Political acceptability Legal, Regulatory feasibility Institutional capacity & Administrative feasibility Transparency & accountability potential	This will vary according to local factors. It should however be noted that EbA and adaptation for conservation has been implemented in wide range of different countries, including ones (see case studies in Section 2.6.5). In many cases EbA can meet multiple policy objectives but fall between different decision makers responsibilities.		
Socio-cultural	Social co-benefits (Health, education) Socio-Cultural acceptability Social & Regional Inclusiveness Benefits for gender equity Intergenerational equity	Multiple benefits to local communities are possible but full engagement and/or leadership of affected members of communities has been shown to be critical. Local Knowledge and Indigenous Knowledge can provide important insights. (Section 2.6.6)		
	Ecological capacity			

FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report
Environmental/ ecological	Adaptive capacity/potential	It is important to assess the benefits for ecosystems in relation to other potential options. In particular for some EbA approaches, it may be possible to achieve a range of different outcomes for biodiversity.
Geophysical	Physical Feasibility Land use change enhancement potential Hazard risk reduction potential	Appropriate measures need to be designed to take account of local geophysical conditions, for example catchment characteristics which define where some habitats can occur. This is also critical for ensuring the effectiveness of EbA in reducing natural bazards

A key element of economic feasibility is the cost of adaptation options. Costs of adaptation vary greatly 3 depending on the actions taken, the location, the methods used, the need for ongoing maintenance and 4 whether land purchase is necessary. At its simplest adaptation may be a matter of taking account of actual or 5 potential climate change impacts in the course of conservation planning and have little or no additional cost. 6 For example, if a species of conservation concern colonises or starts to use a new area as a result of climate 7 change (for example, migrant waterfowl shifting the locations where they overwinter (Pavón-Jordán et al., 8 2020b), protection or habitat management may be re-directed there. At the other extreme large scale 9 restoration can incur significant costs, for example between 1993 and 2015, the EU-LIFE nature programme 10 invested 167.6M € in 80 projects, which aim to restore over 913 km² of peatland habitats in Western 11 European countries (Andersen et al., 2017). This is equivalent to less than 2% of the remaining peatland 12 area, much of which has been affected to at least some extent by human pressures and restoring the total 13 areas will cost considerably more. De Groot et al. (2013) analysed 94 restoration projects globally and found 14 costs varied by several orders of magnitude but terrestrial and freshwater ecosystems mostly in the range of 15 100–10.000 USD per ha. They did however estimate that the majority of these projects provided net benefits 16 and should be considered as high yield investments. Some methods can however be much cheaper than 17 others even in the same type of ecosystems in the same Country: estimated cost of restoring forest cover in 18 Brazil varied between a mean of 49 USD using natural regeneration compared to a mean of 2041 USD per 19 hectare using planting (Brancalion et al., 2019). In assessing costs it is also important to take account of the 20 benefits delivered by different options, both in economic terms and other wider benefits. 21 22

The 'technological' dimension of feasibility in the context of ecosystems can be regarded as the range of 23 techniques available and the capacity to implement them. As described in Sections 2.6.2 and 2.6.3, above, a 24 wide range of techniques have been developed and are starting to be implemented. There is good evidence to 25 support adaptation for biodiversity and EbA in general terms and in many cases adaptation draws on 26 techniques for habitat creation and restoration which have been develop to mee other objectives. However, 27 feasibility needs to be assessed alongside likely effectiveness: a feasible but ineffective scheme is of no 28 value and the evaluation of success for specific interventions remains poorly developed (Morecroft et al., 29 2019). It is often therefore important to proceed with the use of pilot studies and good monitoring and 30 evaluation of outcomes to build confidence before wider deployment of approaches. A linked technical area 31 is the availability of specialist skills and knowledge to implement adaptation which can vary considerably 32 according to the type of adaptation measure. 33

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Institutional dimensions are dealt with more fully in other chapters, but in the specific context of the natural environment, it is notable that EbA is relevant to a wide range of organisations and policy objectives, in addition to environmental departments, NGOs and agencies, which conservation has traditionally been delivered by. Upscaling implementation is likely to be dependent on this wider range of interests. There can however be problems in that appropriate geographies for decision making on ecosystems (such as a catchment) may not directly map onto governance arrangements

Socio-cultural factors are important in adaptation of the natural environment, in reviewing constraints on
 EbA Nalau et al. (2018) found that risk perceptions and cultural preferences for particular types of
 management approaches were frequently identified in studies.

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In the IPCC feasibility assessment framework, one integral dimension is 'Environmental / ecological'. In this
 respect adaptation by and for ecosystems should perform well and this may be a reason to prefer EbA to

1 2	other approaches when there is an alternative. It should however be noted that sometimes apparently environmentally positive approaches such as forest creation can be done in ways which are damaging
3	(Section 2.6.7 and Box 2.2) and impacts need to be critically assessed for local circumstances.
4	
5	Geophysical dimensions are important for ecosystems as they have typically shaped which ecosystems can occur where and feasibility will depend on implementing adaptation options in places where they are
6	appropriate. Paleoecological studies can help inform potential options (Wingard et al., 2017)
7	appropriate. I alcoccological studies can help inform potential options (wingard et al., 2017)
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9 10	[START BOX 2.2 HERE]
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12	Box 2.2: Risks of Maladaptive Mitigation
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14	To hold global temperature rise to well below 2°C and pursue efforts to limit it to 1.5°C as required by the
15	Paris Agreement requires major changes in land use and management. There are many opportunities for
16	Nature-based Solutions, which can provide climate change mitigation and adaptation in ways, which protect
17	and restore biodiversity and provide a wide range of benefits to people (Cross Chapter Box NATURAL this
18	Chapter). There are also new technologies and approaches to develop the bioeconomy in ways, which
19	provide many benefits (Cross-Working Group Box BIOECONOMY in Chapter 5). Nevertheless, renewable
20	energy is a large and essential element of the climate change mitigation and there are adverse impacts on
21	biodiversity associated with some renewable energy, including wind and solar technologies (Rehbein et al.,
22	2020) However, of the most serious emerging conflicts are between land-based approaches to mitigation
23	and the protection of biodiversity, particularly as a result of afforestation strategies and potentially large
24	areas devoted to bioenergy, including bioenergy with carbon capture and storage (BECCS). It is important to
25	recognise the impacts of climate change mitigation at the same time as assessing the direct impacts of
26 27	climate change and ensure that adaptation and mitigation are joined up.
27 28	BECCS is an integral part of all widely accepted pathways to holding global temperature rise to 1.5°C
28 29	(IPCC, 2018b). This requires large areas of land which can conflict with the need to produce food and
30	protect biodiversity (Smith et al., 2018). One study that examined the combined impacts of climate change
31	and land use change for bioenergy and found severe impacts on species were likely if bioenergy were a
32	major component of climate change mitigation strategies (Hof et al., 2018). A study on the potential impacts
33	of bioenergy production and climate change on European birds found that land conversion for biodiversity to
34	meet a 2°C target would have greater impacts on species range loss than a global temperature increase of
35	4°C, if bioenergy were the only mitigation option (Meller et al., 2015). To avoid the worst impacts of
36	BECCS, it will need to be carefully targeted, according to context and local conditions (and other mitigation
37	strategies prioritised so its use can be minimised IPCC, 2019, Special Report on Land; Ohashi, 2019,
38	Biodiversity can benefit).
39	
40	Reforestation of formerly forested areas can bring multiple benefits, but planting trees in places where they
41	do not naturally grow can have serious environmental impacts, including potentially exacerbating the effects
42	of climate change. Savannas, are at amongst those that are at risk from afforestation programmes. Savannas are grass dominated, high diversity ecosystems with endemic species adapted to high light environments,
43 44	herbivory and fire (Staver et al., 2011; Murphy et al., 2016). Interactions between climate change, elevated
44 45	CO_2 and the disruption of natural disturbance regimes have led to widespread woody plant encroachment
т <i>э</i>	(Strang to 1 2010) and in the formation of the legitic state of the wide spread words plant encodermic the

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(Stevens et al., 2016) causing a fundamental shift in ecosystem structure and function with loss of grass, 46 reduced fire frequency (Archibald et al., 2009) and streamflow (Honda and Durigan, 2016). Afforestation 47

exacerbates this degradation (Bremer and Farley, 2010; Veldman et al., 2015; Abreu et al., 2017). Global-48 scale analyses aimed at identifying degraded forest areas suitable for afforestation (Veldman et al., 2019) 49

cannot reliably separate grassy ecosystems with sparse tree cover from degraded forests and local knowledge 50

is essential to ensure tree planting is targeted where it can most benefit and avoid harm. Figure Box 2.2.1 51 indicates where these issues are most likely to arise. 52

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Regions where savannas at potential risk from afforestation

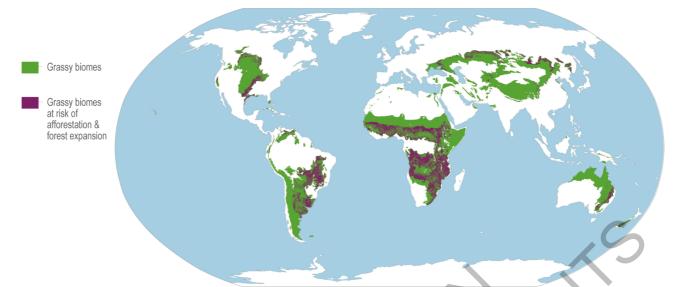


Figure Box 2.2.1: Regions where savannas at potential risk from afforestation. Based on (Veldman et al., 2015)

(Wilson et al., 2014). Mitigation benefits of growing timber are reduced or become negative in these

A similar issue can occur in naturally treeless peatlands which can be afforested if they are drained, but this

leads to the loss of distinctive peatland species and communities as well as high greenhouse gas emissions

conditions by CO₂ emissions from the oxidation of the drained peat - they can be a net sources rather than a

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[START FAQ2.6 HERE]

[END BOX 2.2 HERE]

16 FAQ2.6: Can tree planting tackle climate change?

sinks (Simola et al., 2012; Crump, 2017; Goldstein et al., 2020).

Restoring and preventing further loss of native forests is essential for combating climate change. Planting trees in historically unforested areas (grasslands, shrublands, savannas, some peatlands) can reduce biodiversity and increase the risks of damage from climate change. It is therefore essential to target tree planting to the appropriate locations and use appropriate species. Restoring and protecting forests reduces human vulnerability to climate change, reduces air pollution, stores carbon and builds natural systems resilience.

24 Like all living plants, trees remove carbon dioxide from the atmosphere through the process of 25 photosynthesis. In trees, this carbon uptake is relatively long-term, since much of it is stored in the trees' 26 woody stems and roots. Therefore, tree planting can be a valuable contribution to reducing climate change. 27 Besides capturing carbon, planting trees can reduce some negative impacts of climate change by providing 28 shade and cooling. It can also help prevent erosion and reduce flood risk by slowing water flow. Restoring 29 forest in degraded areas supports biodiversity and can provide benefits to people, ranging from timber to 30 food and recreation. 31 32

There are some areas where replacing lost trees is useful. These include forest that has been recently cut down, and where reforestation is usually practical. However, it is very important to identify correctly areas of forest that are degraded or have definitely been deforested. Reforesting places, especially where existing native forest patches occur, brings benefits both in sucking up carbon from the atmosphere and helping us adapt to climate change. Plantations of a non-native species, although offering economic benefits, do not usually provide the same range of positive impacts and generally have lower biodiversity and carbon uptake and storage.

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Reforestation options include the natural regeneration of the forest, assisted restoration, enrichment planting, 1 native tree plantations, commercial plantations and directed tree planting can occur in agroforestry systems 2 and urban areas. Reforestation with native species, usually contributes to a wide range of sustainability 3 goals, including biodiversity recovery, improved water filtration and groundwater recharge. It can reduce the 4 risks of soil erosion and flood risk. In cities, planting trees can support climate change adaptation by 5 reducing the heat of the area and can promote a wide range of social benefits such as providing shade and 6 benefiting outdoor recreation. Urban trees can also lower energy costs by reducing demand for conventional 7 sources of cooling like air-conditioning, especially during peak-demand periods. It is therefore important to 8 recognise that there are a wide range of different planting and forest management strategies. The choice will 9 depend on the objectives and the location. 10 11 Not everywhere is suitable for tree planting. It is particularly problematic in native non-forested ecosystems. 12 These natural ecosystems are not deforested and degraded but are instead naturally occurring non-forested 13

ecosystems. These areas vary from being open grasslands to densely wooded savannas and shrublands. Here 14 restoring the natural ecosystems instead of afforesting them will better contribute to increasing carbon 15 storage and increasing the areas resilience. It is important to remember, just because a tree can grow 16 somewhere, it does not mean that it should. These systems are very important in their own right, storing 17 carbon in soils, supporting a rich biodiversity and providing people with important ecosystem services such 18 as grazing. Planting trees in these areas destroys the ecosystem and threatens the biodiversity, which is 19 adapted to these environments. They can also impact on ecosystem services such as forage for livestock, on 20 which many people rely. 21

22 Many of these open areas also occur in low rainfall areas. Planting trees there uses a lot of water, can cause 23 reductions in stream flow and groundwater. Many of these locations also burn regularly and planting trees 24 threatens both the establishing trees but can also increase the intensity of the fires from that of a grass-fuelled 25 fire to that of a woody-fuelled fire. Swapping grassy ecosystems for forests may contribute to warming, as 26 forests absorb more incoming radiation (warmth) than grasslands. Aside from the negative impacts to 27 adaptation, it is also questionable just how much carbon can be sequestered in these landscapes asplanting 28 trees in grassy ecosystems can reduce carbon gains. Furthermore, a high belowground carbon store prevents 29 carbon loss to fire in these fire-prone environments. 30

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Another example is with peatlands. Peats store an incredible amount of carbon within them and are therefore important in maintaining and restoring to reduce atmospheric carbon. However, the restoration actions depend on what type of peatland it is and where it is located. Many temperate and boreal peatlands are naturally treeless. Here planting trees is often only possible following drainage, but draining and planting (especially with non-native species) destroys native biodiversity and releases greenhouse gases. Yet many peatlands, especially in the tropics, are naturally forested and restoring these peatlands requires re-wetting and restoring natural tree cover (see Figure FAQ2.2.1) which will increase carbon storage.

40 There are actions we can do instead of planting trees in non-forested ecosystems, and these include:

- Address the causes of deforestation, forest degradation and widespread ecosystem loss;
- Reduce carbon emissions from fossil fuels;
- Focus on ecosystem restoration over tree planting. For example, in restoring tropical grassy
 ecosystems, we can look at actions that cut down trees, enhance grass regrowth, and restore natural
 fire regimes. We have then a much better chance of both enhancing carbon capture and reducing
 some of the harmful effects of climate change.
- In between the two extremes of where planting trees is highly suitable and areas where it is not, it is important to remember that the context matters and decisions to (re)forest should look beyond simply the act of planting trees. We can consider **what the ecological, social and economic goals are of tree planting**. It is then important to verify the local context and then decide **what restoration action will be most effective**. It is also very important to conserve forests before worrying about reforesting.
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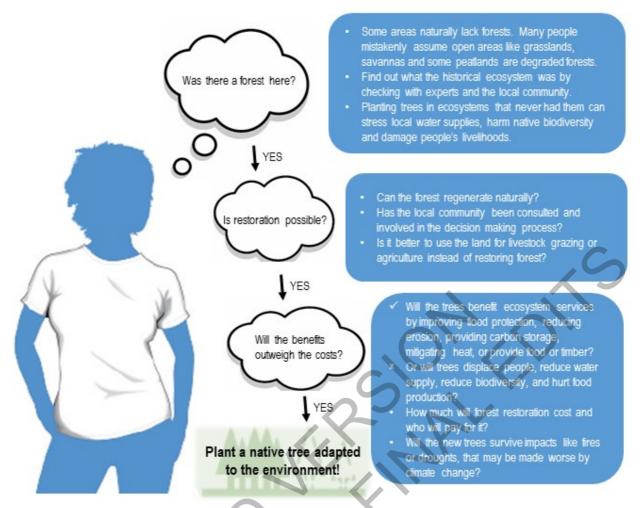


Figure FAQ2.6.1: Some places are more appropriate for tree planting than others and caution needs to be applied when planting in different biomes with some biomes being more suitable than others. This figure highlights some basic biome specific guidelines when planting in natural and semi-natural vegetation.

[END FAQ2.6 HERE]

2.7 Reducing Scientific Uncertainties to Inform Policy and Management Decisions

Research since the IPCC Fifth Assessment Report (Settele et al., 2014) has increased understanding of climate change impacts and vulnerability in ecosystems. Evidence gaps remain and geographic coverage of research is uneven. This section assesses gaps in ecosystem science where research is necessary for environmental policies and natural resource management, including under the UN Framework Convention on Climate Change and the Convention on Biological Diversity.

2.7.1 Observed impacts

Detection and attribution efforts have yet to give robust assessments of the roles of climate change in wildfires, tree mortality and human infectious diseases. Only one fire impact – the increase of the area burned by wildfire in western North America in the period 1984–2017 (Section 2.4.4.2.1) and just a few cases of tree mortality (Section 2.4.4.3) have been formally attributed to anthropogenic climate change. Global changes in soil and freshwater ecosystem carbon over time remain uncertainties in global carbon stocks and changes (Section 2.4.4.4); due to physical inabilities to conduct repeat monitoring and the lack of remote sensing to scale up point measurements, no global methods can yet produce repeating spatial estimates of soil carbon stock changes.

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Despite the growing understanding of the importance of ecosystem services this assessment found limited research on observed impacts of climate change on ecosystem services for 14 of 18 ecosystem services (Table 2.1)

2.7.2 Projected risks

6 A challenge for future projections that continues from previous IPCC reports is accurately characterising and 7 quantifying interactions among climate change and other factors causing ecological change, including 8 deforestation, agricultural expansion, urbanisation, and air and water pollution. Interactions can be 9 particularly complex for invasive species, pests and pathogens, and human infectious diseases. Modelling of 10 risks at the species level requires comprehensive databases of physiological, life-history, and functional traits 11 relevant to ecosystem resilience to climate change. Taxa that particularly lack this basis for model 12 projections include fungi and bacteria. For numerous plant and animal species, research on genotypic and 13 phenotypic diversity as a source of ecosystem resilience would inform projections of risk. 14 15

Soil plays a vital role in ecosystem function, is the habitat of a large number of species and is a large carbon 16 store which is currently a major source of greenhouse gas emissions, it is therefore a priority for climate 17 change research (Hashimoto et al., 2015). Major uncertainties remain in our understanding of soil functions. 18 Earth System Models (ESMs) predict soil respiration to increase with rising temperature (Friedlingstein et 19 al., 2014). However, there is evidence of acclimation post-increase (Carey et al., 2016) as the opposite 20 response of decrease in respiration with warming (Li et al., 2013; Reynolds et al., 2015). Long-term, large 21 scale field observations combined with a better conceptual understanding of factors governing soil process 22 responses to climate change is needed. A better understanding of plant water relations is also necessary, 23 including the response of plant transpiration to increased CO₂, climate warming and changes in soil moisture 24 and groundwater elevation. 25

27 2.7.3 Adaptation and Climate Resilient Development

28 There are significant evidence gaps in developing adaptation, both for biodiversity conservation and EbA. In 29 particular, whilst many adaptation measures have been proposed and implementation is starting, there are 30 very few evaluations of success in the scientific literature (Morecroft et al., 2019; Prober et al., 2019). As 31 detailed in Section 2.6.2, there is a strong literature on conceptual approaches to climate change adaptation 32 for biodiversity, but very little empirical testing of which approaches actual work best. Going forward it is 33 important put in place good monitoring and evaluation of adaptation strategies. For EbA, there are good 34 examples of measuring changes in response to new adaptation measures, but these remain relatively rare 35 globally. 36

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Human factors which promote or hinder adaptation are important as well as the technical issues. There are few studies incorporating climate change and ecosystem services in integrated decision making, and even fewer aimed to identify solutions robust to uncertainty (Runting et al., 2017).

41 Over the last decades, losses from natural disasters including those from events related to extreme weather

have strongly increased (Mechler and Bouwer, 2015). There is a need for better assessment of global

43 adaptation costs, funding and investment (Micale et al., 2018). Potential synergies between international

finance for disaster risk management and adaptation have not yet been fully realised. Research has almost exclusively focused on normalising losses for changes in exposure, yet not for vulnerability, a major gap

- given the dynamic nature of vulnerability (Mechler and Bouwer, 2015).
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References

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Chapter 2

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Chapter 2

Large Tables

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Table 2.2: Global Fingerprints of Climate Change Impacts across Wild Species. Updated from (Parmesan and Hanley, 2015). For each dataset, a response for an individual species

or functional group was classified as (1) no response (no significant change in the measured trait over time), (2) if a significant change was found, the response was classified as

either consistent or not consistent with expectations from local or regional climate trends. Percentages are approximate and estimated for the studies as a whole. Individual analyses

6 within the studies may differ. The specific metrics of climate change analysed for associations with biological change vary somewhat across studies, but most use changes in local or 7 regional temperatures (e.g. mean monthly T or mean annual T), with some using precipitation metrics (e.g. total annual rainfall). For example, a consistent response would be

poleward range shifts in areas that are warming. Probability (P) of getting the observed ratio of consistent: not consistent responses by chance was <10-13 for (Parmesan and Yohe,

9 10 2003; Root et al., 2003; Root et al., 2005; Poloczanska et al., 2013) and was <0.001 for Rosenzweig 2008 (source=publication) (Parmesan and Yohe, 2003; Root et al., 2003; Root et al., 2005; Rosenzweig et al., 2008; Poloczanska et al., 2013). Test were all binomial tests against p=0.5, performed by Parmesan.

N: total numbers of species, functional groups or studies	Species in given system: Terrestrial (T) Marine (M) Freshwater (F)	Types of change	Changes documented	Geographical region	Study allows for attribution to climate change
d Phenological changes					
677 species	461 plants, 168 birds, 35 insects, 9 amphibians, 2 fish	spring phenology	Overall: 9% delay; 27% no trend; 62% advance Mean change 2.3 days per decade advance	global	yes
agricultural crops, fruit trees, wild plants	Т	spring and fall phenology	From 1971-2000 48% responding as expected; spring advance 2.5 days per decade; Mean fall delay = 0.2 days per decade fruit ripening 2.4 days per decade advance; farmer's activities 0.4 days per decade advance	Europe	yes
200 species		spring phenology	Overall advance 2.8 days per decade 20 changes delay, 153 advance, 8 no change; Significantly more advance at higher latitudes	global	yes
55 studies (~100-200 species)	1T: 65% 1M: 13% 1F: 22%	various	90% of changes consistent with local/regional climate change	global	yes
726 taxa	M,T&F birds, moths, aphids, terrestrial plants, marine & FW phytoplankton	spring phenology	83.5 % of "trends" were advances; mean overall advance 3.9 days per decade. Terrestrial plants 93% advancing, mean 5.8 days per decade; FW plants 62% advancing, mean 2.3 days per decade secondary consumers advanced less than primary consumers & producers	UK	no
59 populations, 17 studies	amphibia	phenology	35% statistically significant change; mean advance 6.1 ± 1.65 days per decade; range 17.5 days delay to 41.9 days advance; 65% (n=47 populations) found	global	no
	species, functional groups or studies d Phenological changes 677 species agricultural crops, fruit trees, wild plants 200 species 55 studies (~100-200 species) 726 taxa	species, functional groups or studiessystem: Terrestrial (T) Marine (M) Freshwater (F)d Phenological changes677 species461 plants, 168 birds, 35 insects, 9 amphibians, 2 fishagricultural crops, fruit trees, wild plants200 species55 studies (~100-200 species)1T: 65% 1M: 13% 1F: 22%726 taxa59 populations, 17amphibia	species, functional groups or studiessystem: Terrestrial (T) Marine (M) Freshwater (F)d Phenological changes677 species461 plants, 168 birds, 35 insects, 9 amphibians, 2 fishagricultural crops, fruit trees, wild plantsTspring and fall phenology200 species\$pring phenology species)55 studies (~100-200 species)1T: 65% IF: 22%726 taxaM,T&F birds, marine & FW phytoplankton59 populations, 17amphibia	species, functional groups or studiessystem: Terrestrial (T) Marine (M) Freshwater (F)Overall: 9% delay; 27% no trend; 62% advance Mean change 2.3 days per decade advance amphibians, 2 fish677 species461 plants, 168 birds, 35 insects, 9 amphibians, 2 fishspring phenology phenologyOverall: 9% delay; 27% no trend; 62% advance Mean change 2.3 days per decade advanceagricultural crops, fruit trees, wild plantsTspring and fall phenologyFrom 1971-2000 48% responding as expected; spring advance 2.5 days per decade; Mean fall delay = 0.2 days per decade; famer's activities 0.4 days per decade advance; famer's activities 0.4 days per decade advance200 speciesspring phenologyOverall advance 2.8 days per decade 20 changes delay, 153 advance, 8 no change; Significantly more advance at higher latitudes55 studies (~100-200 species)1T: 65% IK: 13% IF: 22%various90% of changes consistent with local/regional climate change726 taxaM; T&F birds, marine & FW phytoplanktonspring phenology83.5 % of "trends" were advance; rean overall advancing, mean 2.3 days per decade 20 days per decade. Terrestrial plants 93% advancing, mean 2.3 days per decade secondary consumers advanced less than primary consumers & producers59 populations, 17 studiesamphibiaphenology35% statistically significant change; mean advance 6.1±1.65 days per decade; range 17.5 days delay to	species, functional groups or studiessystem: Terrestrial (T) Marine (M) Freshwater (F)regiond Phenological changes677 species461 plants, 168 birds, 35 insects, 9 amphibians, 2 fishspring phenologyOverall; 9% delay; 27% no trend; 62% advance Mean change 2.3 days per decade advanceglobalagricultural crops, fruit trees, wild plantsTspring and fall phenologyFrom 1971-2000 48% responding as expected; spring advance 2.5 days per decade (Mean fall delay = 0.2 days per decade fruit ripening 2.4 days per decade advance; farmer's activities 0.4 days per decade fruit ripening 2.4 days per

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				significant relationship between breeding phenology and temperature; higher latitudes advanced more.		
Gill et al. (2015)	64 studies	T: 100% trees	delay of autumn senescence	Delay averaged 0.33 days yr ⁻¹ and 1.20 days per degree warming; more delay at low latitudes across N-hemisphere; High latitude species driven more by photoperiod than low latitude species	global	no
Ficetola and Maiorano (2016)	n= 66 studies of temperature effects; 15 of precipitation	T/F 100% (amphibians)	phenology and abundances	Population dynamics driven by precipitation while breeding phenology driven by temperature	global	no
Halupka and Halupka (2017)	28 spp multi- brooded, 27 spp single-brooded, some spp. several pops	T 100% (birds)	phenology: Length of breeding season	Shows differences in sign of response between single & multi-broods & migrant vs resident; Season extended by 4 days per decade for multi-brooded, shortened by 2 days per decade for single-brooded; Multi: 26 species; 15 of 34 pops sig extended, none sig reduced	northern hemisphere	yes
Kharouba et al. (2018)	88 species in 54 pair- wise interactions	Not given	T: changes in relative phenologies of consumers and their resources	Asynchrony between consumer and resource has increased in some cases and decreased in others, with no significant trend; the prediction that asynchronies should be increasing in general is not supported.	global	no
Cohen et al. (2018)	127 studies	T 100% (animals)	phenological trends	81% of 127 studies of animals show phenological change in direction of earlier spring. Some studies were multi-species. Mean advance since 1950: 2.88 days per decade	Europe North America Australia Japan	no
Keogan et al. (2018)	145 populations, 209 time series	seabirds - terrestrial breeding sites	phenological trends	No change in breeding dates between 1952 and 2015	global	yes
Radchuk et al. (2019)	4835 studies, 1413 species	non-aquatic animals	phenological trends	Greatest phenological advancements in amphibians, followed by insects and birds in that order	global but most in N hemisphere	no
Piao et al. (2019)	Review	plants	spring and autumn phenologies	Rate of advance slowing down across northern hemisphere and reversed in parts of western N America in response to regional cooling since 1980s	Global	no
Menzel et al. (2020)	53 species in Germany, 37 in Austria, 21 in Switzerland (includes overlaps)	plants	spring and autumn phenologies	Long time series 1951–2018. Autumn leaf colouring: mean delay 0.36 days per decade. Spring phenology (leaf infolding) mean advance 0.24 days per decade. Summer phenology (fruit ripening)	Europe	yes

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				mean advance 0.26 days per decade. Growing season length mean increase 0.26 days yr ⁻¹ but farming season length decreased by 0.02 days yr ⁻¹		
2.2b. Observ	ed Changes in distributio	ons, abundances and le	ocal population extine	tions		
Parmesan and Yohe (2003)	920 species	T: 85.2% M: 13.5% F: 1.3%	Distibutions and abundances	50% of species (n=460/920) showed changes in distribution or abundances consistent with local or regional climate change	global	yes
Root et al. (2003)	n=926 species	T: 94% M: 5.4% F: 0.6%	Distibutions and abundances	52 % of species (n=483/926) showed changes in distribution or abundances consistent with local or regional climate change	global	yes
Rosenzweig et al. (2008)	n=18 studies	T: 65% M: 13% F: 22%	Distibutions and abundances	90% of studies showed changes in distribution or abundances consistent with local or regional climate change	global	yes
Pöyry et al. (2009)	48 species	T, butterflies	range shifts	From 1992-2004 37 ranges shifted poleward; 9 shifted equatorial; 2 no change. Non-threatened species expanded poleward by 84.5 km, threatened species showed no significant change (-2.1 km)	Finland	yes
Tingley et al. (2009)	53 species	T, birds	elevational range shifts	Resurvery (2003—2008) of historical elevational transects (1911—1929). 90.6% of species track their climate niche (temperature and/or precipitation) with regional climate change. Lower elevation species (mean range centroid = 916m) tracked only precipitation; high elevation species (mean range centroid = 1,944m) tracked only temperature; Species that tracked both T and P had mid-elevation range centroids (1,374—1,841)	California	yes
Chen et al. (2011)	24 taxonomic group x region combinations for latitude, 31 for elevation	T>264 M>10 F>34	range shifts: elevation and latitude	Mean upward elevation shift 11.0 m per decade Poleward shift 16.9 km per decade	pseudo-global	no
Grewe et al. (2013)	90 species	dragonflies	shifts of northern range boundaries	48 poleward shifts; 26 equatorial; 16 no change from 1988 to2006	Europe	no

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				southern lentic (standing water) species expanded 116 km polewards; southern lotic (running water) & all northern species stable.		
Mason et al. (2015)	21 animal groups, 1573 species	T: Birds, Lepidoptera F: Odonates	range shifts in 3 time periods	Northward shifts 23 km per decade (1966–1975) and 18 km per decade(1986–1995), with significant differences among taxa in rates of change	UK	yes
Gibson- Reinemer and Rahel (2015)	13 studies, 273 species: plants birds mammals marine inverts	T M	Range shifts in 2 or 3 areas for each species. shift measured as change of limit OR centroid	50% shifts of cold limit inconsistent with each other in within species despite similar warming. Percent species showing inconsistent shifts (including stable vs directional or different directions) = 47% plants,; 54% birds; 46% marine invert; 60% mammals. Large difference in magnitude of range shifts when in same direction (mean difference 8.8x);	global	no
Ficetola and Maiorano (2016)	n= 66 studies temperature effects; 15 precipitation	T/F 100% (amphibians)	phenology and abundances	Population dynamics driven by precipitation, breeding phenology driven by temperature	global	no
Scheffers et al. (2016)	n= 94 ecological processes	all	all possible types and levels of ecological change	82% of ecological processes affected by climate change	global	no
Wiens (2016)	976 species	all	population extinction rates near warm latitudinal and elevational range limits	47% of species suffered climate-related local extinctions: fish 59%; insects 56%; birds 44%; plants 39%; amphibians 37%; mammals 35%	global	yes
Bowler et al. (2017)	1,167 populations, 22 communities	T: 48% M: 61% F: 35%	abundance; population trends	Terrestrial species with warm temperature preference performed better than cool preferers; Freshwater and marine species: no effect of temperature preference on performance. 47% of species with significant abundance changes: 61%M, 48%T, 35% FW	Europe	yes
Pacifici et al. (2017)	873 mammals; 1272 birds	T 100% (birds and mammals)	multiple: range change, abundance, reproductive rate, survival, body mass	Estimated negative impacts (range contraction, reduced reproductive rates, or other measures of fitness estimates) for IUCN threatened species based on actual observed change in more common, related species. 47% threatened mammals and 23% birds	global for bird; mammals N America	unclear - complex methods

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				negatively impacted by climate change in part of their ranges		
MacLean and Beissinger (2017)	21 studies 26 assemblages of taxonomically related species	plants and animals	Range shifts in latitude and altitude related to species' traits: dispersal, body size, habitat, diet specialization & historic range limit.	High latitude/altitude range boundaries shifted less than lower laatitude/altitude boundaries. Author explanation is that habitat limits being reached (eg mountain top). Magnitudes of shifts positively related to dispersal traits and habitat breadth.	global	no
Ralston et al. (2017)	46 species	birds	Shifts in climate niche breadth, filling of climate space and overall abundance.	Species increasing in abundance were also increasing breeding climate niche breadth and niche filling. Declining species were opposite: niche breadths narrowing and greater climate debt.	north America	no
Rumpf et al. (2019)	1,026 species	terrestrial plants, invertebrates, vertebrates	Comparison of rates of range limit shift at leading & trailing elevational edges.	No difference in mean rate of shift of leading and trailing edges; elevational range sizes not changing systematically. Greater lags in regions with faster warming.	global	no
Freeman et al. (2018)	975 species, 32 elevational gradients	terrestrial plants, endotherms, ectotherms	Comparison of rates of range limit shift at leading & trailing elevational edges.	Mean change at warm limit 92±455 m per °C; cool limit 131±465 m per °C; (± SD, not significantly different from each other. Available area and range sizes decreased for mountaintop species.	global	no
Anderegg et al. (2019b)	meta-analysis 50 studies >100 species	T: 100% woody plants	mortality at dry range edges	100 indiv. species + a community of 828 species mortality at range edges due to drought was 33% greater than for core populations	apparently global	yes; drought not warming
Román- Palacios and Wiens (2020)	10 studies, 538 species, 581 sites	terrestrial plants and animals	analysis for drivers of population extinctions at warm range edges	44% of species had suffered local population extinctions near warm range limits. In temperate regions sites with local extinction had greater increases in maixmum temperature than those without (0.456° C vs. 0.153° C, P < 0.001 , n = 505 sites) and smaller increases in mean temperatures (0.412° C vs 1.231° C, P < 0.001). In tropical regions, range edges with local extinction also had greater increases in maximum temperatures (0.316° C vs 0.061° C P < 0.001 , n = 76) but changes	global	yes

in mean temperatures were similar between edges with and without extinctions (0.415°C vs 0.406°C, P	FINAL DRAFT	Chapter 2	IPCC WGII Sixth Assessment Report	
= 0.9			in mean temperatures were similar between edges with and without extinctions (0.415°C vs 0.406°C, P = 0.9	
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